



The Abdus Salam
International Centre for Theoretical Physics



2018-28

Winter College on Optics in Environmental Science

2 - 18 February 2009

Atmospheric Monitoring, Differential Optical Absorption Spectroscopy

DOAS I, Basics

Platt U.
University of Heidelberg
Germany

Atmospheric Monitoring, Differential Optical Absorption Spectroscopy – DOAS I, Basics

Ulrich Platt

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- The principle of DOAS
- The (many) Variants of DOAS
- Examples
- Sample Results
- (preliminary)Summary



WINTER COLLEGE ON OPTICS IN ENVIRONMENTAL SCIENCE
2 – 13 February 2009, Trieste, Italy

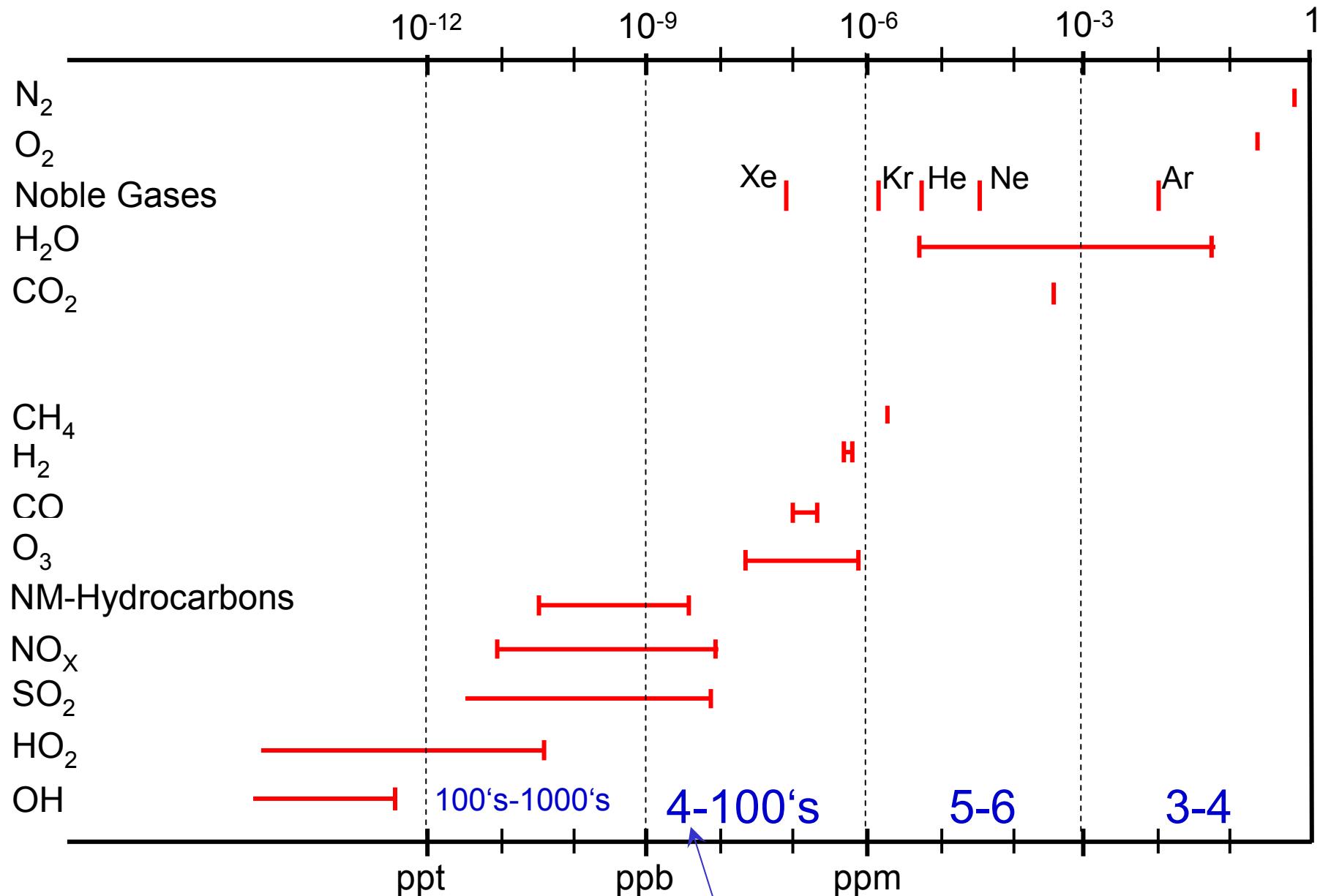


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Mixing Ratios of Atmospheric Trace Gases



Requirements for Measurement - Techniques

- **Sensitivity:** Species (like OH, IO) present at mixing ratios of $\approx 10^{-13}$, (0.1 ppt, about 2×10^6 molecules/cm³) can significantly influence atmospheric chemistry.
- **Specificity:** The result of the measurement of a particular species must not be influenced by any other trace species present in the air.
- **Spatial coverage:** In-situ vs. remote sensing
- **Time resolution**
- **Calibration** should be easy (inherent?), stable, ...
- Further desirable properties:
 - Simplicity of design and use of the instruments based on it.
 - Possibility of unattended operation.
 - Weight, portability, and dependence of the measurement on ambient conditions.



Diversity of Measurement Tasks

Long Term Observations note 'operator dilemma'

- Stratospheric ozone trend
- Change of Stratospheric chemistry (NDACC)
- Stratospheric (chlorine) source-gases in the troposphere
- Tropospheric ozone trend (GAW)
- Greenhouse gases

Regional Episodic Events

Remember Lectures
by Profs.
Singh and Wagner ...

- Pollution monitoring
- Urban plume evolution
- Continental plumes
- Antarctic Ozone Hole
- Polar boundary-layer ozone loss events

Fast in-situ (Photo)chemistry (process studies)

- Free radical (OH/HO_2) photochemistry



Advantages of Spectroscopic Techniques ...

- Direct identification of trace gas molecules
- Remote sensing capability
- Inherent calibration (ideally ...)
- Real time measurements
- Non – contact measurement
-

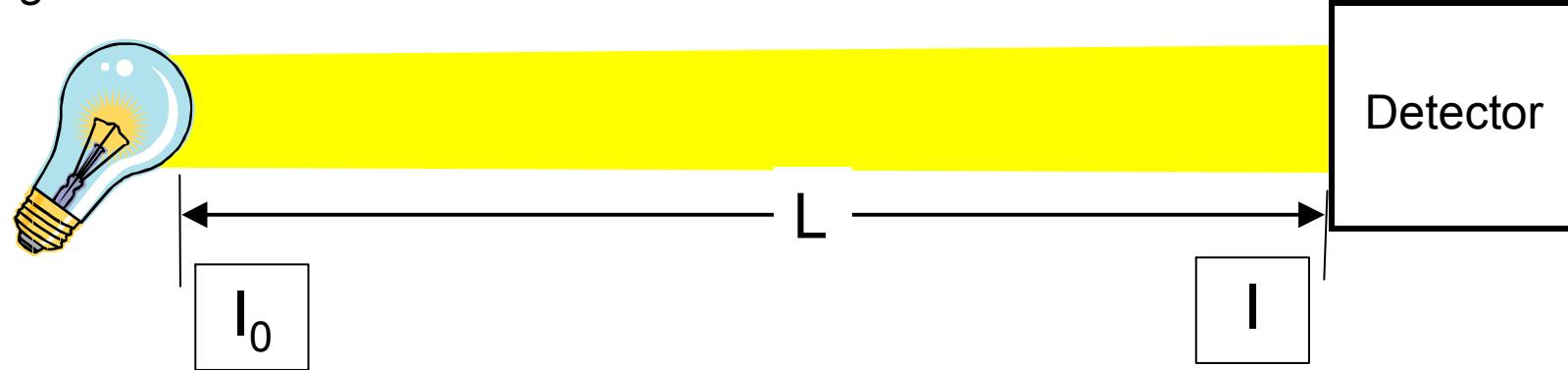
Remember Lectures by
Prof. Svanberg ...



Spectroscopic Measurement of Atmospheric Trace Gases

The Principle

Light source



$$I = I_0 \cdot e^{-\bar{c} \cdot L \cdot \sigma} \text{ Lambert-Beer's law}$$

\bar{c} = average trace gas concentration

L = length of light path

σ = absorption cross section

(or Bouguer-Lambert law, it was discovered by Pierre Bouguer in 1729, Johann Heinrich Lambert in 1760, and August Beer in 1852)



