



**The Abdus Salam  
International Centre for Theoretical Physics**



**1968-6**

**Conference on Teleconnections in the Atmosphere and Oceans**

*17 - 20 November 2008*

**Aspects of the theory of teleconnections**

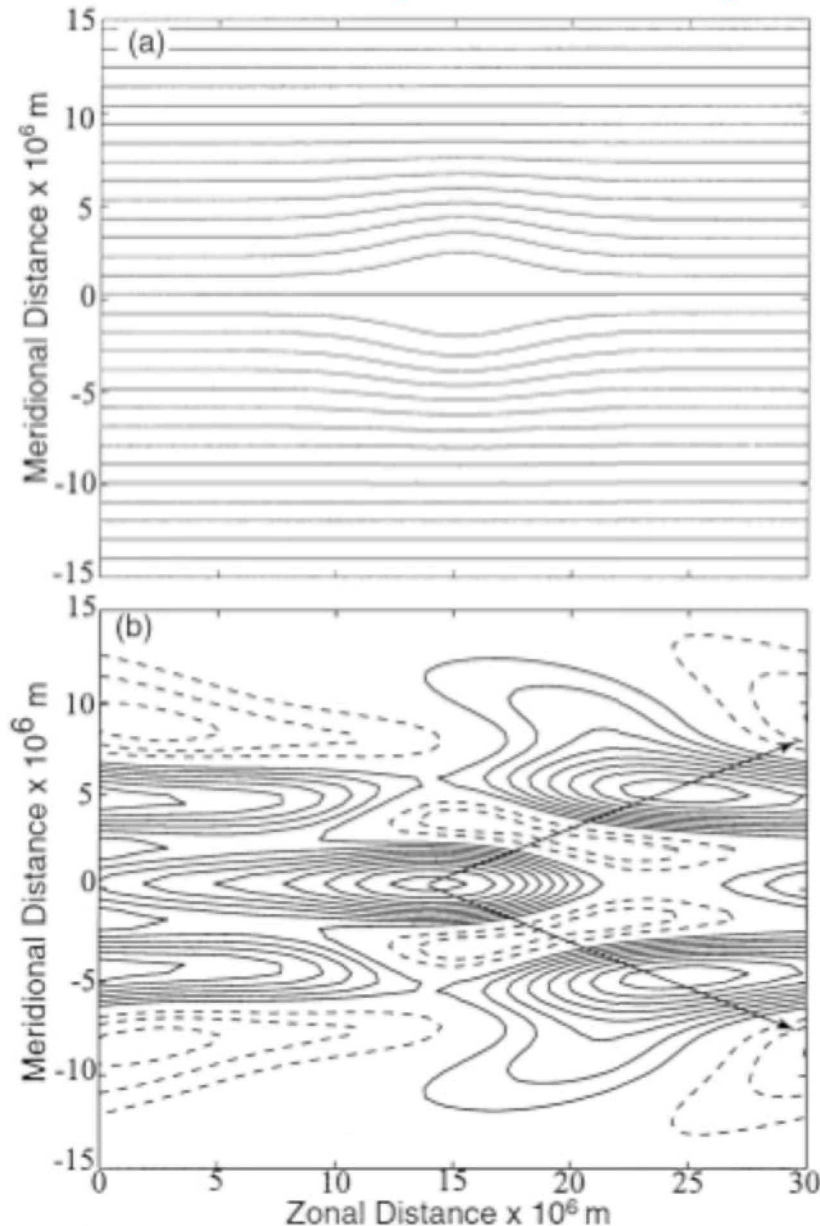
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# Some aspects of the theory of Teleconnections deserving of more attention

- Barotropic Rossby edge wave dynamics
- Effective linear operators
- Pseudo-momentum

Isaac Held  
Trieste, Nov. 16, 2008

- **Barotropic Rossby edge wave dynamics**



Barotropic instability of  
zonally asymmetric flow

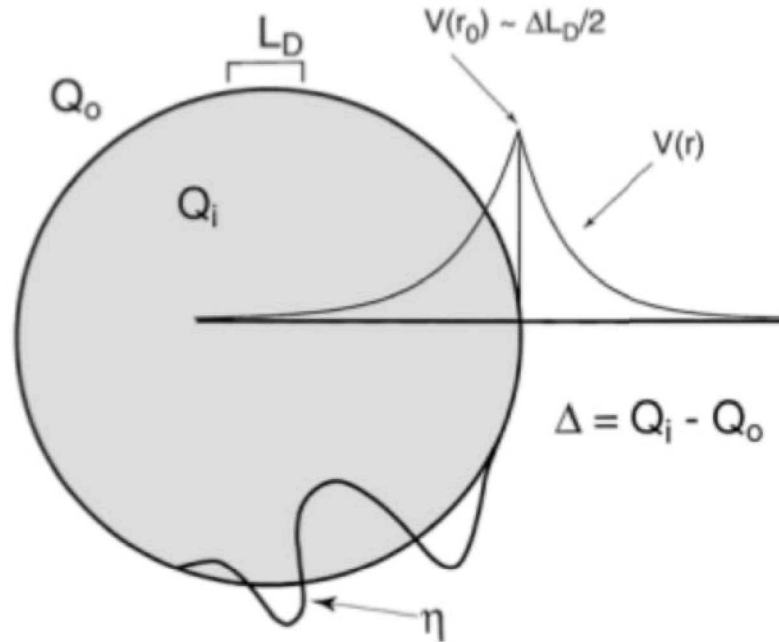
Instability often takes the PNA-like  
form of waves emanating from  
jet exit region

