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9th Workshop on Three-Dimensional Modelling of Seismic Waves Generation, Propagation and their Inversion

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Ray Theory (Overheads)

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CLASSEAL RAY THEORY

U = ANGLE RAY MAKESLOCALLY WITH VERTICAL

$$
v d\theta = dr \tan i = dr \frac{Py}{y1 - p^2 c^2}
$$

\n
$$
\Rightarrow 0 = \int \frac{p_c}{p_c} (1 - p^2 c^2)^{-k} dr
$$

\n
$$
dt = \frac{dr}{\omega c} = \frac{qr}{c} (1 - p^2 c^2)^{-k}
$$

\n
$$
E = \int \frac{1}{c} (1 - p^2 c^2)^{-k} dr
$$

 \hat{q}

 $\frac{1}{2}$

WHAT IS THE ENERGY FLUX?

we need to find the Rate or modulus
of one is the of a name L' x-axis on
The 0TME.
Here
$$
\underline{n}
$$
 be a unit vector in x-direction
 $\underline{n} = (1, 0, 0)$
That of the x is the right of the x
where $dS = element of area$
Rate of working = ts in j is it: (force x velocity)

P-wave
Energy flux = Re $\{(\lambda + 2\mu)$ ik U $e^{i(\omega t - kx)}$
 $\times Re \{i\omega U e^{i(\omega t - kx)}\}$

$$
\Rightarrow \frac{Energy FUX AVERAGED OVER A CYCLE}{= \frac{1}{2} |U|^2 W^2 \rho \alpha}
$$

UNITS: ENERGY PER UNIT TIME PER UNIT AREA.

S-WAVE SIMILARLY

$$
\tau_{xy} = \mu ikVe^{i(\omega t - kx)}
$$

Energy flux = Re { μ ik Ve^{i(\omega t - kx)} }
×Re{ i\omega Ve^{i(\omega t - kx)}}

ENERGY FLUX AVERAGED OVER A CYCLE

 $=$ $\frac{1}{2}$ $|V|^2$ wk μ

 $=$ $\frac{1}{2}$ $|V|^2$ $\omega^2 \rho \beta$

ASYMPTOTIC THEDRY

1-DIMENSIONAL CASE

THE BASIC IDEH OF THE ASYMPTONC OR RAY THEORIES IS THAT IN MEDIA IN WHICH THE WAVE VELOCITIES AND DENSITY VARY SCOULY WAVES PROPAGATE IN MUCH THE SAME WAY AS IN HOMOGENEOUS MEDIA.

CONSIDER A P-WAVE PROPAGATING IN THE X-DIRECTION IN A MEDIUM IN WHICH DENSITY AND RUAVE SPEED ARE ALSO FUNCTONS OF X. WAVE EQUATION:

$$
\frac{\partial}{\partial x} (1+2\mu) \frac{\partial u}{\partial x} = \rho \frac{\partial^2 u}{\partial t^2}
$$
\n
\ni.e.
$$
\frac{\partial}{\partial x} (\rho a^2 \frac{\partial u}{\partial x}) = \rho \frac{\partial^2 u}{\partial t^2} \qquad \rho = \rho(x)
$$
\n
$$
SEEK AN APPROWMATE SOLUTION OF THE FORM\nFORM\n
$$
\frac{u(x,t) = U(x) e^{i\omega(t-\theta(x))}}{U=U(x)}, \qquad \theta = \theta(x) \qquad \text{To BE DERRMMED}
$$
\n
$$
\frac{\omega}{\omega} \text{ is consuperED TO BE A} \perp ARE PARMMED
$$
$$

SUBSTITU TING, WE FIND

$$
\frac{\partial u}{\partial x} = (-i\omega \frac{\partial \theta}{\partial x} U + \frac{\partial U}{\partial x}) e^{i\omega(t-\theta)}
$$
\n
$$
\frac{\partial}{\partial x} (P^{\alpha^2} \frac{\partial u}{\partial x}) = \{-\omega^2 P^{\alpha^2} U(\frac{\partial \theta}{\partial x})^2 - i\omega \frac{\partial \theta}{\partial x} \frac{\partial U}{\partial x} P^{\alpha^2}
$$
\n
$$
-i\omega \frac{\partial}{\partial x} (\frac{\partial \theta}{\partial x} U P^{\alpha^2}) + \cdots \frac{\partial}{\partial x} (e^{i\omega(t-\theta)})
$$