Gustav Robert
Kirchhoff

1824 - 1887

in Heidelberg:
1854 - 1874

Robert Wilhelm
Bunsen

1811 - 1899

in Heidelberg:
1852 - 1899

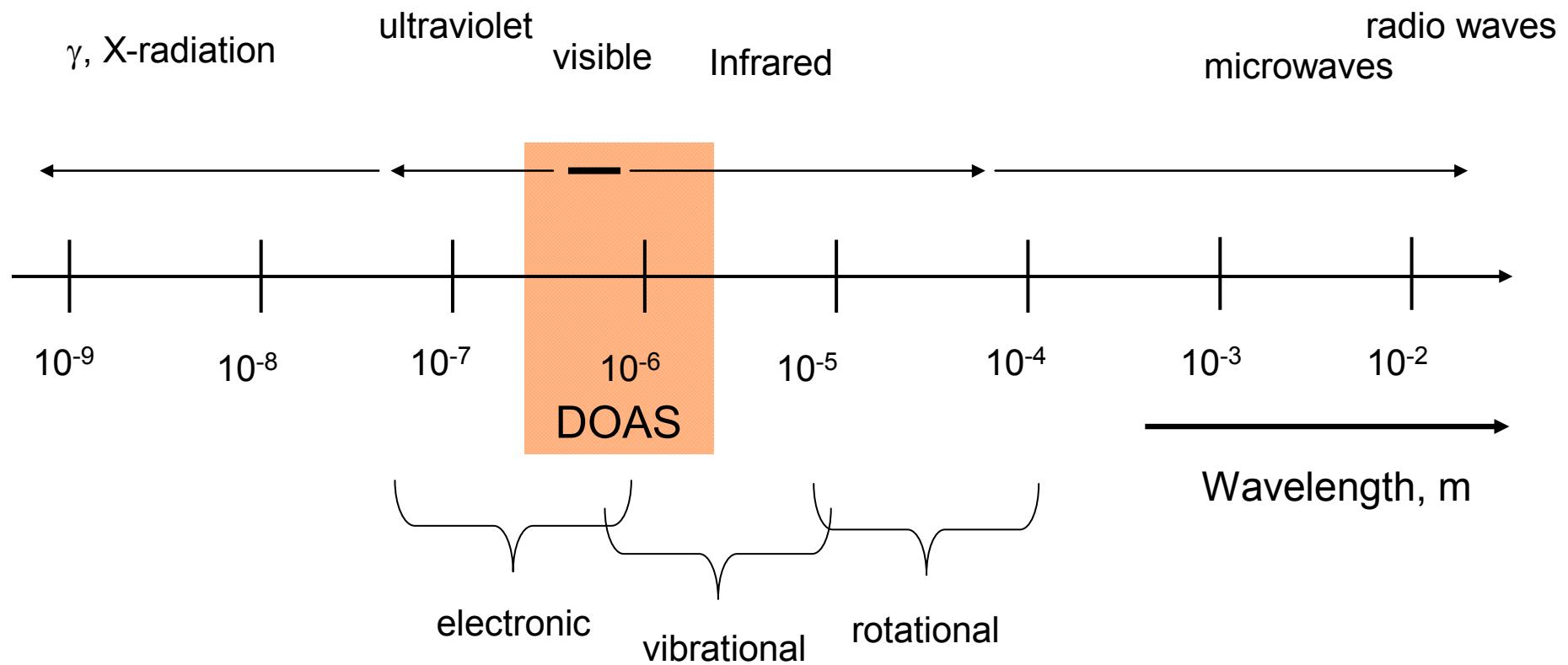


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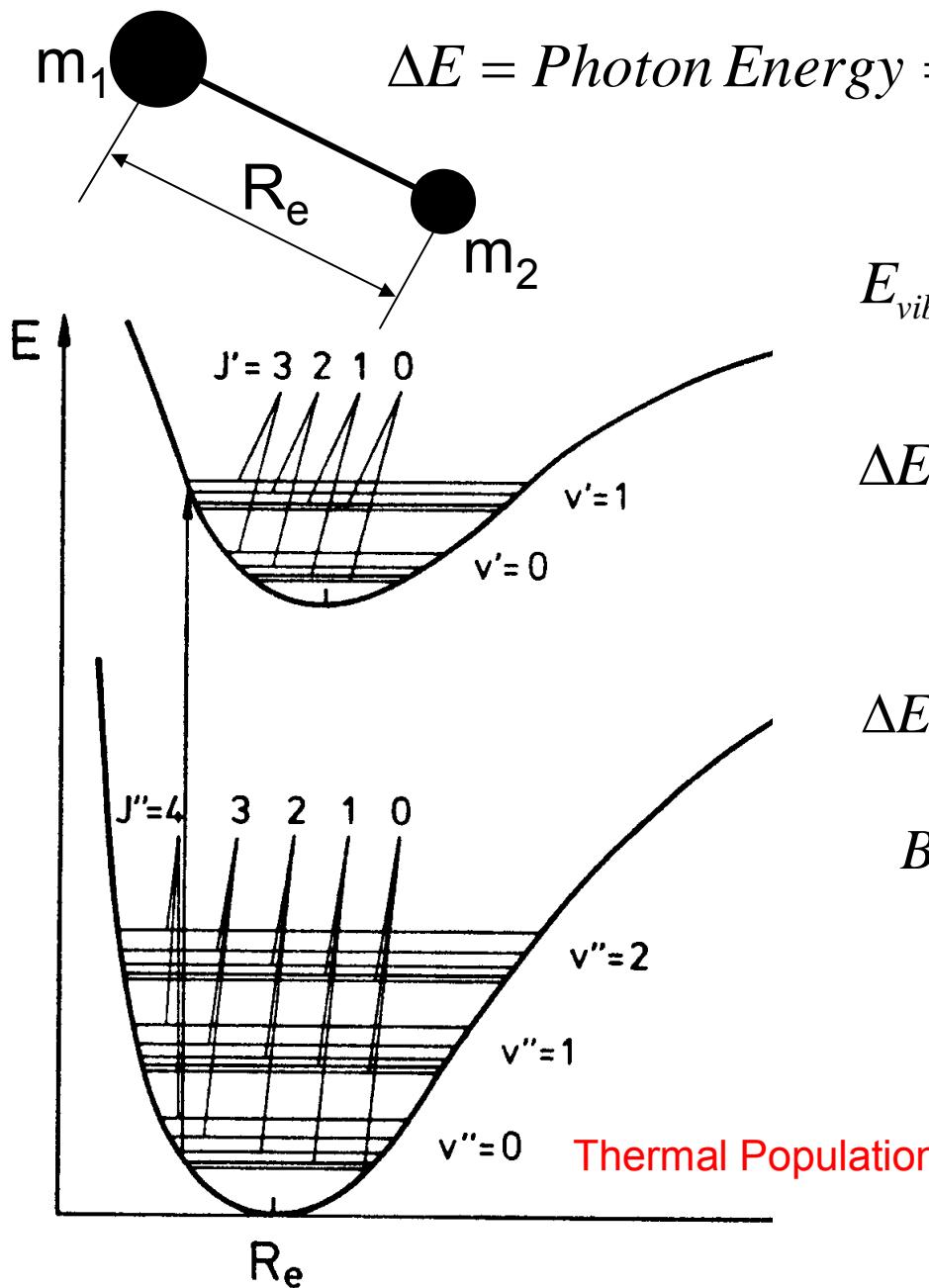
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Electromagnetic Radiation



Molecules: Electronic – Rotational –Vibrational Spectra



$$E_{\text{vibration}}(v) = \left(v + \frac{1}{2}\right)hv'_0$$

$$\Delta E_v = \left(v' + \frac{1}{2}\right)hv'_0 - \left(v'' + \frac{1}{2}\right)hv''_0$$

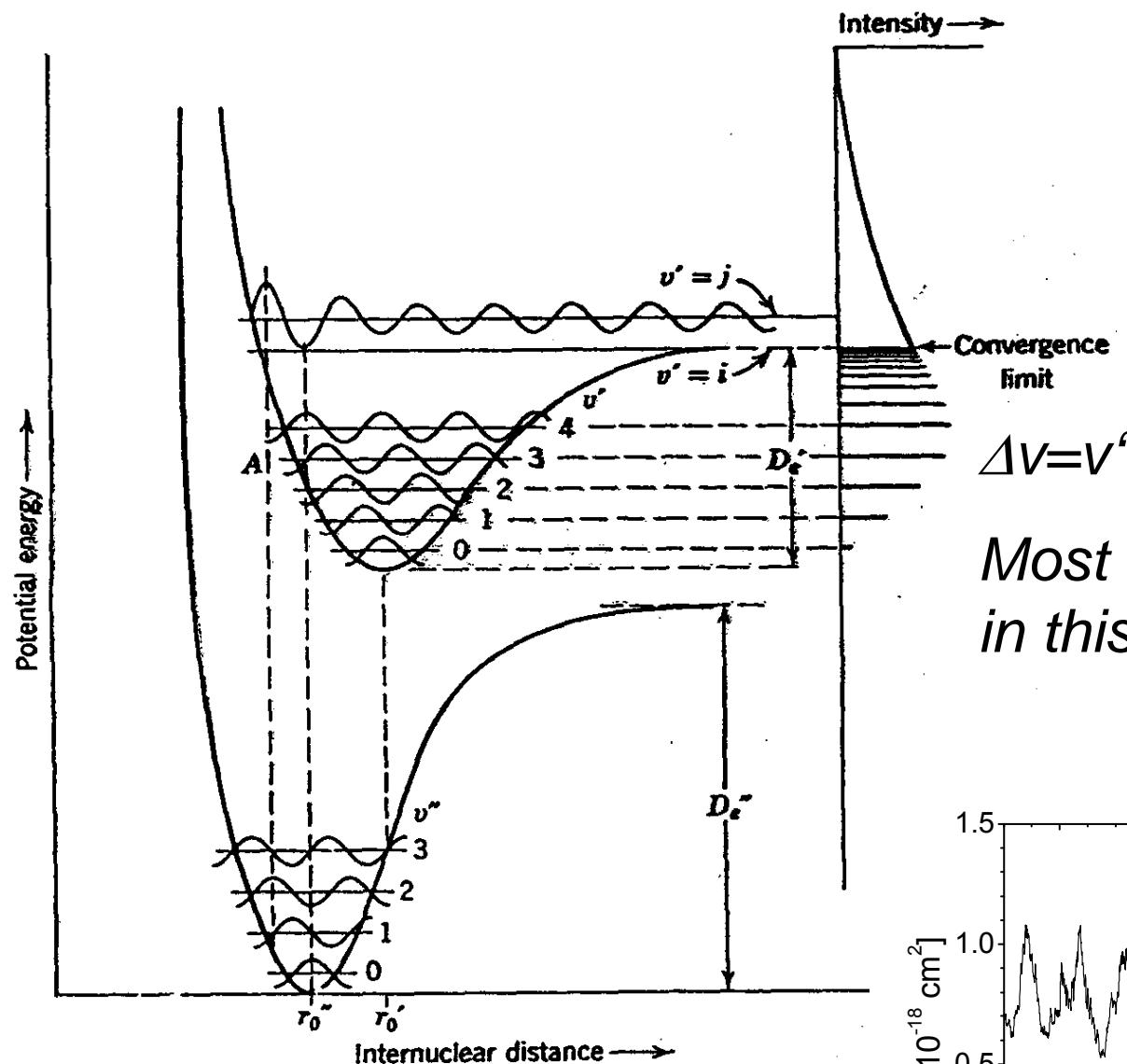
$$\Delta E_{\text{rotation}}(J', J'') = B'J'(J'+1) - B''J''(J''+1)$$

$$B = \frac{\hbar^2}{2\Theta} = \frac{\hbar^2}{2r_e^2 \frac{m_1 m_2}{m_1 + m_2}}$$

Rule

Atoms: Line Spectra
Molecules: Band Spectra

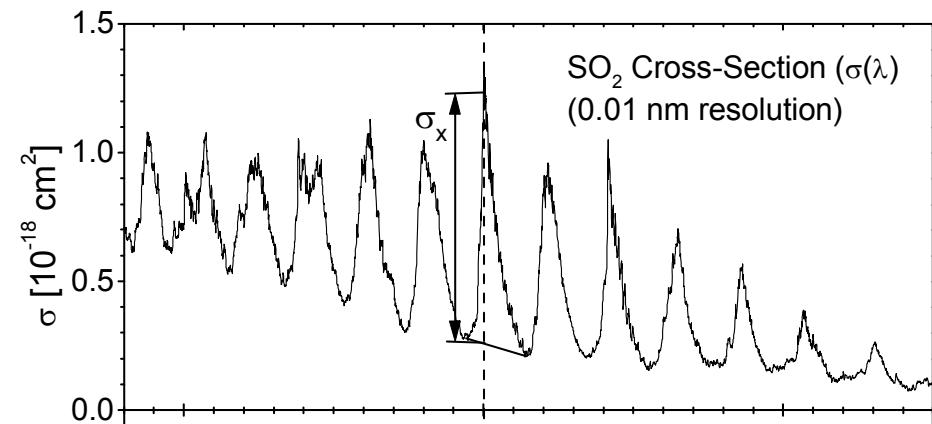
The Franck – Condon Principle



Electronic transition associated with absorption/emission of a photon occurs fast compared to the vibrational period.

