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College of Soil Physics

22 October - 9 November, 2007

Factors affecting the water erosion process.

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WATER EROSION: FACTORS

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Sheet erosion (splash erosion and overland flow)



Transport due to wind-driven rain



Transport due to wind-driven rain



u_* and KE_z or M_z

Sheet erosion



Interrill erosion



Rill erosion



Gully erosion



Natural factors influencing soil erosion (by water)

- rainfall erosivity
 - intensity
- soil erodibility
 - particle size (silt and fine sand)
 - clay content
 - organic matter content

aggregate stability

- (embedded) stones
- topography
 - slope steepness
- vegetation
 - surface cover
 - → reduce rainfall erosivity
 - canopy cover
 - roots → reduce soil erodibility



Erosion by water is a natural process ...

...but can be influenced by human practices

Human factors

- harvest of vegetation
 - loss of soil cover
- soil tillage
 - decrease in organic matter and biological activity
 - compaction of subsurface layers
- surface sealing (e.g. roads, pavement)
 - increase of overland flow



Soil loss tolerance value:

depends on the on and off site effects of soil erosion

- *on site*: depth of fertile soil layer
 - \rightarrow values proposed by the Soil Conservation Service (Logan, 1982)

soil depth (cm)	soil loss tolerance (t/ha yr)
< 25 till solid rock	2.2
< 25 till sand	4.5
25 to 50 till solid rock	4.5
25 to 50 till sand	6.7

- off site:
 - water quality (e.g. eutrophication risk)
 - silting up of reservoirs and ditches
 - risk of flooding

What are the consequences?

- On site
 - decrease of soil quality
 - filling up of gullies
 - loss of yield, nutrients, pesticides
- Off site
 - mudflows
 - silting up of reservoirs
 - deterioration of water quality

Loss of soil quality



Formation of gullies



Loss of yield



<u>Mudflows</u>



Silting up reservoirs



Deterioration of water quality (eutrophication)



How can we measure it?

- Laboratory:
 - soil pans
 - rainfall simulations
 - flume experiments
- Field
 - erosion plots
 - rainfall simulations
 - flume experiments
- Watershed
 - deposition in reservoirs
 - sediment discharge at the outlet

• Laboratory rainfall simulation



• Field rainfall simulation



Field plot

 (flume with
 pressure
 transducer +
 collector drum)



• Watershed

(ultrasonic sensor to determine the water level)



 Watershed (sampling device)



 Watershed (turbidity sensor)



Universal Soil Loss Equation – USLE (1978) Revised USLE – RUSLE (1997)

$A = R \times K \times LS \times C \times P$

- A = soil loss (sediment yield) per unit area (Mg ha⁻¹ year⁻¹)
- R = rainfall-runoff erosivity factor (MJ mm ha⁻¹ h⁻¹ year⁻¹)
- K = soil erodibility factor (Mg ha h⁻¹ MJ⁻¹ mm⁻¹)
- LS = slope-length factor
- C = cover-management factor
- P = supporting practices factor

R – Rainfall-runoff erosivity factor

- Potential of rain to detach and transport sediment
- Rainfall is characterised by
 - amount of rainfall (precipitation)
 - duration of rain storm
 - intensity of rainfall
 - momentum or kinetic energy of raindrops
- Wischmeier and Smith (1958): $R = E I_{30}$

Kinetic energy E

• <u>KE of single raindrop</u>

$$E_i = \frac{1}{2} m_i v_i^2$$

• <u>KE of rainstorm</u>

$$E = \sum E_i$$

impossible \rightarrow group raindrops into classes of equal size (and hence m and v)

- <u>Determine raindrop-size distribution</u>
- stain technique: paper + appropriate dye
- photographic methods
- immersion method: heavy grade oil
- flour pellet: flour on plate
- impact assessment: tension balance, acoustic device, pressure transducer (distrometer)
- optical spectro-pluviometer: laser technique

Drop-size distribution depends largely on <u>intensity</u> (!!!), origin of raindrop, type of cloud – max. 7 mm

For each drop size (class): calculate $m = 1/6 \pi d^3 (\rho_w - \rho_a)$

determination of v: from Laws (1949)

heights of fall (m)



 \rightarrow m + v \rightarrow E

very labor intensive

- "Direct" measurement of KE (needs calibration)
- Ellison splash cups (1947)
- Sensit KE meter: piezo-electrical crystals





still too labor intensive for practical purposes

• Indirectly from intensity: Renard et al. (1991)

KE of pluviophase j

$$E_{pj} = 0.29 [1 - 0.72 exp(-0.05 I_j)] \Delta P_j$$

$$E_{pj} = KE of pluviophase j (MJ ha-1)
 $I_j = rainfall intensity of pluviophase j (mm h-1)
\Delta P_i = depth of rainfall of pluviophase j (mm)$$$