 \mathcal{S}^-

where ... indicates terms of lower order in W

Thus, FROM W²

$$
\left(\frac{\partial \theta}{\partial x}\right)^2 = \frac{1}{\alpha^2}
$$

 $\left(\frac{\partial \theta}{\partial x}\right)^2 = \frac{1}{\alpha^2}$
 $\left[\frac{\cos \theta}{\cos \theta}\right]$
 $\left[\frac{\cos \theta}{\cos \theta}\right]$

AND FROM W TERMS

$$
\frac{\partial \theta}{\partial x} \frac{\partial U}{\partial x} \rho x^2 + \frac{\partial}{\partial x} \left(\frac{\partial \theta}{\partial x} U \rho x^2 \right) = 0
$$

 $\frac{\partial}{\partial x} \left(\frac{\partial \theta}{\partial x} U^2 \rho \alpha^2 \right) = 0$
 $\frac{\partial}{\partial x} (U^2 \rho \alpha) = 0$ = constant energy

 $3-D$ THEORY - WORKS SIMILARLY (KARRL KKELLER, J. Aroust. Soc. Am., 21, 694, 1959) SEEK ASOLUTION OF EQUIS. OF MOTION IN FORM $u_i = U_i(x, y, z) e^{i\omega(t - \theta(x, y, z))}$ substitute into equation of motion, identify leading passers of ω (ω^2). DETAILS ARE COMBICATED.

b

WE FIND THAT EITHER

 $\theta_{ji} \theta_{ji} = \frac{1}{\alpha^2}$ ω *m* U_i \parallel θ , OR θ_{i} : θ_{i} : = $\frac{1}{3}$ w_{1} w_{i} \perp \mathcal{O}_{i} ;

THUS WE GET TWO KINDS OF SOLUTIONS CORRESPONDING TO P-WAVES AND TO SLWAVES

THUS, IM BOTHCASES WE OBTAIN FOR

THE "TRAVEL TIME" $S(x)$ an EQUATION

OF THE FORM

$$
(P\theta)^2 = \frac{1}{c^2}
$$
 EXAMPLE

WHERE C = & FOR P.WAVES, OR C = B FOR S-WAVES.

IMAGINE A PATH EVERYWHERE II TO DO

WE HAVE $\theta = \theta_0 + \int \frac{1}{c} ds$ HOW CAN WE DETERMINIE SUCH PATHS?

DIFFERENTIATING THE EIKONAL EON.

2
$$
\theta_{1i} \theta_{2i} = \frac{\partial}{\partial x_{j}} (t_{2})
$$

ie. 2 θ_{3} i θ_{3} ji = $\frac{2}{2x_{3}}\left(\frac{1}{5}t\right)$ but O₃: is parallel to Γ ie $\theta_{ij} = \frac{1}{c} \frac{dx_{i}}{ds}$ Ą \therefore 2 abi = 2. (20)

or
$$
dQ_i = \frac{c}{2} \frac{\partial}{\partial x_i} (\frac{1}{c^2}) = \frac{\partial}{\partial x_i} (\frac{1}{c})
$$

(B)

THUS, FROM (A) &(B)

RAY-TRACING **BUUATONS**

ALTERNAIVELY, WRITING

$$
\frac{d}{ds} = \frac{1}{c} \frac{d}{d\theta}
$$

$$
k_{ii} = \omega \theta_{ii}
$$

WE GET

$$
\frac{dx_i}{d\theta} = \frac{c^2}{\omega}k_i = c\frac{e^2}{k}
$$
\n
$$
\frac{dk_i}{d\theta} = \omega c \frac{\partial}{\partial x_i} \left(\frac{1}{c}\right) = -\frac{\omega}{c} \frac{\partial c}{\partial x_i} = -k \frac{\partial c}{\partial x_i}
$$
\n
$$
\therefore \frac{\omega n \epsilon \epsilon}{k} k = (k_i, k_i)^{k_i} \frac{\omega}{c}
$$
\n
$$
\frac{k_i}{k} = -k \frac{\partial c}{\partial x_i}
$$

THESE REPRESENT THE MOTION OF A "PARTICLE" TRAVELLING AT THE LOCAL WAVE SPEED C, SUFFERING DEFLECTIONS FROM A STRAIGHT-LINE TRAJECTURY DUE TO VELOCITY GRADIENTS THAT ARE NOT II TO THE PATH

 $R_i = \frac{\partial \Psi}{\partial x_i}$

THE LOCAL DISPERSION RELATION IS THEN GIVEN BY A FUNCTION W (Ri, Xi), so THE PHASE Y (X) SANSFIES AN EQUATION OF THE FORM $\omega = \omega(\frac{\partial y}{\partial x_i}, x_i)$