But no examples are simple enough  
that they provide useful intuition  
into what characteristics of the flow  
control various features of the  
instability

Swanson, JAS, 2000

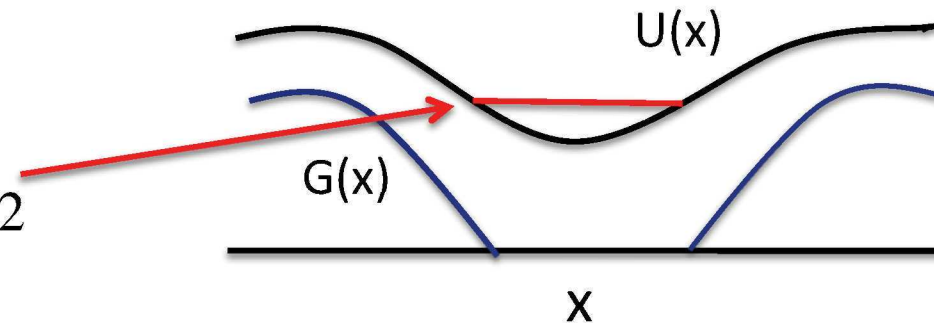
Quasi-geostrophic  
shallow water  
vortex

K. Swanson, JAS, 2000  
QJRMS, 2001  
J. Clim, 2002

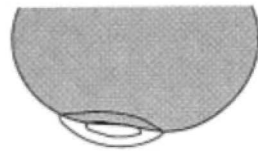
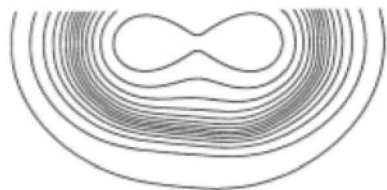


$$\omega = Uk - \frac{\Delta k}{2\sqrt{k^2 + L_D^{-2}}}$$

$$G_x = \frac{\partial \omega}{\partial k} \Rightarrow 0 \quad \text{as} \quad U \Rightarrow \Delta L_D / 2$$

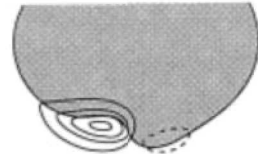
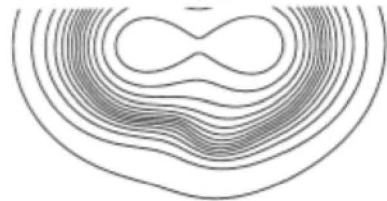


Absolute instability – growth in place typically requires that group velocity be zero at some frequency



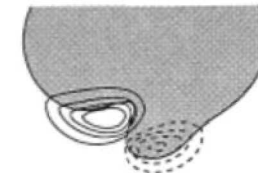
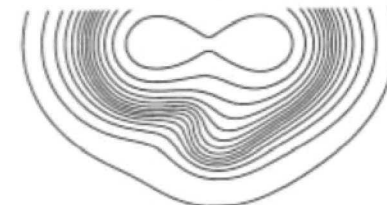
*t = -6 days*

Nonlinear evolution when vortex is perturbed by topography so that group velocity  $\Rightarrow 0$

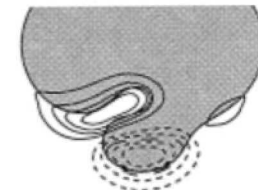


*-4.5 days*

Swanson, J. Clim, 2002

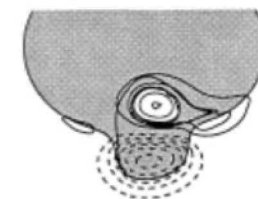
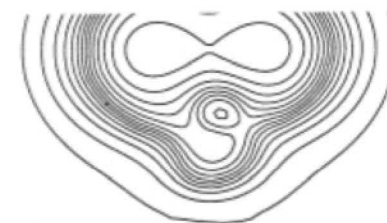


*-3 days*



*-1.5 days*

Typically evolves towards “blocked” state, but solutions are more wave-like when propagation is enabled by additional contours



*Onset*

- **Effective linear operators**

Linear stationary wave responses

- linearizing about time mean flow –
- not very robust – often need to add arbitrary damping

Complexities

- interactions with transients
- interactions with latent heating
- variability in “mean flow”

Is there a short cut to developing effective linear operators that incorporate all such effects?