$$\Delta v = v' - v'' = 4:$$

Most likely transition in this case

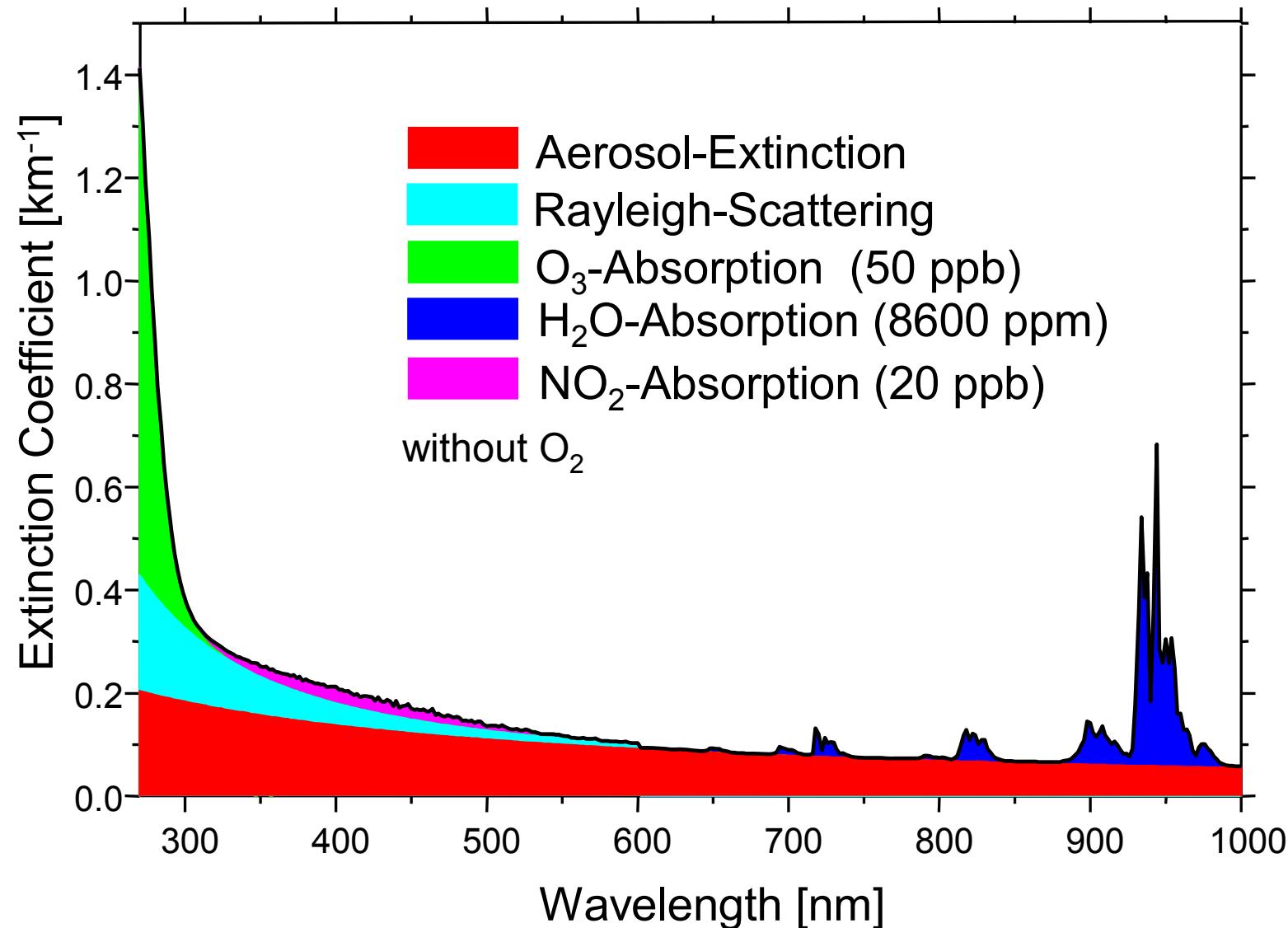


Problems with ‘Simple’ Absorption Spectroscopy

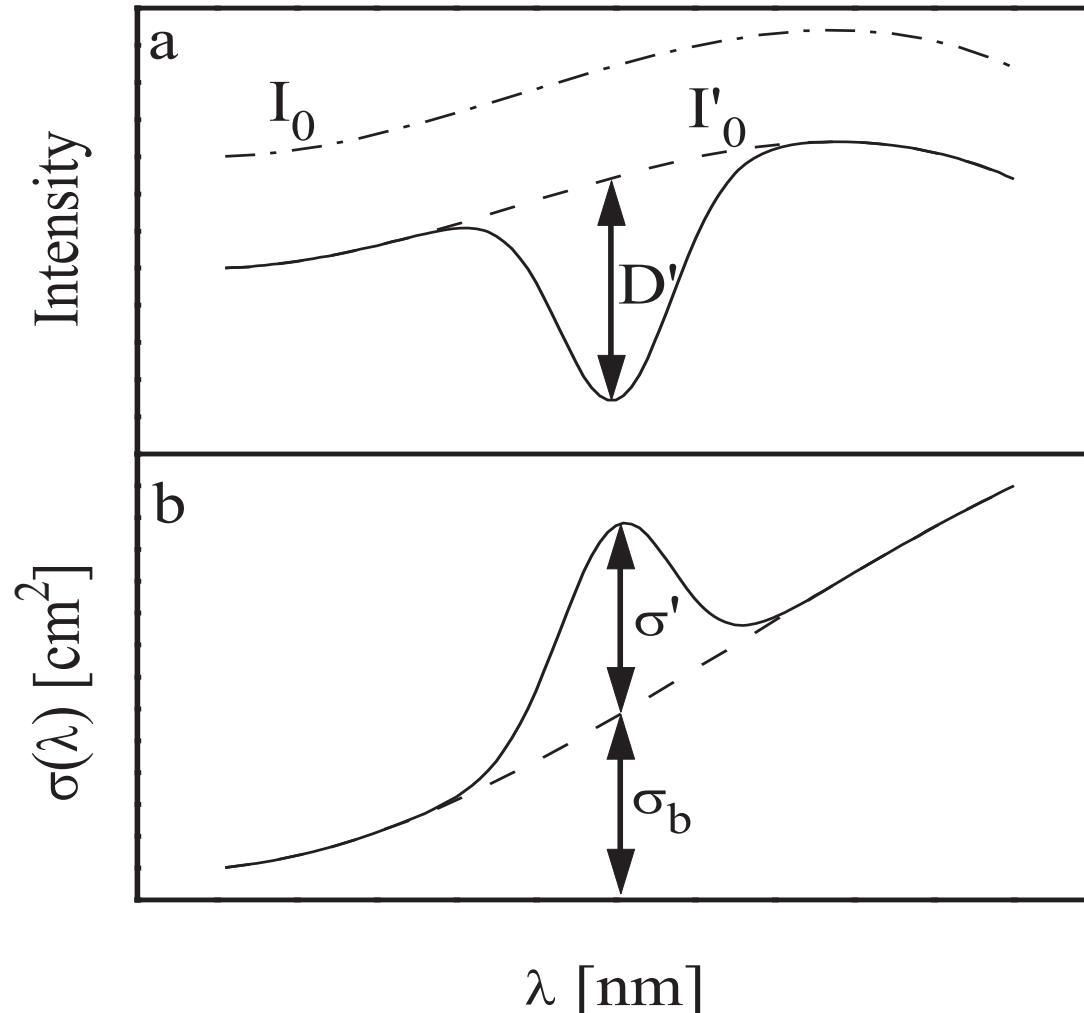
- How to distinguish extinction due to scattering by aerosol particles (Mie scattering) and by air molecules (Rayleigh scattering) from molecular absorption?
- How can absorptions due to more than one trace gas be separated?
- Effects of degradation of light source, optics, or detector?



How Transparent is the Air?



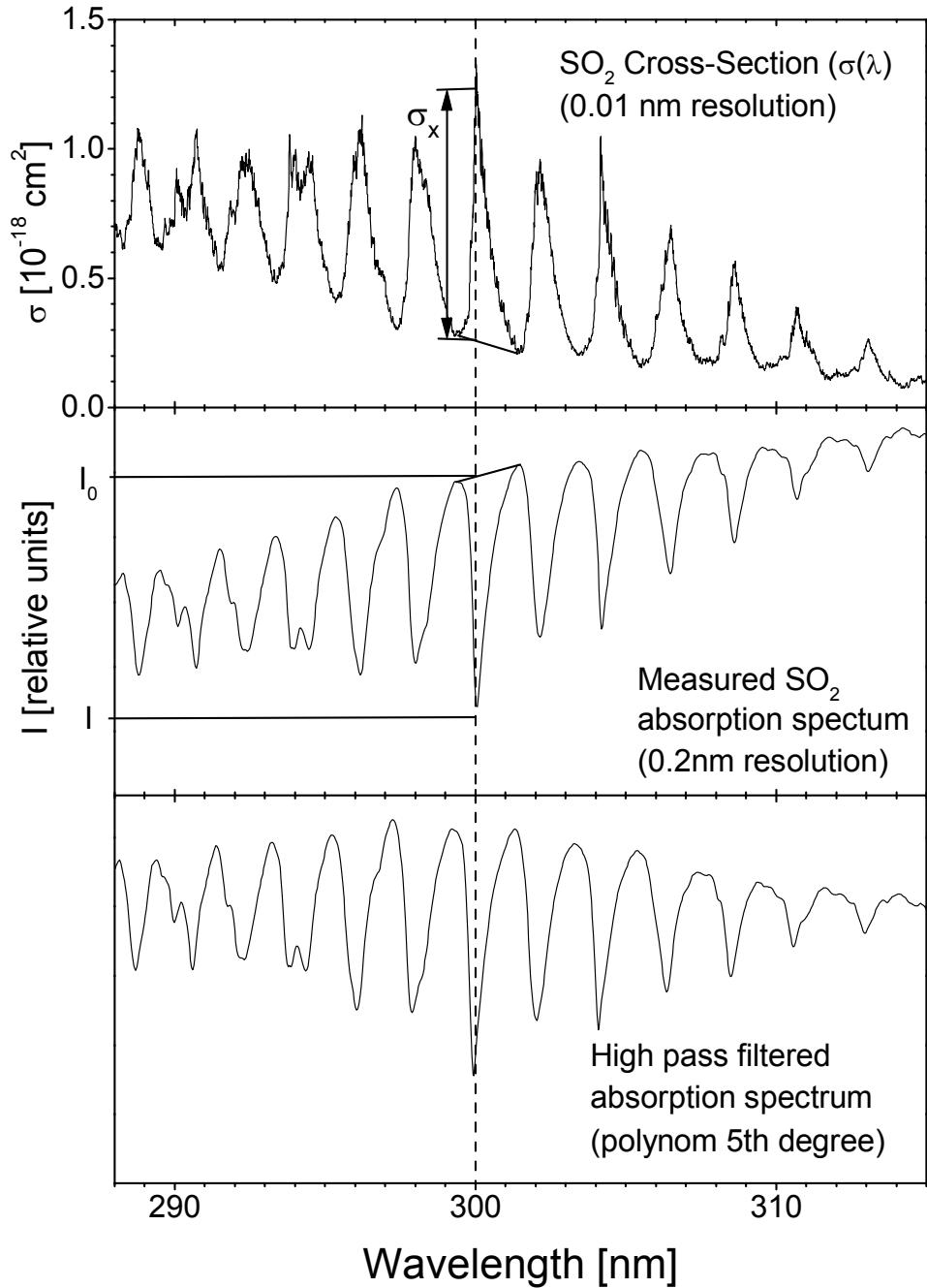
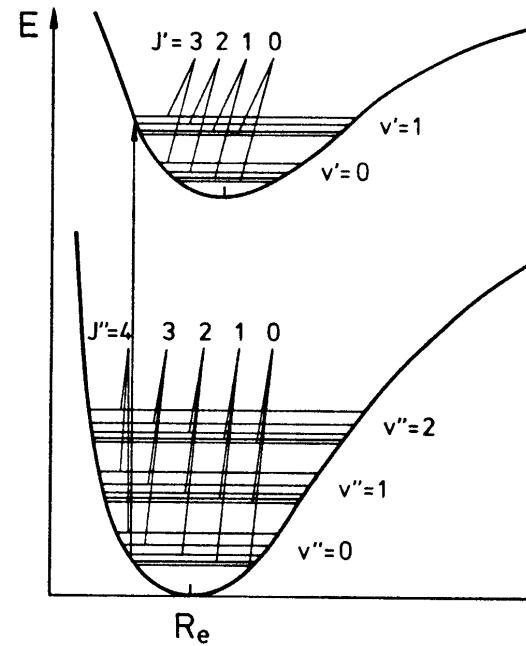
Differential Optical Absorption Spectroscopy – The Idea



- Use **differences** of intensities at different wavelengths
- Record the intensity in **many** (typ. several 100) wavelength **channels** (entire spectra)
- **High pass-filtering** of spectra
⇒ remove continuum
- **Fit reference spectra**
⇒ Make use of all spectral information



Example: Absorption Cross Section and DOAS – Spectrum of SO₂



Differential Optical Absorption Spectroscopy

Lambert-Beer's Law:

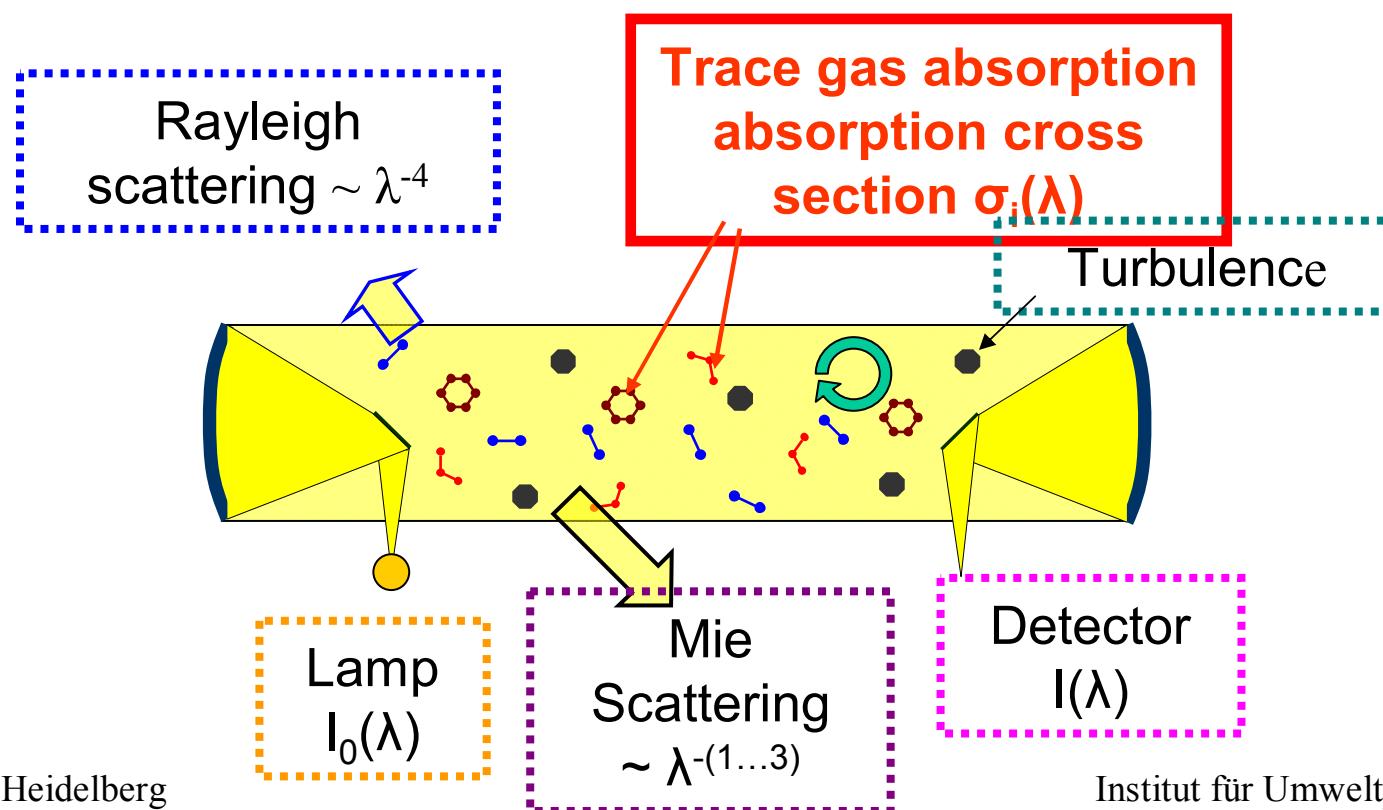
$$I(\lambda) = I_0(\lambda) \cdot e^{-\sigma(\lambda) \cdot c \cdot L}$$

$$I(\lambda) = I_0(\lambda) \cdot e^{-[\sum \sigma'_i(\lambda) \cdot c_i \cdot L + (\sigma_{bi} \cdot c_i + \varepsilon_{Ray}(\lambda) + \varepsilon_{Mie}(\lambda)) \cdot L] \cdot T(\lambda)}$$

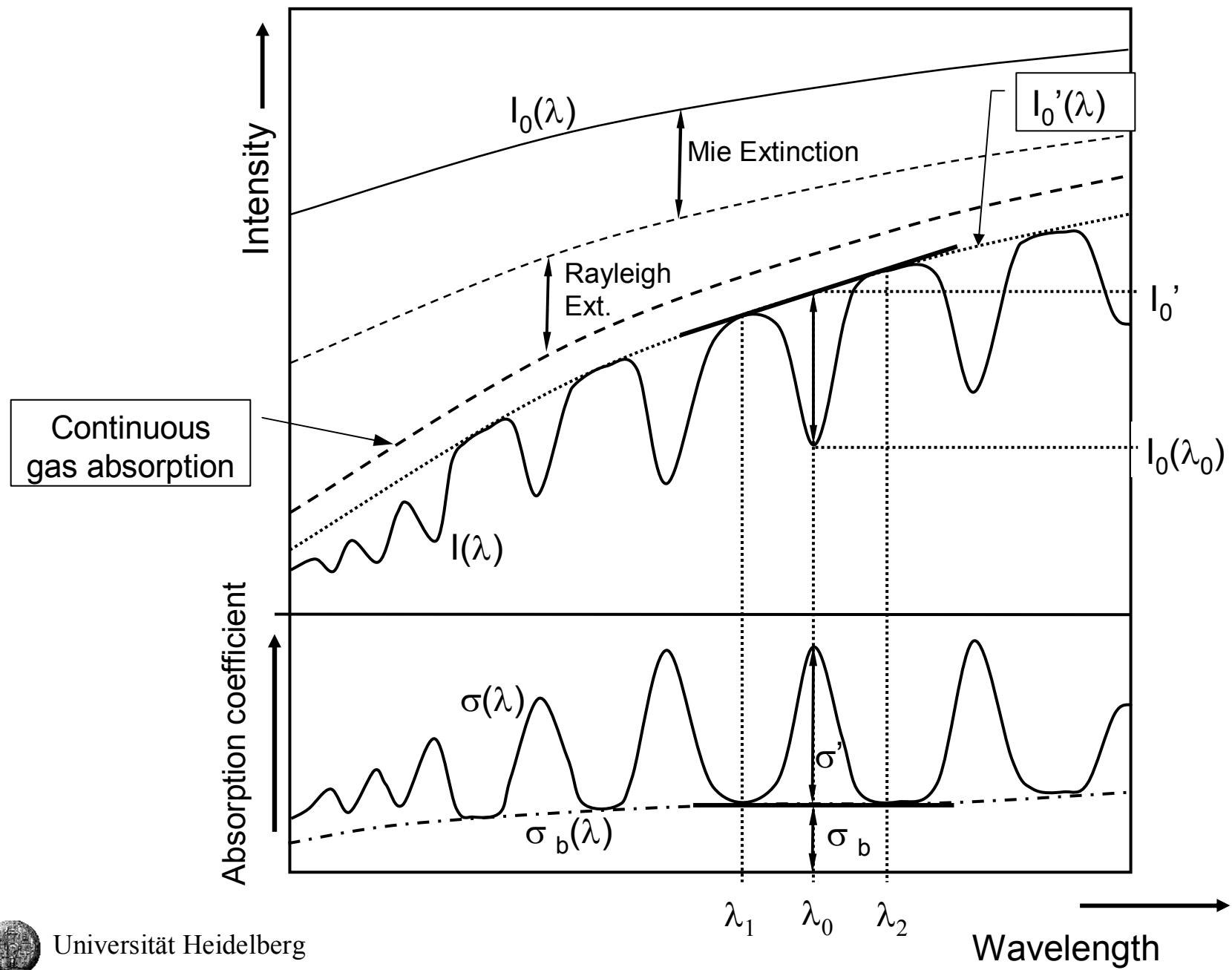
narrow-

Remove by high-pass filtering

wide band extinction



From Total Absorption to Differential Absorption



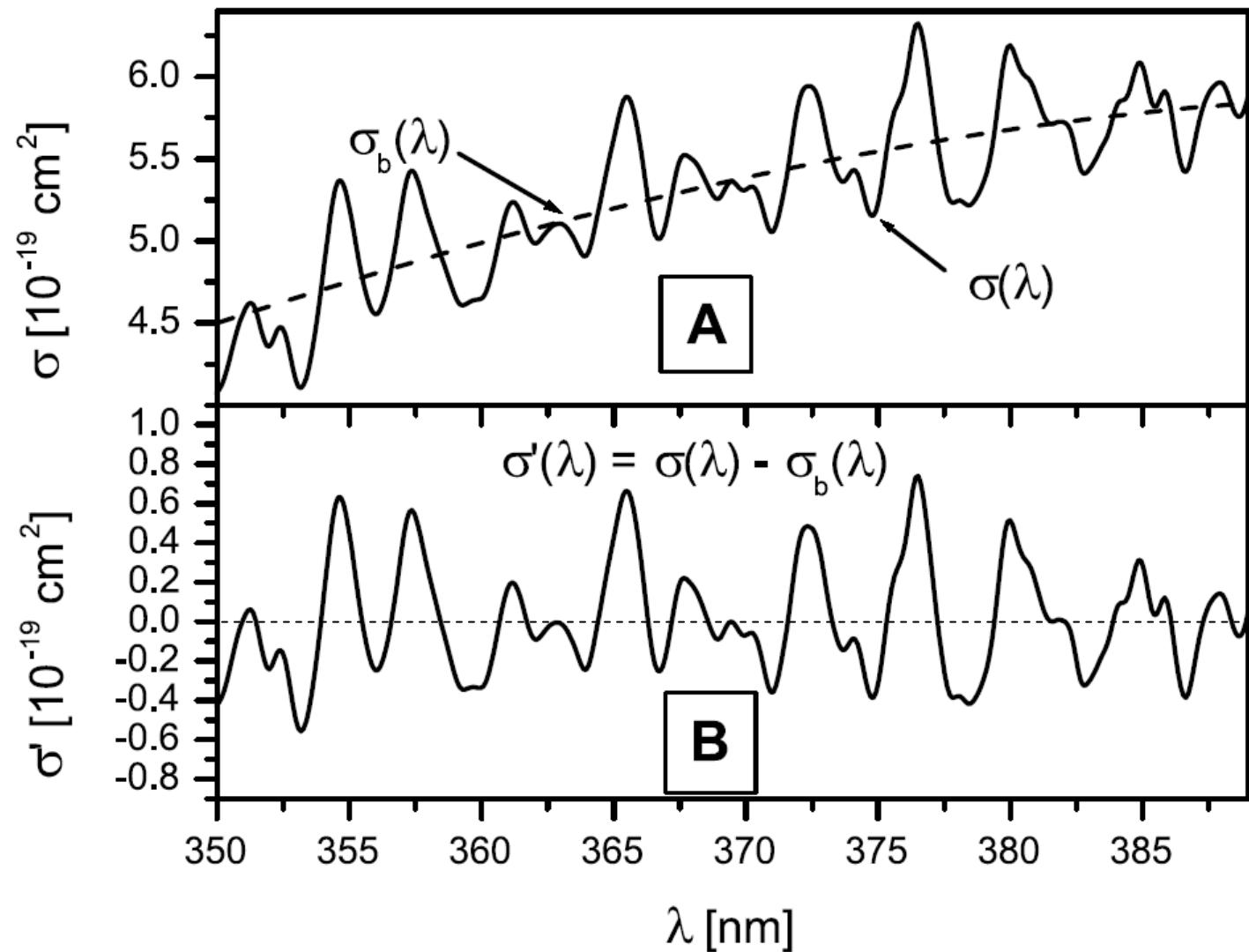
Total- and Differential Absorptipn Cross Sections

Example for the separation of the cross section into its differential and broad-banded parts using an NO_2 cross section.