THE METHOD OF CHARACTERISTICS (ESSENTIALLY THE METHOD GIVEN ABOVE) THEN LEADS TO HAMILTON'S EQUATION

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HAMILTON'S EQUATIONS

GIVEN A LOCAL DISPERSION RELATION

 ω = ω (k, z)

THE RAY EQUATIONS ARE

$$
\vec{x}_i = \frac{\partial \omega}{\partial k_i}
$$

$$
\vec{k}_i = -\frac{\partial \omega}{\partial x_i}
$$

C.f. HAMILTON'S EQNS. FOR A MECHANICAL SYSTEM: GIVEN THE HAMILTONIAN $H(P_i, q_i)$ THE EVOLUTION OF THE SYSTEM $q_i = \frac{\partial H}{\partial p_i}$
 $\dot{p}_i =$ IS GOVERNED BY

$$
\frac{3}{2}c = -\frac{3H}{2Q}
$$

- Q .: "GENERALISED COORDINATES
- Pi = "GENBRALISED MOMENTA-"

LET US USE THIS IDEA TO RE-DERIVE THE RAY EQUATIONS. THE LOCAL DISPERSION RELATION IS OF THE SIMPLE FORM

 $w = c(z)|k|$

FOR BODY WAVES IN AN ISOTROPIC MEDIUM $(c : \alpha \quad \text{or} \quad c = \beta)$ ω = c(x) $(k_i k_j)^2$ le l

SVIE ENOTALIZATIONS CIVE

$$
\dot{x}_i = c \underline{k_i}
$$

\n $\dot{\overline{k}}_i = -\overline{k} \frac{\partial c}{\partial x_i}$
\n $\dot{\overline{k}}_i = -\overline{k} \frac{\partial c}{\partial x_i}$
\n $\overline{\overline{k}}_i$
\n $\overline{\overline{k}}_i$

LET US WRITE DOWN RAY EQUATIONS FOR AN ANISOTROPIC MEDWM. WE MAVE $(c_{ijkl} u_{kl})$, $+ \omega^2 u_i = 0$ => -ikj cijke (-ikg) un tw² ui=0 ie (Cijke kekj - w² Sir) u_{k} = 0 THUS THE LOCAL DISPERSION RELATION 15 Det (Cijkekek; -w² Sik) =0 THE DERIVATIVES DW OW CAN BE FOUND FROM STANDARD PERNRBATOON THEORY (RAYLEIGH'S PRANCIPLE) WE FIND $\dot{R}_{m} = \frac{\partial \omega}{\partial x_{m}} = \frac{1}{2\omega}$ $\frac{\partial C_{ij}}{\partial x_{m}}$ $v_{i}v_{k}k_{k}k_{j}$

 $\dot{x}_{m} = \frac{\partial \omega}{\partial k_{m}} = \frac{1}{2\omega} (C_{ij}k_{m}k_{j} + C_{imkk}k_{l})\sigma_{i}\sigma_{k}$

where v_i is a (local) unit eigenvector (CORRESPONDING TO THE WAVE OF INTEREST)

ANOTHER ELEGANT PROPERTY OF HAMILTON'S EQUATIONS IS THAT THEY CAN BE WRITTEN DOWN IN ANY COORDINATE SYSTEM

SUPPOSE THAT WE WATST TO DO 3-D RAYTRACING IN SPHERICAL CUURDINATES

 $R_v = \frac{\partial v}{\partial v}$, $R_\theta = \frac{\partial v}{\partial \theta}$, $R_\phi = \frac{\partial v}{\partial \theta}$ have We $k = (k_1^2 + \frac{1}{2}k_0^2 + \frac{1}{r^2m^2}k_0^2)^2$ and

 (r, θ, ϕ)

with the usued disportion relation

 $\omega = c(r, \theta, \phi)k$

WE OBTAIN RAY-TRACING EQUATIONS:

$$
\dot{\tau} = \frac{k}{k}c
$$
\n
$$
\dot{\theta} = \frac{1}{r^2} \frac{kg}{k}c
$$
\n
$$
\dot{\phi} = \frac{1}{r^2} \frac{kg}{k}c
$$

ILB

$$
\frac{k_{r}}{2^{r}} = -\frac{2c}{2^{r}}k + \frac{1}{kr}(\frac{1}{r^{2}}k_{0}^{2} + \frac{1}{r^{2}sin\theta}k_{1}^{2})
$$
\n
$$
\frac{k_{0}}{2^{r}} = -\frac{2c}{20}k + \frac{cot\theta}{kr^{2}sin^{2}\theta}k_{1}^{2}
$$
\n
$$
\frac{k_{0}}{2^{r}} = -\frac{2c}{20}\frac{1}{\theta}
$$