How is this **effective linear operator** related to linearizing about time mean flow

$$\frac{dx}{dt} = N(x) + \delta F$$

$$\delta \bar{x}^{-t} = L \delta F \quad \text{or} \quad L^{-1} \delta \bar{x}^{-t} = \delta F$$

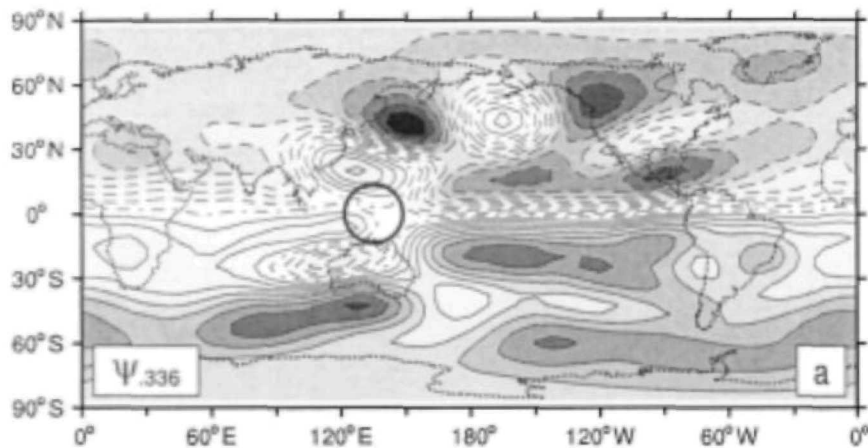
$$L = \int_0^{\infty} C(\tau) C^{-1}(0) d\tau$$

Leith, JAS, 1975

$$C_{ij}(\tau) = \overline{x_i(t) x_j(t + \tau)}^t$$

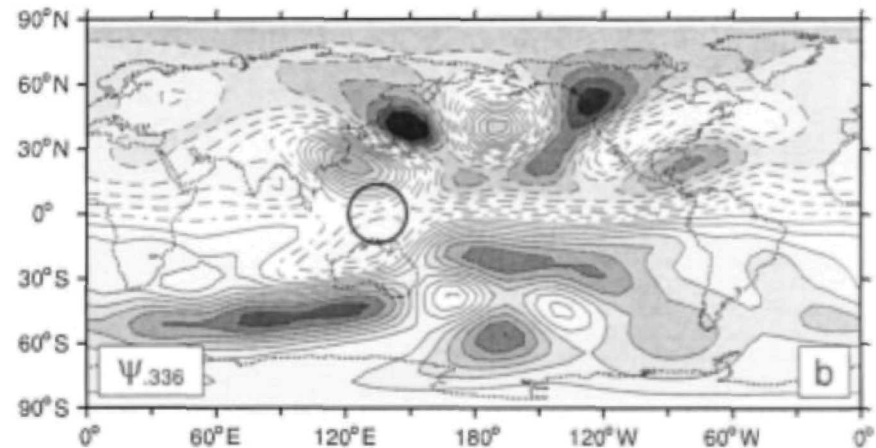
GCM response to localized  
equatorial heating

CCM0



response obtained from  
fluctuation-dissipation relation

FD



upper tropospheric streamfunction

Gritsun and Branstator, JAS 2007



- Pseudo-momentum

External Rossby waves:

equivalent barotropic;

