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Targeted Training Activity: Seasonal Predictability in Tropical Regions to be followed by Workshop on Multi-scale Predictions of the Asian and African Summer Monsoon

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Monsoon ISV: modelling issues and influence of global warming

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Modelling ISV and the Impacts of Climate Change

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Outline



- Role of the ocean
 - Thermodynamic
 - Dynamic
- Influence of model resolution
- Organised convection and the coupling between physics and dynamics
- Monsoon ISV and climate change



Air-sea interaction and the MJO







Increased power on intraseasonal timescales from high resolution SSTs





Some background – observed air-sea interaction





Some background – modelled air-sea interaction







CGCM has a propagating convective signal compared with standing oscillation in AGCM. Coherent variations in SST in CGCM in agreement with observations, *provided* the mean state (i.e. eastward edge of Warm Pool) is realistic

Inness and Slingo J. Clim. 2003 Walker Inness et al. J. Clim. 2003

Mixed layer model studies of the diurnal cycle: Sensitivity to vertical resolution



 1m resolution (CTR) gives good simulation of diurnal and intraseasonal variability

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- 10m resolution of most ocean models will not resolve diurnal variability of SST
- Intraseasonal variability is ~0.4°C less than CTR
- Implies 40% underestimate of the strength of air-sea coupling



Bernie et al. J. Clim. 2005



- N48L30 HadAM3 coupled to L300 ORCA2 with 1 meter resolution in near surface ocean
- Two 50-year simulations: 3-hourly coupling vs. daily coupling
- 3-hourly coupling gives improved tropical Pacific mean state, ocean currents and seasonal cycle



Amplitude of SST diurnal cycle in HadOPA (L300)







Bernie et al. Clim. Dyn. 2008

Lag-lead composites of MJO precipitation for various base points across the Indo-Pacific Warm Pool 3-hourly Daily HDC : Composite Lag corrolation of cyrain at 60.0000E HDM : Composite Lag corrolation of cyrain at 60.0000E 20 20 Lag (days) Lag (days) 0 -10-20 -20-30 -30 5 100 50 150 200 50 100 150 200 Longitude Longitude HDC : Composite Lag corrolation of cyrain at 90.0000E HDM : Composite Lag corrolation of cyrain at 90.0000E 30 20 10 Lag (days) Lag (days) -10-20 -20-30 -30 50 100 150 200 150 50 100 200 Longitude Longitude HDC : Composite Lag corrolation of cyrain at 120.000E HDM : Composite Lag corrolation of cyrain at 120.000E 30 E 20 -20 10 Lag (days) Lag (days) 0 0 -10 -20 -20-30 E -30 150 Longitude 50 100 150 200 50 100 200 Longitude HDC : Composite Lag corrolation of cyrain at 150.000E HDM : Composite Lag corrolation of cyrain at 150.000E 30 20 20 10 Lag (days) Lag (days) 0 \bigcirc -10Ē -10-20 -20 -30 -30 50 100 150 200 50 100 150 200 Longitude Longitude



Saith and Slingo IJC 2006





July 2002 Monsoon **Break:**

Evolution of surface winds and sea surface height

- Westerly wind event associated with the passage of the MJO produces downwelling ocean Kelvin waves and positive SSH anomalies in eastern Indian Ocean.
- Kelvin wave activity subsequently propagates north and south along coasts.

120°E

27

18







July 2002 Monsoon Break: Equatorial Indian Ocean Temperature Anomalies

• Build up of heat in Eastern Indian Ocean through oceanic response to westerly wind anomalies associated with the MJO.

• Accumulation of heat in Eastern Indian Ocean acts to prolong monsoon break conditions?

Krishnan et al. GRL 2006







Analysis of Active-Break cycles in 100-year HadCM3 and the role of the Indian Ocean

- Composite evolution of active and break cycles defined in terms of OLR dipole index (Vecchi)
- Observed quadrapole structure of convective anomalies is reasonably captured
- Model shows northward and eastward propagation of convective anomalies

Mohanty et al., in prep.







Analysis of Active-Break cycles in 100-year HadCM3 and the role of the Indian Ocean

- Composite evolution of SST anomalies associated with active and break cycles
- Active phases show coherent SST anomalies across the Indo-Pacific Warm Pool consistent with observations
- Break phases suggest that SST plays a less important role - at least in this model?

Mohanty et al., in prep.

140E

140E







Analysis of Active-Break cycles in 100-year HadCM3 and the role of the Indian Ocean

- Composite anomalies in the depth of the 20C isotherm for peak active (upper) and break (lower) phases
- Active phase suggests large dynamic adjustment of the equatorial ocean

Mohanty et al., in prep.



Active-Break cycles in 100-year HadCM3: Role of the Indian Ocean



Vertical profiles of the difference, Active minus Break, for the West and East Indian Ocean



Changes in mixed layer structure and dynamic adjustment of the ocean lead to differing responses in temperature and zonal currents



Mohanty et al., in prep.