TO MAKE CONTACT WITH CLASSICAL RAY THEORY IN THE SPHERICAL EARTH LET US NOW SIMPLIFY THESE FOR THE CASE COCU) TAKE SOURCE AT O=0, R&EO

$$
\vec{r} = k_y c
$$
\n
$$
\vec{r} = k_y c
$$
\n
$$
\vec{r} = \frac{1}{6}c
$$
\n
$$
\vec{r} = \frac{1}{6}c
$$
\n
$$
\vec{r} = 0
$$
\n $$

THUS WE OBTAIN THE CLASSICAL RAY TATEGRALS

 $t = \int \frac{1}{2}(1-\frac{c^{2}p^{2}}{r^{2}})^{-\lambda} dr$
 $\theta(\frac{r}{r^{2}}) = \int \frac{pc}{r^{2}} (1-\frac{c^{2}p^{2}}{r^{2}})^{\frac{1}{2}} dr$

AMPLITUDES AND WAVEFORMS

BECAUSE RAY THEORY (FOR BODY WAVES) IS FREQUENCY-INDEPENDENT, IT PREDICTS THAT WAVES PROPAGATE WITHOUT ANY CHANGE TO THE WAVEFORM

(JUST AS IN A MOMOGENEOUS MEDIUM)

THE ASYMPROTIC THEORY CAN BE USED TO DERIVE WAVE AM PLITUDES (BY INVESTIGATING THE TERMS & W) THE DERIVATION WILL NOT BE GIVEN HERE (SEE LITERATURE) THE RESULT 21 THAT ENERGY FLUX IN A RAY TUBE

IS CONSTANT

RECALLING THAT

ENERGY FLUX of polle² (For Pwayes) THIS MEANS THAT RAY AMPLITUDES VARY INVERSELY AS VPa AND ALSO AS 1/VA WHERE A is the CROSS-SECTIONAL AREA OF THE RAY TUBE.

TRAVEL TIME IS STATIONARY WITH RESPECT TO PERTURBATIONS OF THE PATH

APRICATION IN TOMOGRAPHY

NOTE THAT IT IS NOT TRUE

THAT THE PERTURBATION OF THE RAY PATH IS 2ND ORDER.

2 ATTENUATION AND PHYSICAL DISPERSION OF SEISMIC WAVES (RECALL DR. YANOVSKAYA'S LECTURES RNOTES) Stress E a Strain $\mathcal C$ RCREEP FUNCTION É Ĉ. Strain TA SEFESS RELAXATION €. FUNCTI UN $\dot{\psi}$ (t) È FOR ASINUSOIDAL SHEAR DISTURBANCE

 $u = U_{c}e^{i\omega t}$

$$
\mathcal{T}(t) = \mu(\omega) \epsilon(t)
$$

WHERE $\mu(\omega)$ is COMPLEX AND

FREQUENCY DE PENDENT

$$
T
$$

$$
S
$$

$$
T(t) = K(w)E(t)
$$

$$
W = \frac{1}{2\pi} \int_{-\infty}^{\infty} 4(t) e^{-i\omega t} dt
$$

$$
W = \frac{1}{2\pi} \int_{-\infty}^{\infty} 4(t) e^{-i\omega t} dt
$$

$$
W = \frac{1}{2\pi} \int_{-\infty}^{\infty} 4(t) e^{-i\omega t} dt
$$

$$
\mu(\omega) = i\omega \overline{\psi}(\omega)
$$

IT IS CONVENTIONAL TO DEFINE

$$
Q_{\mu\nu}(\omega) = \frac{R_{\mu}(\omega)}{\Gamma_{\mu\nu}(\omega)}
$$
 > > 1

BUT OFTEN MURE CONVENIENT TO USE

$$
P_{\mu\nu}(\omega) = \frac{1}{Q_{\mu}(\omega)} \quad \ll 1
$$

Writing
$$
\frac{1}{v_5} = \sqrt{\frac{p}{\mu(w)}} = \bar{s_0} - i s_2
$$

= Re($\frac{1}{v_5}$)(1 - $\frac{1}{2}i\varphi_{\mu}$)