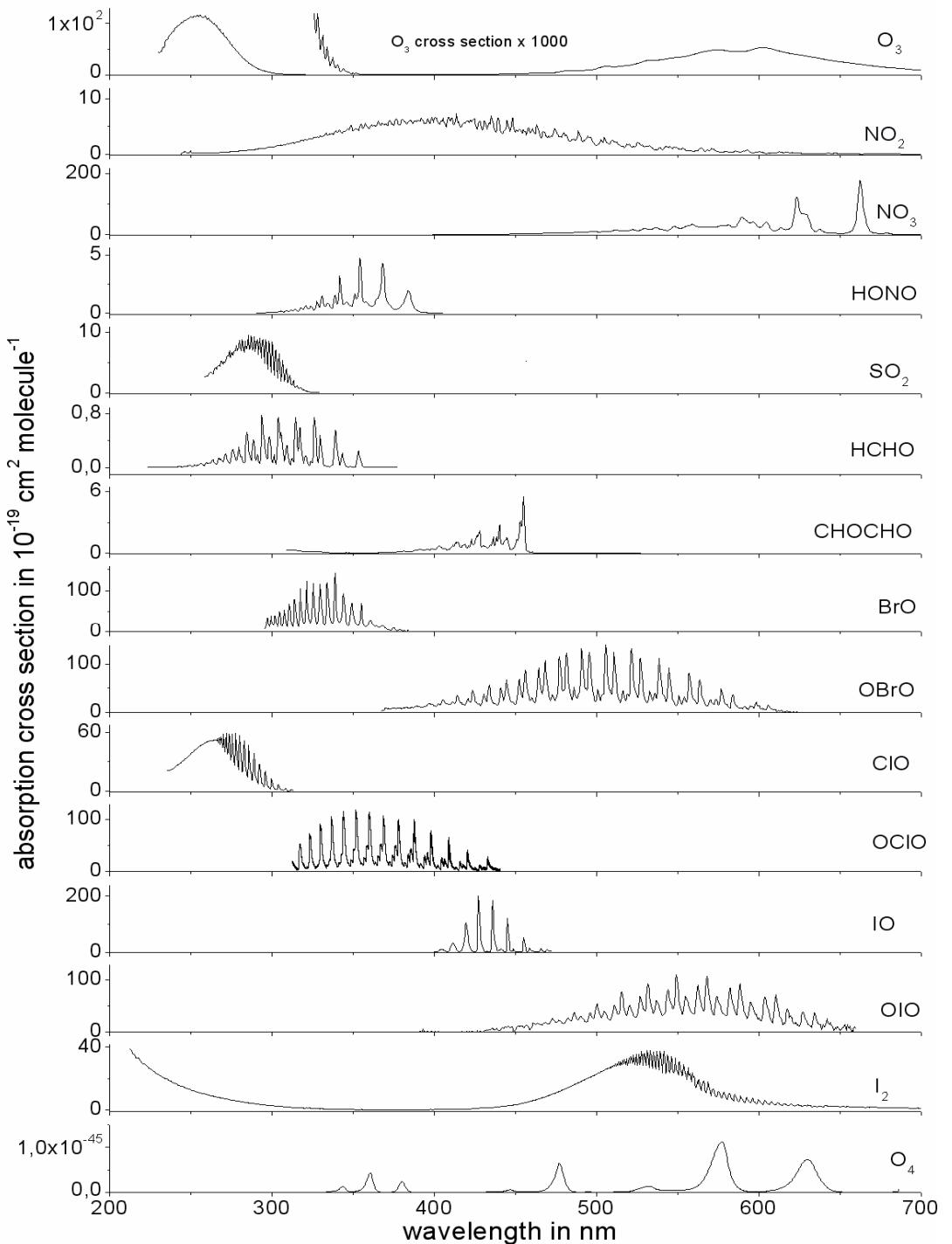
Total cross section (solid line) and broad-band structure (dashed line), derived by interpolating a 3rd order polynomial.

Differential cross section.

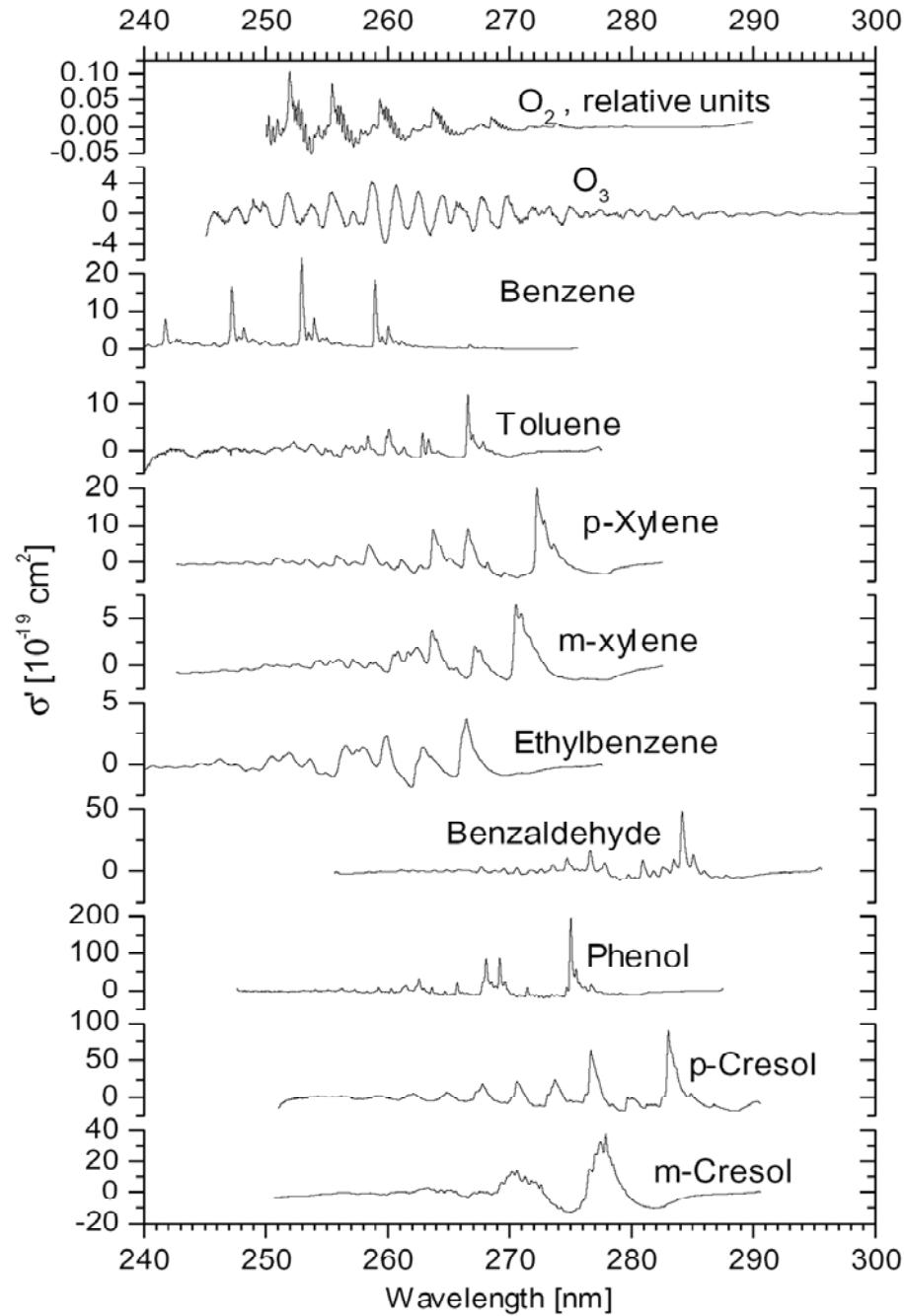
From: Friess 2000



Differential Cross Sections of Atmospheric Molecules

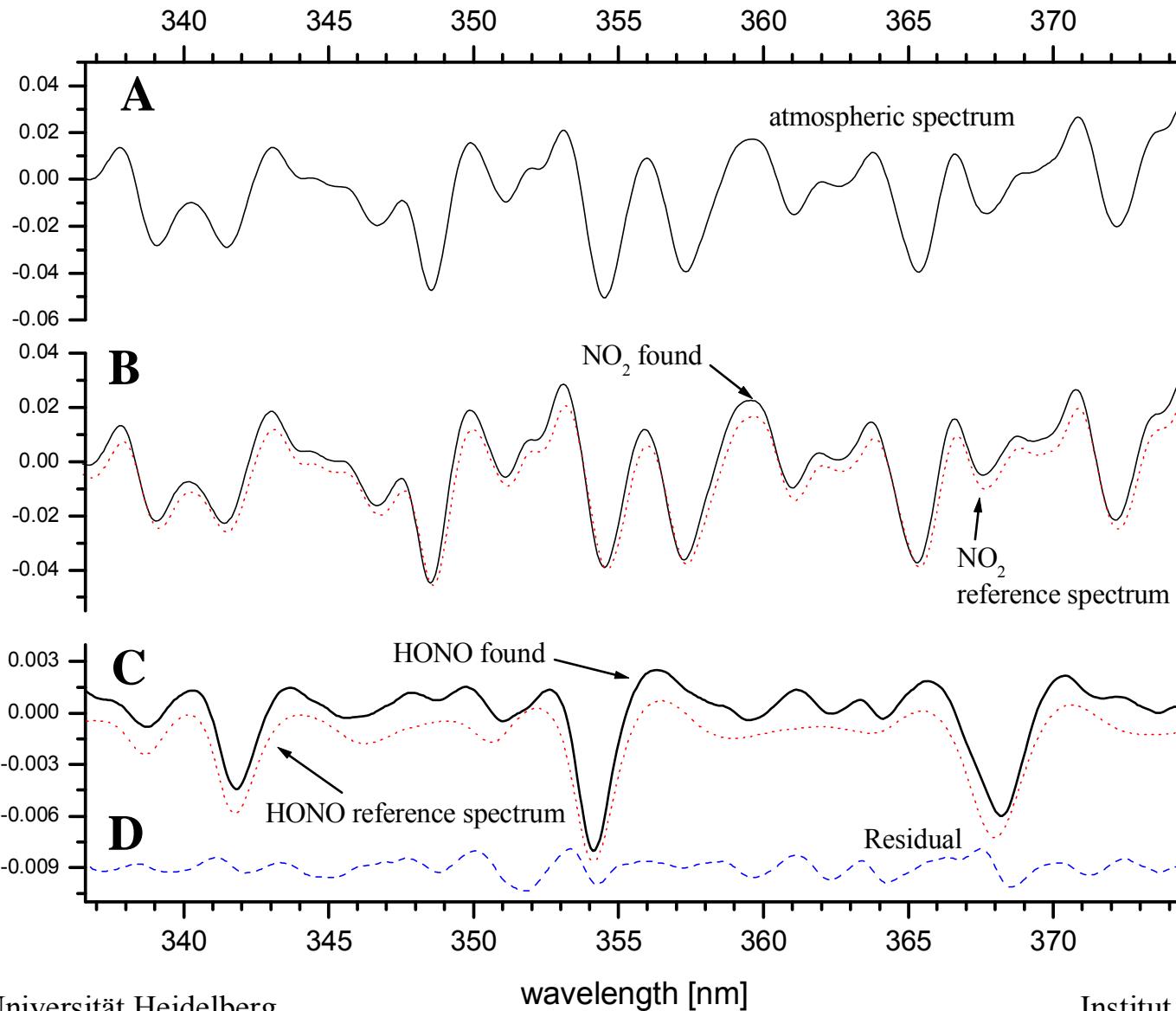


Absorption Cross sections of a Series of Monocyclic Aromatic Hydrocarbons and of O₂ in the Near UV



Deconvolution of NO₂ and HONO Spectra

PiPaPo – Campaign, Milano, 1998



Very weak
lines
typ. << 1%

Lots of
photons
needed
 $>> 10^4$

Note
suppressed
zero!