- Increasing evidence that upper ocean variability plays a key role in defining the propagation characteristics and amplitude of ISV.
- In many regions this is dominated by the thermodynamic response of the upper ocean, but there may be an important role for fast dynamical adjustment, especially in the Indian Ocean during the boreal summer.
- That being the case, coupled models require a much improved representation of the upper ocean.
- Future plans include a detailed analysis of the upper ocean behaviour in HiGEM with 1/3 degree ocean





	HadCM3	HadGEM1	HiGEM1	NU-GEM	HiGEM2
	1998	2004	2006	2007/8	2009/10
Atmosphere	~300km	~150km	~90km	~60km	~40km
	19 levels	38 levels	38 levels	38 levels	60-70 levels
Ocean	1.25 [°] x 1.25 [°]	1º x 1º (1/3º)	1/3º x 1/3º	1/3º x 1/3º	1/4º x 1/4º
	20 levels	40 levels	40 levels	40 levels	50+ levels
Integration Length	1000s	100s	100s	10s	Seasonal to Decadal
	IPCC AR3	IPCC AR4			Prediction
Relative Computing	1	10	100	Earth Simulator	HECToR
Power					Met Office







Resolution improves annual mean biases











20-100 day band pass filtered variance of OLR (Oct-Apr)



Courtesy of Pete Inness

Convectively Coupled Equatorial Waves





Yang, Slingo and Hoskins, Submitted to J. Clim.



Convectively coupled equatorial waves

Wind/divergence solutions of theoretical equatorial wave equation on a b-plane, after Matsuno (1966).

Spatial scale assumes zonal wave no. 6 and meridional trapping (y_0) of 6^0 .

If winds represent lower tropospheric flow, then green/blue shading corresponds to regions of convection.

See Yang, Hoskins & Slingo, 2003, 2007a,b,c: J. Atmos. Sci.

WISHE working in equatorial waves?



Kelvin wave and T_b field regressed on equatorial T_b





Yang, Slingo and Hoskins, Submitted to J. Clim.





- Dynamical equatorial waves are simulated but the coupling of these waves with convection tends to be weak, especially in the vicinity of the equator.
- Model fails to capture the positive feedback between equatorial zonal winds and convection in the Kelvin (and R1) wave. Suggests lack of WISHE mechanism in model?
- Poor vertical structure of the simulated Kelvin wave indicates that the observed energy cycle, in which the upper tropospheric wave is amplified by convection, is missing in the models.
- The next step is to pin down those aspects of the moist physics that contribute to the poor dynamical coupling and energy cycles in the model.







Simulations of CCEW and ISV in the MRI/JMA 20km AGCM

- Equatorial Kelvin Wave captured better at high resolution. Other waves are still poor
- MJO is not improved despite better Kelvin Wave activity

From Rajendran et al., J. Clim. 2008













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Cascade: **Cloud System Resolving Modelling of the Tropical** Atmosphere

NCAS-Climate, Reading **Department of Meteorology, Reading Environmental Systems Science Centre, Reading** School of Environmental Sciences, UEA Institute of Atmospheric Science, Leeds Met Office, Exeter

Cascade Workplan

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- Develop modelling framework for large domain (up to 100⁰ longitude by 30⁰ latitude), high resolution (1-2 km) simulations.
- Perform numerical case studies of organised convection over West Africa and the Indo-Pacific Warm Pool.
- Evaluate these simulations using advanced satellite and in situ observations of cloud structures.
- Analyse the simulations in terms of scale-dependent energy and momentum budgets.
- Use idealised case studies to explore the links between convection and equatorial wave modes.
- Bring this new understanding into the development of new representations of tropical convection for global weather and climate models.



Monsoon ISV and Climate Change



- Hadley Centre coupled model HadCM3:
 - Atmosphere $3.75^{\circ} \times 2.5^{\circ} \times 30$ levels
 - Ocean 1.25° x 1.25° x 20 levels
- 100-year pre-industrial control integration
- 100-year (stable) 2xCO₂ integration



Composite evolution of active cycle of the ISM







Rainfall Anomalies (mm/day) during active and break phases

Turner & Slingo, QJRMS (Sub.) 2008



Absolute changes in rainfall during active and break phases

Monsoon breaks may become more severe – impacts on agriculture

Change in number of wet days



precipitation intensity (mm/day)

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Turner & Slingo, QJRMS (Sub.) 2008

Changes in the intensity of extreme Indian daily rainfall with climate change





Turner & Slingo, QJRMS (Sub.) 2008



- New results suggest that climate change may increase the severity of monsoon breaks and the intensity of extreme rainfall events. Both have serious implications for the vulnerability of Indian society
- More skillful simulations and more confident predictions of ISV are urgently needed.
- What's needed for to achieve this goal:
 - Improved representation of the upper ocean to capture the thermodynamic coupling on intraseasonal timescales
 - Better understanding and simulation of the dynamic adjustment of the ocean to ISV
 - Probably higher resolution in both the atmosphere and ocean, especially if it improves the mean state.
 - Better understanding and simulation of the coupling between physics and dynamics in organised tropical convection