THUS THE EXPRESSION FOR A PLANE WAVE TRAVELLING IN THE X - DIRECTION IS OF THE FORM

$$
w \sim U_0 e^{i\omega (t - x/v_s)}
$$

= $U_0 e^{-i\omega x s_x} e^{i\omega (t - x s_s)}$

with s_{2} = Re $(\frac{1}{2}) \cdot \frac{1}{2}$ g/m

DECAY IN ONE WAVELENGTY $exp\{-\omega \frac{2\pi}{\omega s} \frac{1}{2} g_{\mu} s_1\} = exp(-\pi g_{\mu})$ AMPLITUDE DECAY FOR S-WAVE $= e^{-\pi/Q_{\mu\nu}}$ PER CYCLE Op is ALSO SOMETIMES DENOTED BY QB (= Q FOR S-WAVES)

CORRESPONINGLY

AMPLITUDE DECAY FOR P-WAVE = e^{-TT/Q2} PER CYCLE

where
$$
Q_{\alpha} = \frac{2Re(Yv_{p})}{Im(Yv_{p})}
$$

Of COURSE WE MAVE

$$
K = (Re K)(1 + i q_K) \quad etc.
$$

SO IT IS EASY TO FIND EXPRESSIONS FOR QU IN TERMS OF QK, QM. IN PARTICULAR

IF GK=0 (USUALLY A FAIR ASSUMPTION) WE OBTAIN

$$
q_{\alpha} = \frac{q}{3} \frac{v_0^2}{v_0^2} \cdot q_{\beta}
$$

TO AN ADDITIONAL AMPLITUDE DECAY

[NOTE STRONG DAMPING OF MIGHT FREQUENCY WAVES

PHYSICAL DISPERSION

WE SAW THAT $\mu(\omega) = i\omega \overline{\psi}(\omega)$ WHERE $\overline{\psi}(\omega)$ = F.T. OF RELAXATION FUNCTION $\psi(t)$ WAVE VELOCITY $(\sqrt{5})$ is RELATED TO Re(M) AND DAMPING TO IM (M). BUT SINCE M(W) IS THE TRANSFORM OF A SINGLE REAL (CAVSAL) FUNCTION Re(p) AND In (p) ARE RELATED.

EG. FOR THE STANDARD LINEAR SOLID (SEE "WAVE PROPAGATION" NOTES FROM DR. WAVOUSKAYA)

$$
T + T_2 \dot{\tau} = \mu_0 (\epsilon + T_{\epsilon} \dot{\epsilon})
$$

\n
$$
\Rightarrow \mu(\omega) = \mu_0 (1 + i \omega T_{\epsilon})
$$

\n
$$
1 + i \omega T_{\epsilon}
$$

This can be used to find both
$$
q_{\mu}
$$
 (ω)
AND $v_5(\omega) = R_1 \sqrt{\frac{\mu(\omega)}{\rho}}$

VE INCREASES THROUGH THE THUS ABSORDTION BAND. FOR MANY ABSORPTION BANDS US INCREASES THROUGHOUT THE RANGE OF CONSTANT

QUANTITUELY IT CAN BE SHOWN THAT APPROXIMATELY, AND WITHIN THE BAND OF CONSTANT Q,

OR (INTEGRATING) FOR WI, WI WITHIN THE BAND

$$
ln \frac{v_s(\omega_2)}{v_s(\omega_1)} = \frac{1}{\pi} q_{\mu}^{const} ln(\omega_2)
$$

THESE LEAD TO A RELATIONSHIP BETWEEN THE DELAY OF AWAVE OF GIVEIN FREQUENCY AND 27'S DECAY, -THE PHENOMENON IS KNOWN AS PHYSICAL DISPERSION

I SEE LIU, ANDERSON, KANAMORI, GJ. 1976 AND REFERENCES CITED THEREIN]

WE CAN ALSO WRITE FOR THE COMPLEX VELOCITY

$$
U^{\prime}(\omega)=U_{o}\left(1+\frac{a_{1}}{\pi}\ln\frac{\omega}{\omega_{o}}+\frac{1}{2}iq\right)
$$

WHERE UT IS THE (REAL) VELOUTY AT REFERENCE FRÈQUENCY WO.

CONSEQUENTLY THE EFFECT ON THE SIGNAL IS REPRESENTED BY

$$
exp\{-\frac{1}{2}\omega t^{*}(1-\frac{2i}{\pi}\ln\frac{\omega}{\omega_{0}})\}\
$$

THIS REPRESENTS (HPPROXIMATELY, AND ASSUMING THAT THE ENTIRE SIGNAL IS WITHIN THE CONSTANT QM BAND) THE TOTAL AFFÈCT OF ATTENUATION ON THE SIGNAL.