KE of rainstorm

$$E = \Sigma E_{pj}$$

R – Rainfall-runoff erosivity factor

- Potential of rain to detach and transport sediment
- Rainfall is characterised by
 - amount of rainfall (precipitation)
 - duration of rain storm
 - intensity of rainfall
 - momentum or kinetic energy of raindrops
- Wischmeier and Smith (1958): $R = E I_{30}$

 I_{30} = maximum rainfall intensity in a 30 minute period within the rainstorm (mm h⁻¹)

- pluviograph
- tipping bucket pluviometer (rain gauge)


time

R – Rainfall-runoff erosivity factor

- Potential of rain to detach and transport sediment
- Rainfall is characterised by
 - amount of rainfall (precipitation)
 - duration of rain storm
 - intensity of rainfall
 - momentum or kinetic energy of raindrops
- Wischmeier and Smith (1958): $R = E I_{30}$

• Calculation of mean annual rainfall erosivity method 1: R = 0.5 P with P = mean annual rainfall remark: outcome is in american units \rightarrow R' = 17.3 x R (Mg mm ha⁻¹ h⁻¹) method 2 (for KE > 25): E = 9.28 P - 8.838 R = El₃₀/100

• Exercise: Calculate R if the mean annual precipitation P equals 2695 mm.

method 1: $R = 0.5 \times P$

method 2 (for KE > 25): Use $I_{30} = 75 \text{ mm h}^{-1}$ E = 9.28 x P - 8.838

• Exercise: Calculate R if the mean annual precipitation P equals 2695 mm.

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method 1: R = 0.5 x 2695
                = 1347.5
                = 1347.5 \times 17.3
                = 23 311.8
method 2 (for KE > 25): Use I_{30} = 75 mm h<sup>-1</sup>
              E = 9.28 \times 2695 - 8.838
                  = 16 171 J m<sup>-2</sup> = 16 171.6 Mg ha<sup>-1</sup>
              R = (16 \ 171.6 \ x \ 75)/100
                 = 12 128.7
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Monthly rainfall Pm and rainfall erosivity R at Ukkel (Belgium, period 1967-1993)



problem: many weather stations do not have pluviograph or tipping bucket pluviometer

e.g. Belgium (Michiels and Gabriels, 1993)

 $R = 11.55 \exp(0.0254 \text{ MFI})$

MFI = Modified Fournier Index

$$=\sum_{i=1}^{i=12} \frac{p_i^2}{P}$$

 p_i = mean monthly rainfall in month i P = mean annual rainfall e.g. Belgium (Michiels and Gabriels, 1993)

 $R = 11.54 \exp(0.00215 P)$

e.g. Ethiopia (Hurni, 1985)

R = -8.12 + 0.562 P

K – Soil erodibility factor

- Vulnerability or susceptibility of soil against erosion
- Erodibility depends on
 - soil texture (silt + very fine sand fraction)
 - organic matter content
 - soil structure
 - permeability hydraulic conductivity
 - initial soil-water content
- K = soil loss (sediment yield) per unit of R on a standard plot (22.1 m long; 9% slope), continuous fallow for 2 years, tilled up and down the hill

$$A = R \times K \times LS \times C \times P$$

$$\rightarrow$$
 K = A/R LS = C = P = 1

• from erosion plots



• in practice (RUSLE):

(1) via "nomogram"

 $K = 0.1317 [2.1 \ 10^{-6} (12 - OM) M^{1.14} + 0.0325 (S - 2) + 0.025 (P - 3)]$

- M = (% of 2-100 μ m fraction) x (100 % of < 2 μ m fraction)
- OM = % of organic matter
- S = structural class
- P = permeability class

valid for % of 2-50 μ m fraction < 70%



Solution: K = 0.041.

• Exercise:

70 % silt + fine sand, 20 % coarse sand, OC = 1 %, medium granular structure, slow to moderate permeability

K-Value = 0.62



(2) via mean weighted diameter Dg (Declercq and Poesen, 1992)

 $K = 0.0035 + 0.0388 \exp[-0.5((\log(Dg)+1.519)/0.7584)^{2}]$

 $Dg = exp[0.01 \Sigma(f_i \ln(m_i))]$

Dg = mean weighted diameter (mm) f_i = fraction of particles belong to diameter class i (mass%) m_i = mean diameter of diameter class i (mm)

USLE: considers K is constant over season

RUSLE: considers seasonal variation (structure, antecedent soil-water content, frost, ...)

by weighing instantaneous estimate of K in proportion to EI (the percent of annual R) for 15-day intervals

- soil structure
 - organic matter
 - soil conditioners
 - \Rightarrow increase of aggregate stability against raindrop impact
 - \Rightarrow more infiltration
 - breaking of compacted ploughsole
 - controlled traffic

 aggregate stability at different organic carbon contents (clay content = 19 %)



 infiltration with and without soil sealing (resp. low and high aggregate stability)



time (min)