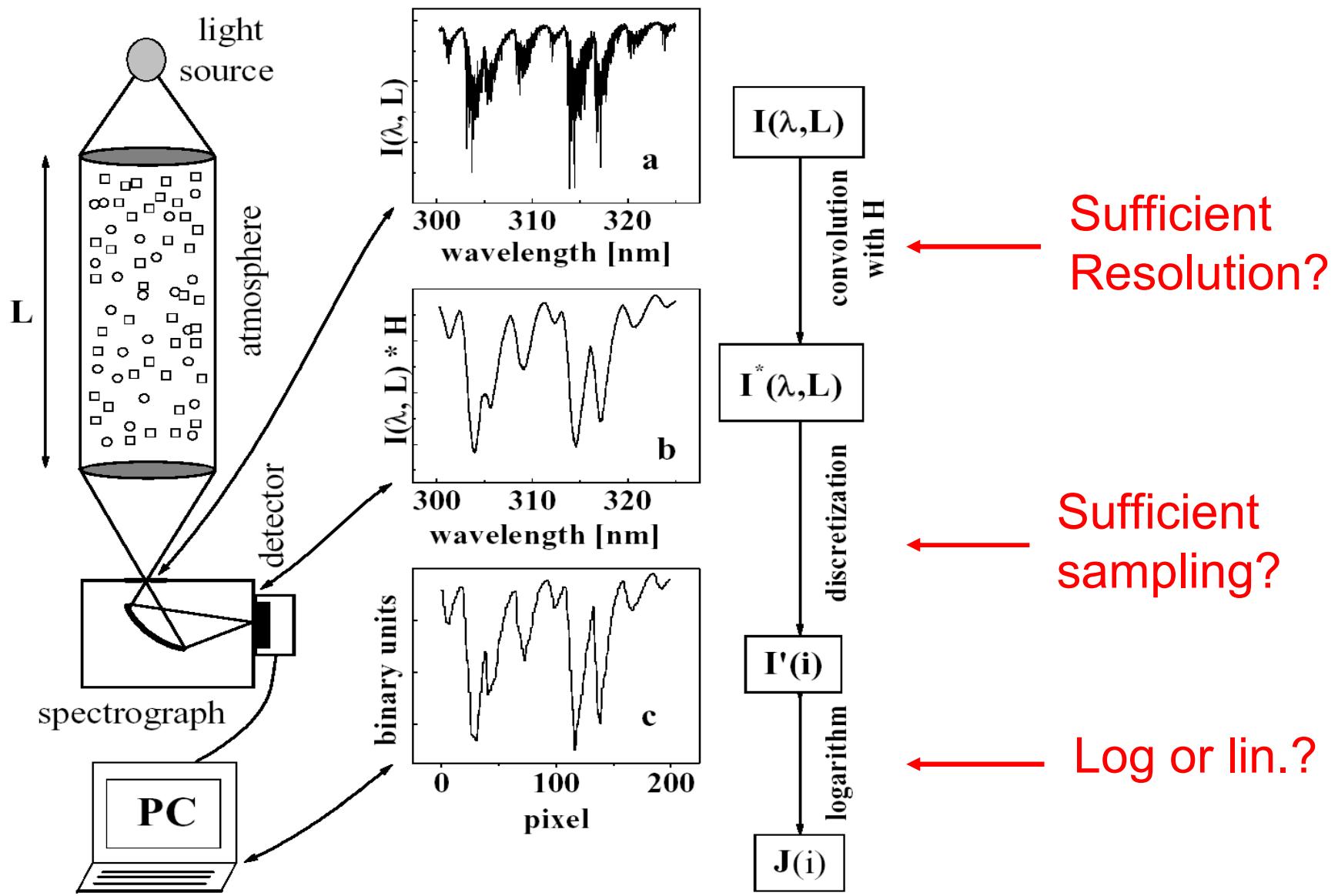
Quantitative Evaluation of DOAS Spectra

How is it actually done?

- “Technical correction”
 - **High pass filtering**
 - **Remove Fraunhofer Spectrum**
 - **Remove “Ring Effect”**
 - **Fit known absorption spectra**
 - **Account for spectral shift**
 - Account for change in dispersion
 - Correct sampling artefacts
 - Correct for “ I_0 effect”
- } Passive DOAS only



From Analog Light to Digital Spectrum

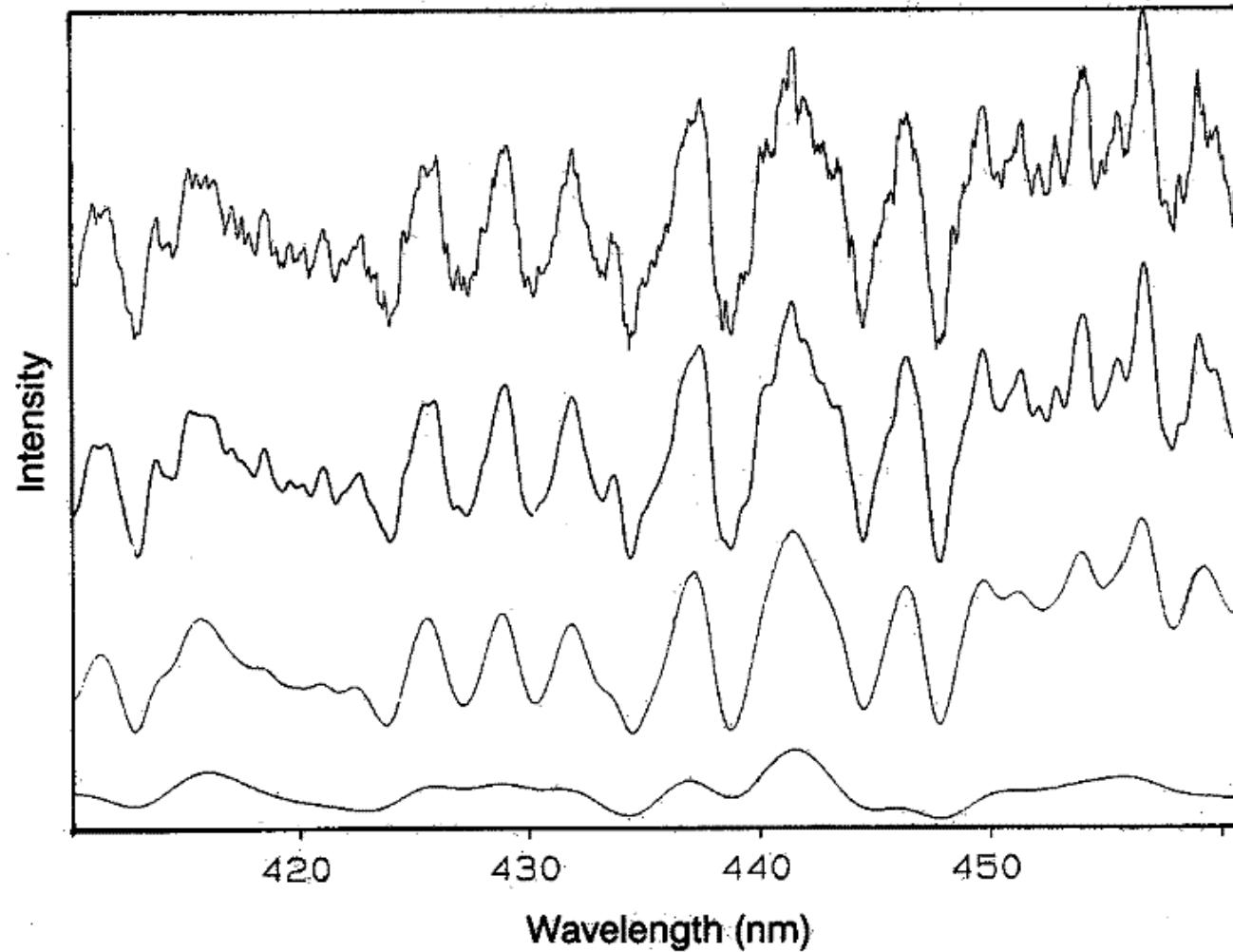


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Effect of Spectral Resolution - Key to Absolute Calibration (Example: NO₂)



Spectral
Resolution:

0.01 nm

0.3 nm

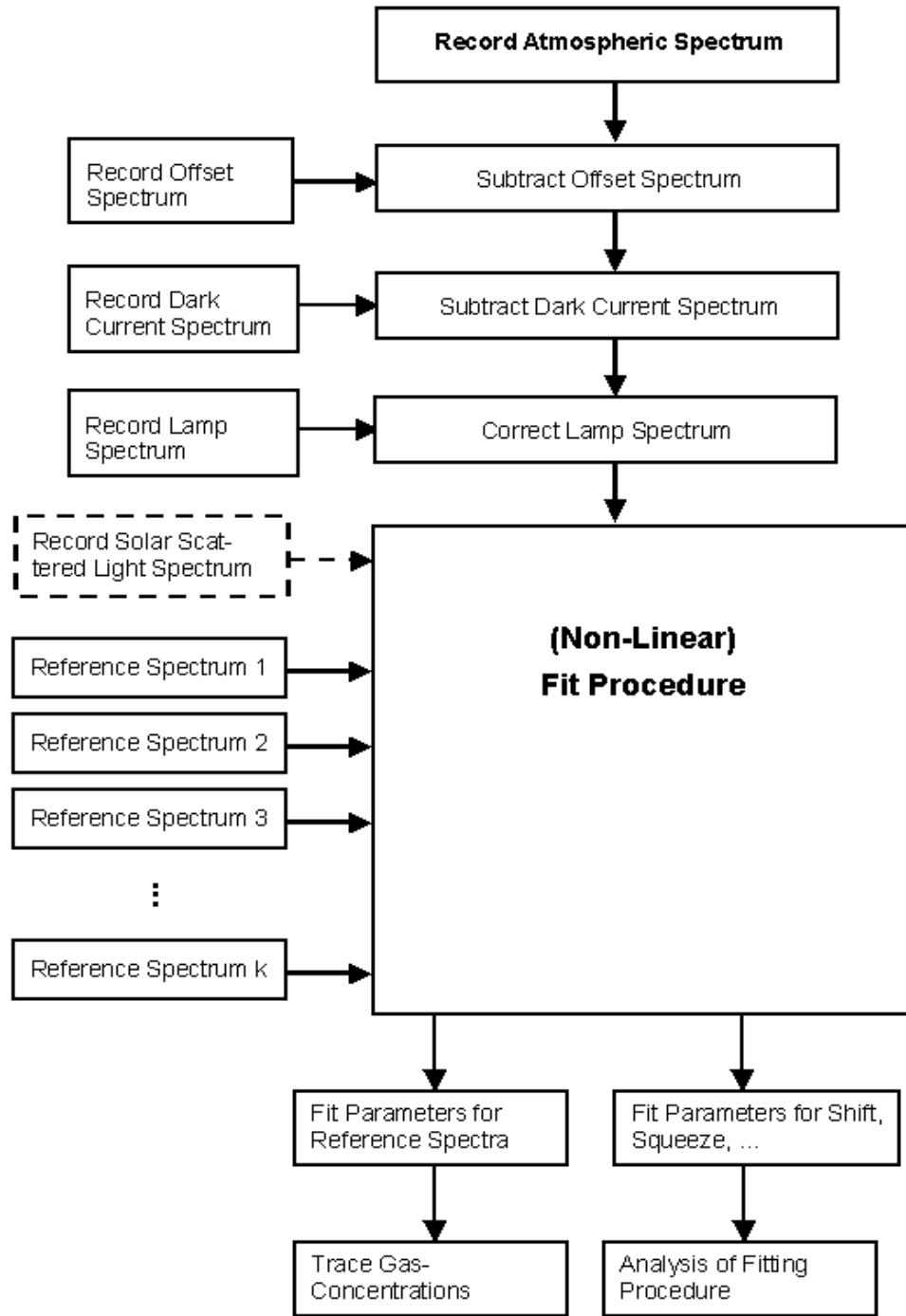
1.0 nm

3.0 nm

Diff. Abs.
Cross sect.