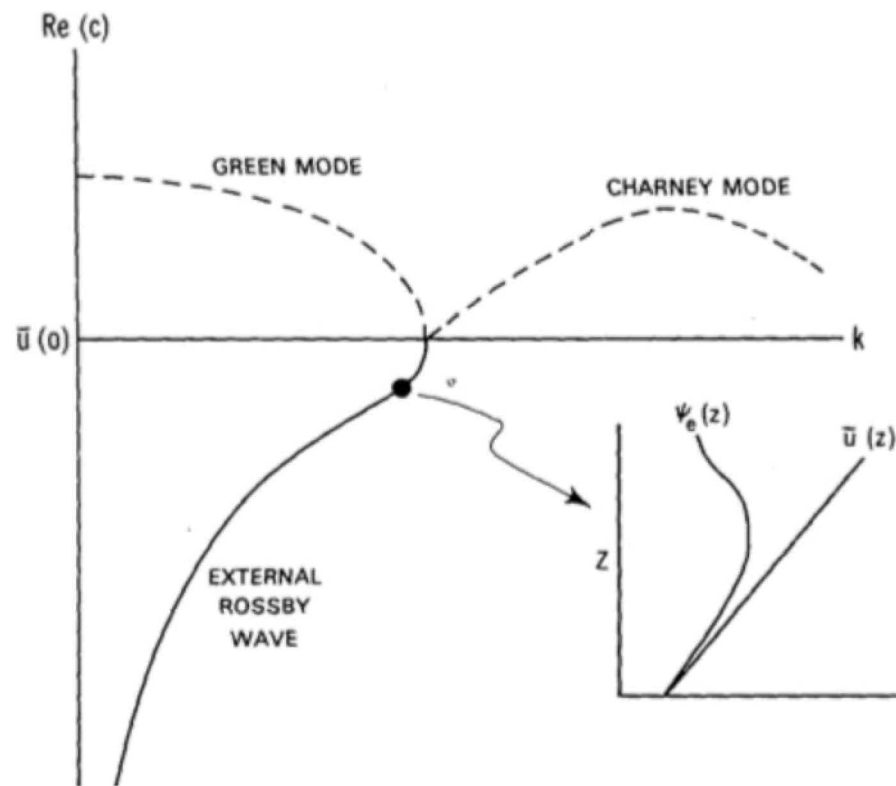
EAPE comparable EKE;

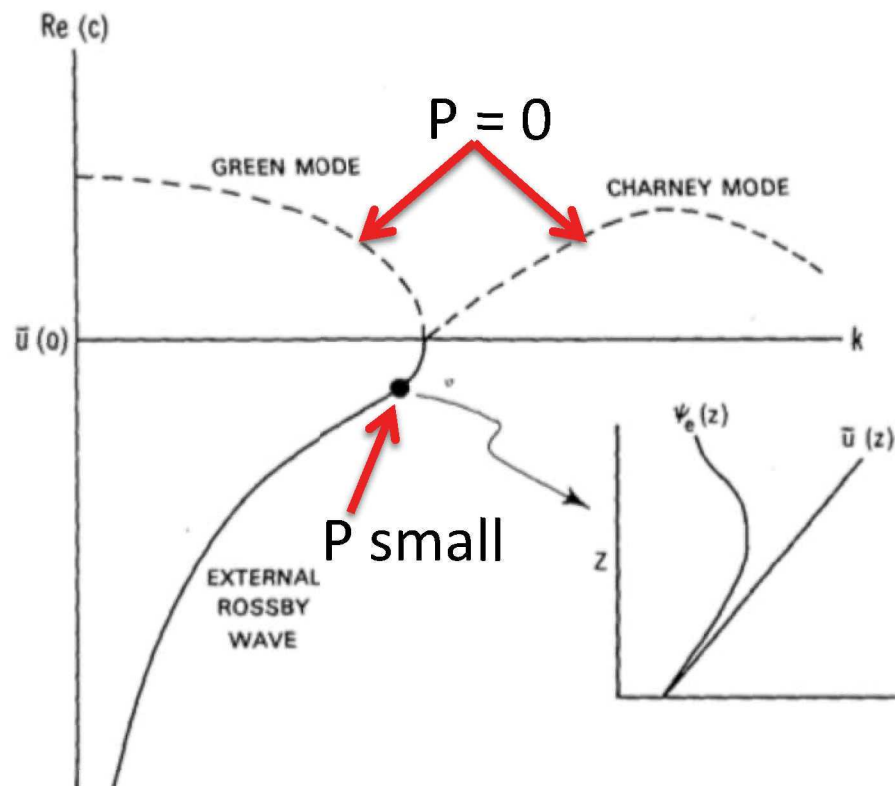
top-heavy in EKE, but meridional particle displacements  
are of the same order in upper and lower troposphere:

$$\eta' \sim \psi'/U$$

Small pseudo-momentum

# Dispersion relation for Charney's model





$$P = \int \eta'^2 \frac{\partial q}{\partial y} dz - \eta'^2 \frac{\partial \Theta}{\partial y} \Big|_{z=0}$$

*Implications of the Pseudomomentum of an external Rossby wave being small:*

- *Wave is much easier to excite than upper level barotropic disturbance in isolation*
- *Damping at low levels increases wave amplitude*
- *Vertically averaged zonal mean flow deceleration that occurs when wave propagates into a region is much smaller than would be the case of an upper level disturbance in isolation*

- **Barotropic Rossby edge wave dynamics**

*Best chance for simple models of the dynamics of low-frequency variability?*

- **Effective linear operators**

*Promise to provide more robust linear theories.*

*Can one generate these directly from observations?*

- **Pseudo-momentum**

*Low level dynamics can be as important as upper level dynamics for the external Rossby waves that underlie wavelike teleconnection patterns*