- soil compaction:











- LS Slope-length factor
 - expresses the effect of topography



USLE:

LS = $(\lambda/22.1)^{m}$ (65.41 sin² θ + 4.56 sin θ + 0.065)

- λ = slope length (m)
- θ = angle of slope
- m = 0.5 if percent slop is < 5%, 0.4 on slopes of 3.5-4.5%, 0.3 on slopes of 1-3% and 0.2 on slopes < 1%

RUSLE: uses four separate slope length relationships

• Value obtained from nomograph



- Non-uniform slopes:
 - Slope is divided in *n* segments with uniform slope
 - For each segment: topographic factor calculated
 - Total topographic factor:

$$\begin{split} L \cdot S &= \left(1/L_n \right) \cdot \sum_{i=1}^{i=n} \left(L_i / 22.13 \right)^{m_i} \cdot S_i \cdot (L_i) \cdot \left(1 - \left(L_{i-1} / L_i \right)^{m_i+1} \right) \\ \text{with} \quad & (L_i / 22, 13)^{\text{mi}} \cdot s_i = \text{topographic factor} \\ & (L_i) (1 - (L_{i-1} / L_i)^{\text{Mi+1}} = \text{weight factor} \end{split}$$

• After re-arranging:

$$L \cdot S = (1/L_n) \cdot \sum_{i=1}^{i=n} S_i \cdot (L_i^{m_i+1} - L_{i-1}^{m_i+1}) (22.13^{m_i})$$
$$S_i = 0.0065s_i^2 + 0.0454s_i + 0.065$$

with $L_n = total slope length$

- L_i = length from top of slope to bottom of segment i
- I_i = length of segment i
- S_i = slope factor of segment i
- s_i = slope of segment i (%)
- m_i = slope length exponent

• Exercise:

Calculate LS for a uniform slope of 12 % and contour bunds at 20 m spacing.

Formula:

$$LS = \left(\frac{x}{22.13}\right)^n \left(0.065 + 0.045s + 0.0065s^2\right)$$

• Exercise:

Calculate LS for a uniform slope of 12 % (= 7°) and contour bunds at 20 m spacing.

$$LS = \left(\frac{x}{22.13}\right)^n \left(0.065 + 0.045s + 0.0065s^2\right)$$

LS = 1.46



C – Crop-management factor

- expresses degree of soil loss compared to cleantilled fallow
- varies from near 0 (very well protected soil) till 1.5 (finely, tilled fallow surface) for clean-tilled fallow surface: C = 1

Influence of soil cover on sediment concentration (Rose, 1994)



- types of soil cover

- protection of soil by potato, sugar beets and maize is rather limited due to the low amounts of crop residues and stem flow (maize)
- sowing of grass under maize, double drilling, green manure
- mulching

•USLE: distinction between

-forest (e.g. C = 0.001 if 90% of soil surface covered by humus); values from tables

-<u>pasture</u> (e.g. C = 0.003 for grass with 95% of cover); values from tables

-<u>cropland</u>: different growth stages are considered; each with different % of surface cover and multiplied with the monthly fraction of annual R e.g. development stage is June: $C_{June} = 0.5 \times R_{June}$

mulch





Mulching: crop residues corn
- Field plot measurements (Nukerke, Belgium):
 - 1 March -3 April 2001: 91 mm rain
 - 98 % less soil loss by green manure (rye grass, 60 % cover)



 84 to 99 % less total phosphorus losses by green manuring



79 to 98 % less nitrate losses (NO₃-N + NO₂-N) and 64 to 94 % less ammonium losses (NH₄-N)



• RUSLE:

$C = PLU \times CC \times SC \times SR$

- PLU = prior land use
- CC = crop canopy
- SC = surface cover
- SR = surface roughness

C is calculated and weighted at 15-day intervals

P – Supporting practice factor

- expresses degree of soil loss compared to up-anddown-hill culture
- max. P-value = 1
- values for contouring, strip cropping, terracing, conservation practices, ... from tables;

e.g. contouring on 9 to 12 % slope: P = 0.60

- soil tillage
 - to retain crop residues at the surface
 - to increase soil biological activity
 - tillage practices:
 - conventional tillage
 - no tillage
 - reduced tillage
 - subsoiling
 - contouring



No tillage



Subsoiling

subsoiling





Difference in soil moisture content after conventional (left) and reduced tillage (right)

- grass buffer strips
 - do not prevent soil erosion on site but reduce off site effects
 - slow down runoff
 - \Rightarrow deposition of sediment







– Measurements on field plots (Nukerke, Belgium):

- 12 to 16 February 2001: 26,7 mm rainfall
- 94 % less soil loss
- 93 % less runoff



– Measurements on field plots (Nukerke, Belgium):







Small and large scale hydraulic structures

 terraces, dikes, reservoirs and retention ponds

 \rightarrow to protect villages, channels, ...



retention ponds