Flow Chart of a typical Active DOAS Evaluation



DOAS Evaluation Procedures (1)

$$I'(\lambda) = I'_0(\lambda) \cdot e^{-L \cdot \sum_j \sigma'_j(\lambda) c_j - L \left[\sum_i \sigma_{0j}(\lambda) c_j + \varepsilon_R + \varepsilon_M \right]} \cdot A(\lambda)$$

$$D'_{\text{meas.}}(\lambda) = \ln \left(\frac{I'_0(\lambda)}{I'(\lambda)} \right) = \underbrace{L \cdot \sum_j \sigma'_j(\lambda) c_j}_{\text{Differential Part}} + \underbrace{L \cdot \left[\sum_j \sigma_{0j}(\lambda) c_j + \varepsilon_R + \varepsilon_M \right] + \ln A(\lambda)}_{\text{Continuous Part}}$$

$$D'_{\text{Fit}}(\lambda) = \sum_j a_j \sigma'_j(\lambda) + \sum_k b_k \lambda^k = \text{'Modelled Spektrum'}$$

Determine Coefficients a_i , b_i such, that:

$$\begin{aligned} \chi^2 &= \sum_i \left[D'_{\text{Fit}}(\lambda_i) - D'_{\text{meas.}}(\lambda_i) \right]^2 = \\ &= \sum_i \left[\sum_j a_j \sigma'_j(\lambda_i) + \sum_k b_k \lambda_i^k - D'_{\text{meas.}}(\lambda_i) \right]^2 = \min. \end{aligned}$$



DOAS Evaluation Procedures (2)

Coefficients a_j , b_j for minimum χ^2 can be calculated as long as D' is linear in a_j , b_k :

$$\frac{d\chi^2}{da_j} = \sum_i \frac{d}{da_j} \left[\sum_j a_j \sigma'_j(\lambda_i) + \sum_k b_k \lambda_i^k - D'_{\text{meas}}(\lambda_i) \right]^2 = 0$$

Since $\frac{dx_l}{da_j} = 0$ unless $l = j$

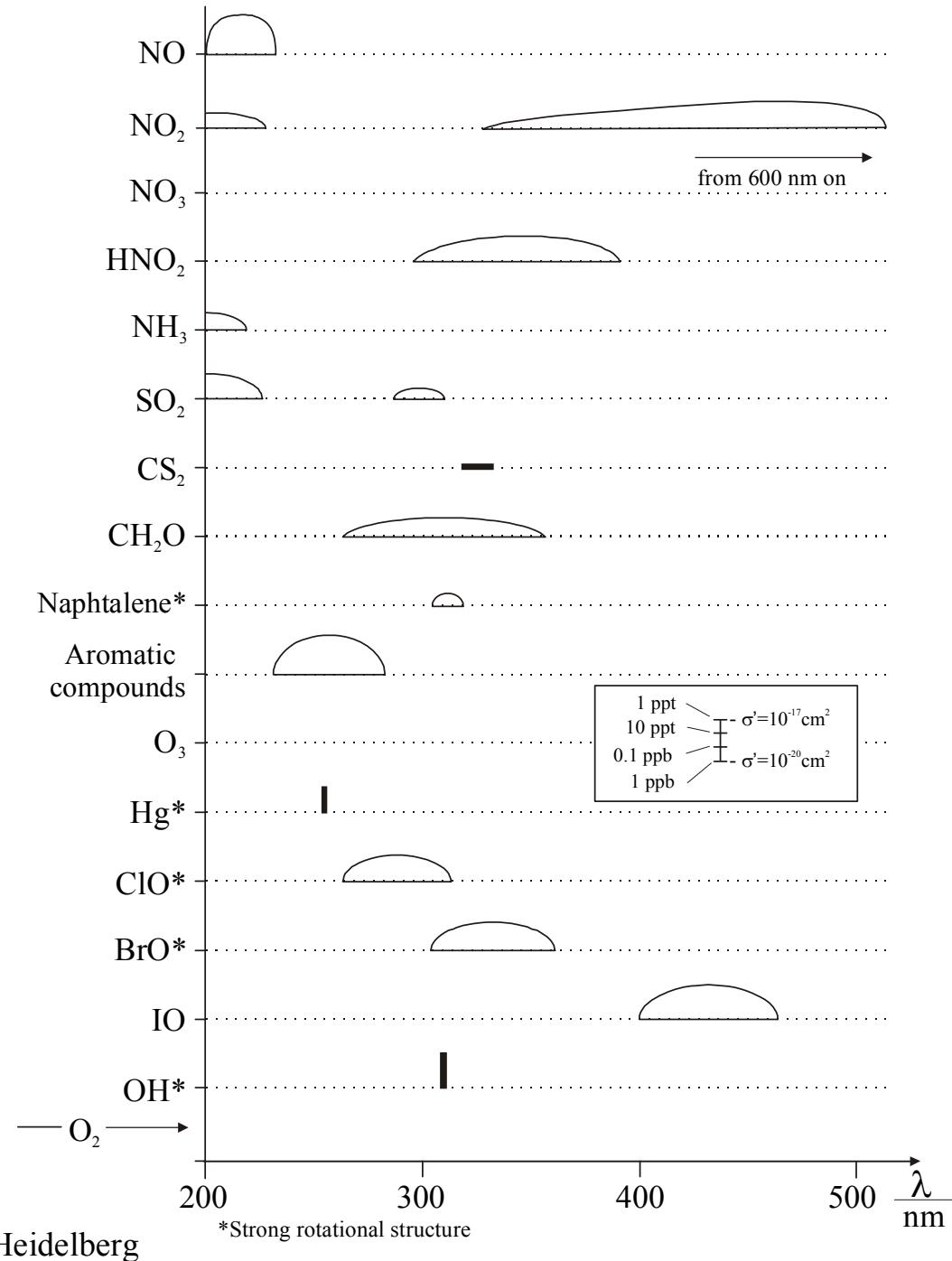
this leads to a system of N linear equations ($N = \text{No. of reference spectra } \sigma_j + \text{degree of polynomial } b^k \lambda_i^k$).

Only if „non-linear“ parameters, like shift in wavelength are included „fitting“ algorithms like Levenberg-Marquardt are required.



Atmospheric Species Measurable by DOAS

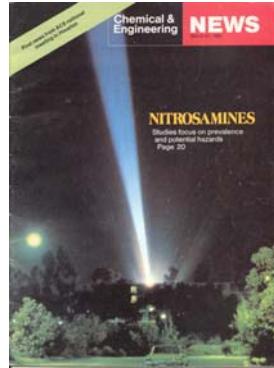
Many species can be detected at ppt levels
(mix. Ratio of 10^{-12}) or below



Some DOAS Characteristics

- Overlapping absorption structures due to different species can be separated
- Species not anticipated can be measured
- Warning against unexpected absorbers (residual)
- Immune against continuous (broad band) extinction due to e.g. aerosol or molecules
- High sensitivity, since many trace gas lines (bands) are used.





Aktive vs. Passive DOAS

Active:

Artificial Light Source (Xe-arc, incandescent lamp, LASER, LED ...)

Advantages:

- Day and night operation
- Wavelengths below 300nm accessible
- Well defined light path (mostly)

Disadvantages:

- Complex set-up (bi-static design)
- 3rd dimension difficult
- aerosol measurements difficult

Passive:

Natural Light Source (Sun, Moon, Stars, ...)



Advantages:

- Simple set-up (mono-static design)
- 3rd dimension accessible
- aerosol measurements possible

Disadvantages:

- Daytime only operation
- Only wavelengths above 300nm accessible
- Ill defined light path

Active DOAS Spectroscopy in the Atmosphere

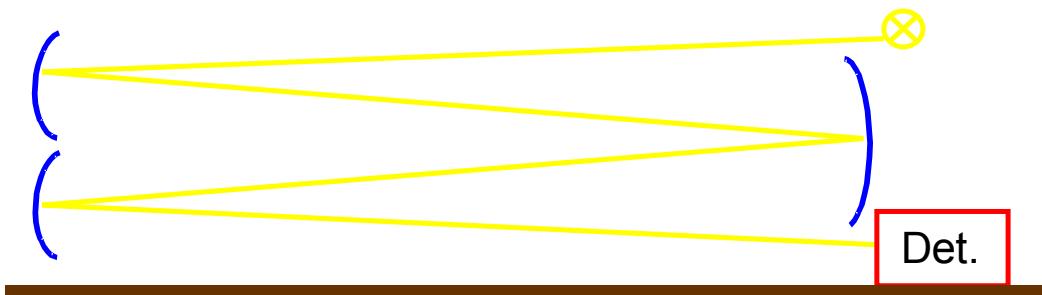
Active DOAS use artificial light source (e.g. Xe-arc)

1) Long-Path DOAS (LP-DOAS)



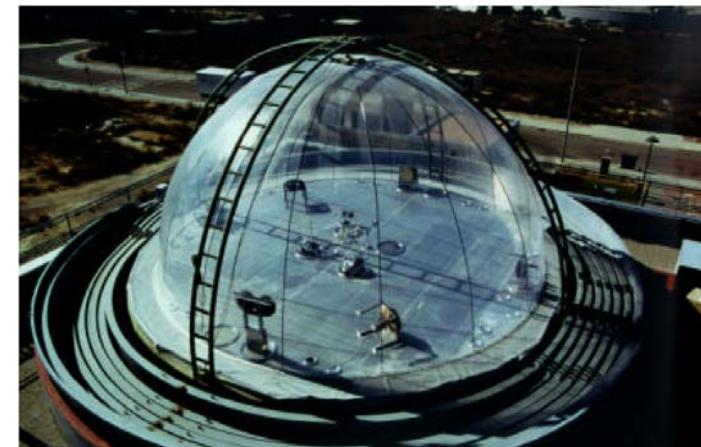
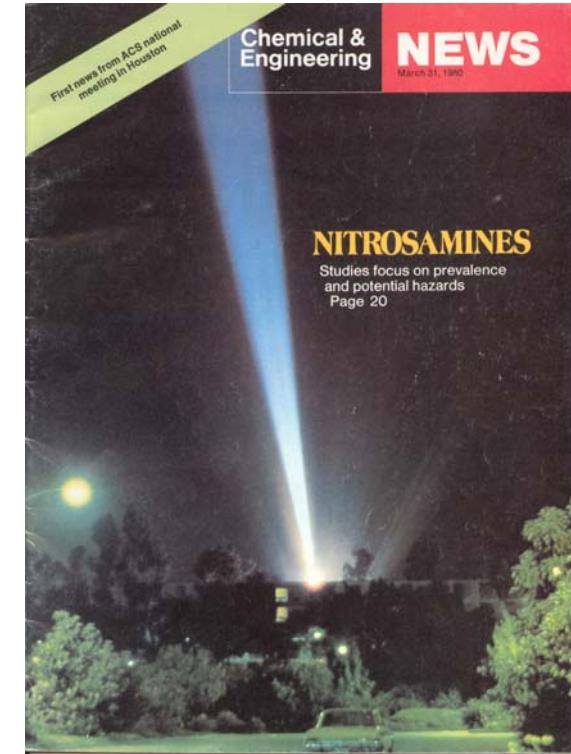
$$c = \frac{\ln \frac{I_0}{I}}{\sigma \cdot L}$$

2) Folded-Path DOAS



EUPHORE

SAPHIR



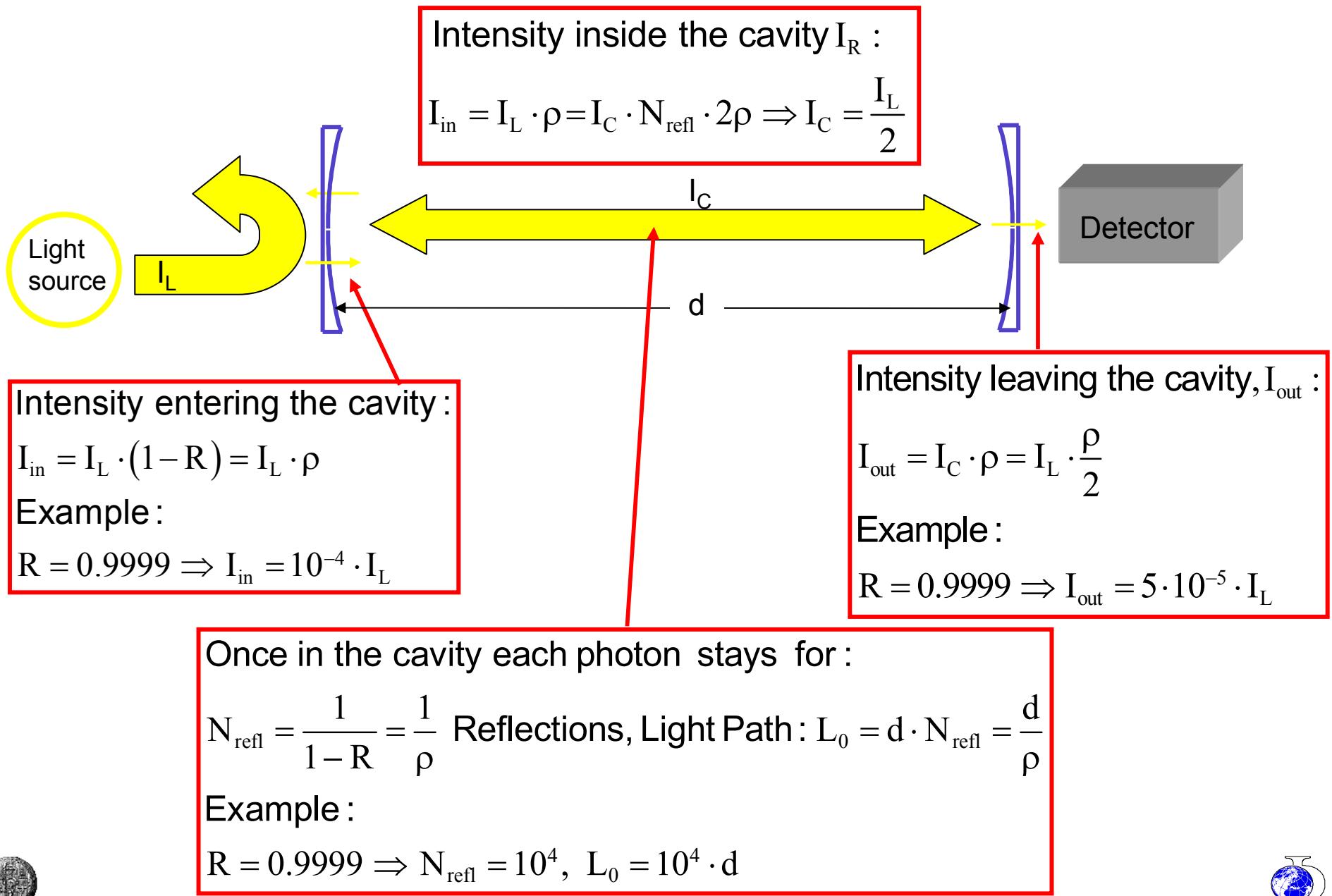
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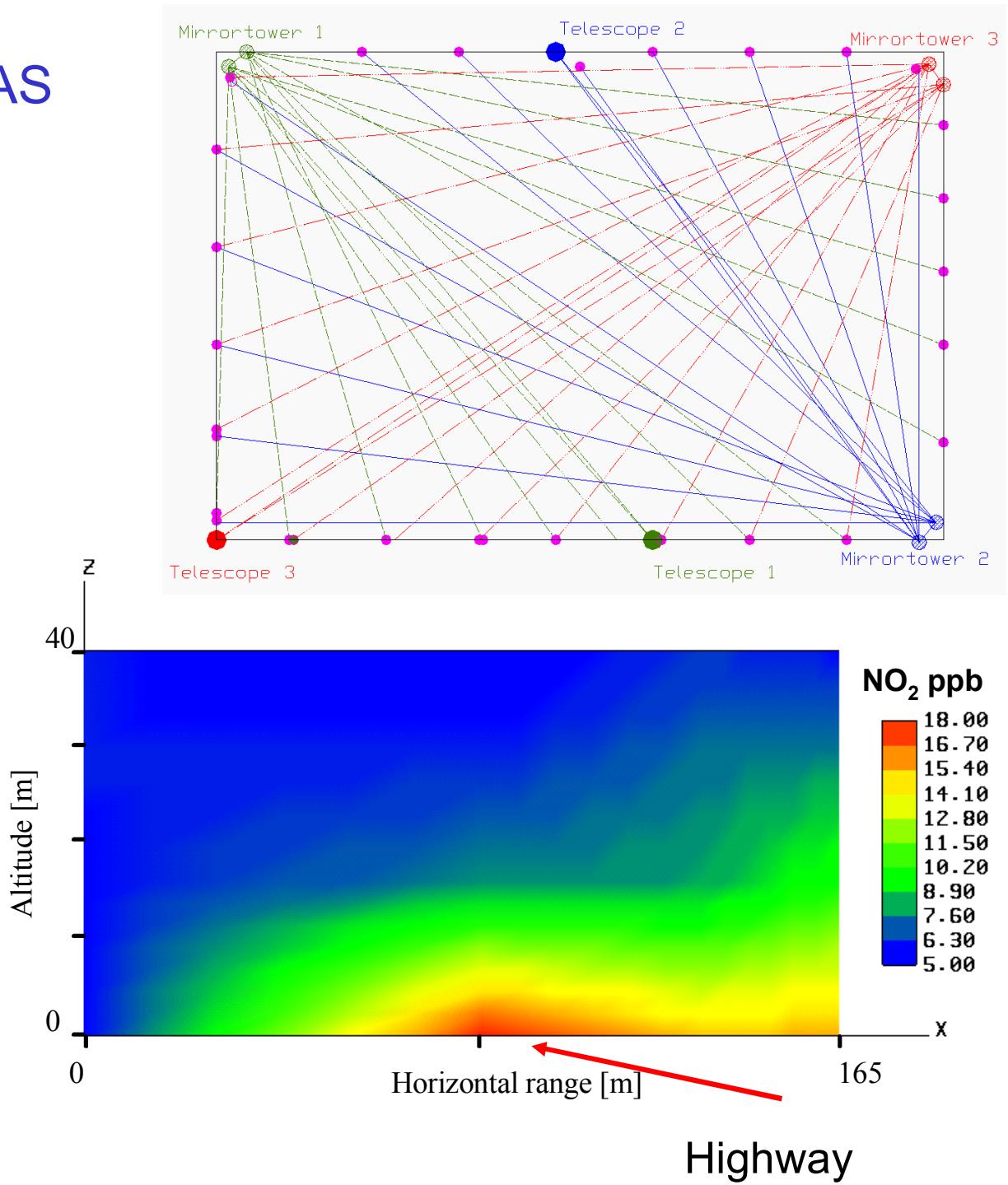
3) Cavity Enhanced DOAS



4) Tomographic DOAS

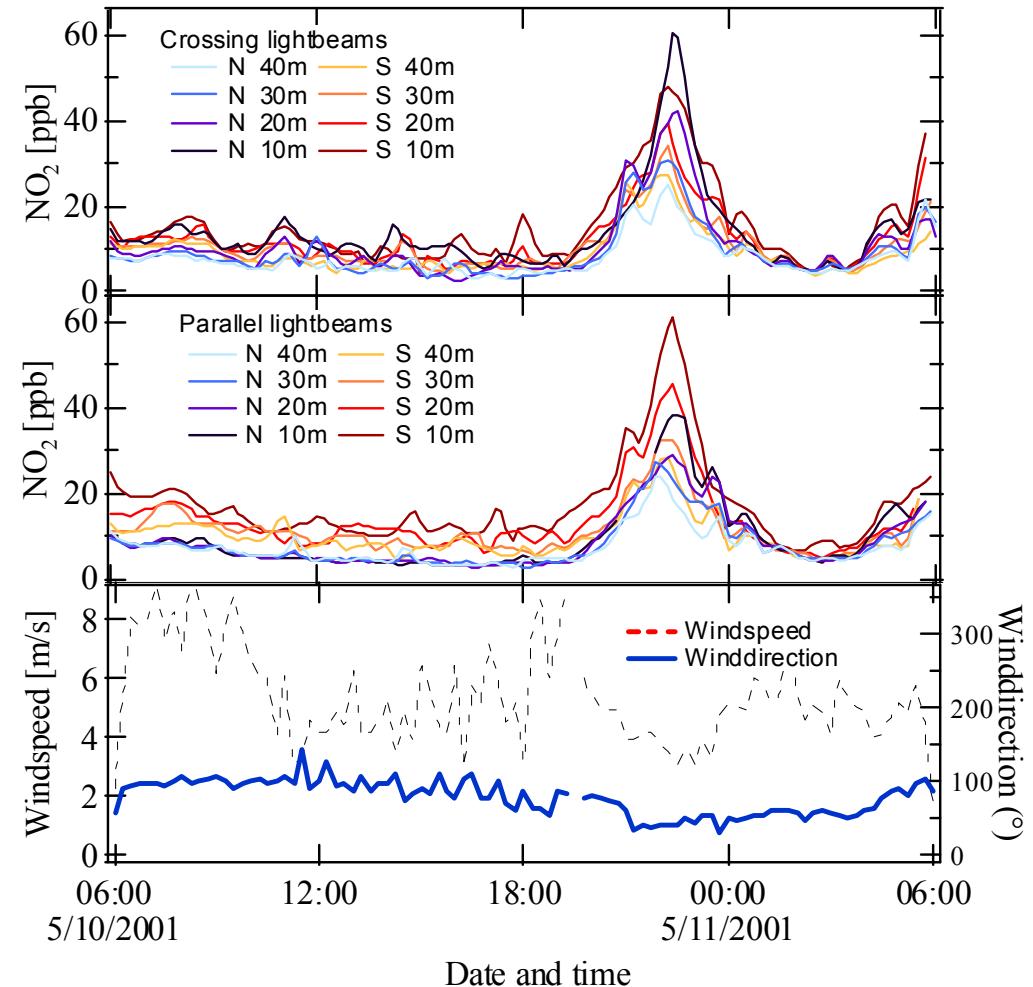
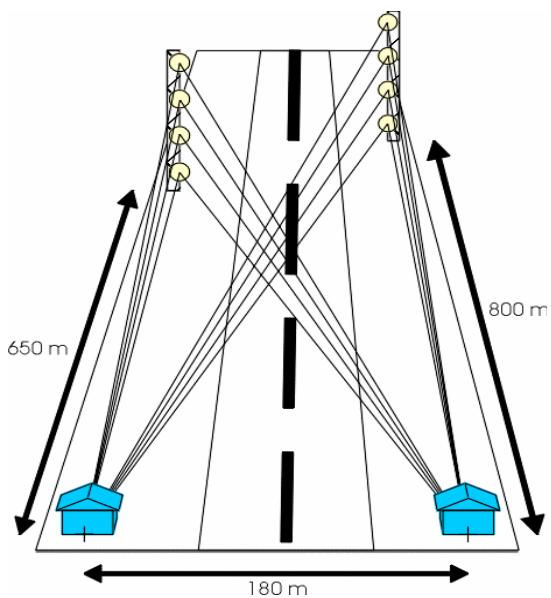
Examples:

- a) Tomographic indoor validation study
(Mettendorf et al. 2005)
- b) NO₂ Distribution at a Highway near Heidelberg
(Pundt et al. 2003)
- c) Distribution of trace gases in Heidelberg
(Pöhler 2009)



TOMO-DOAS at the Autobahn

Pundt et al. 2003



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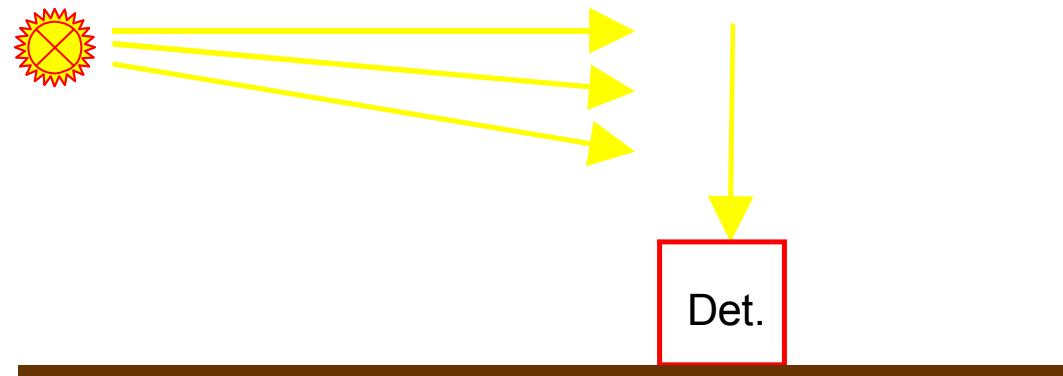
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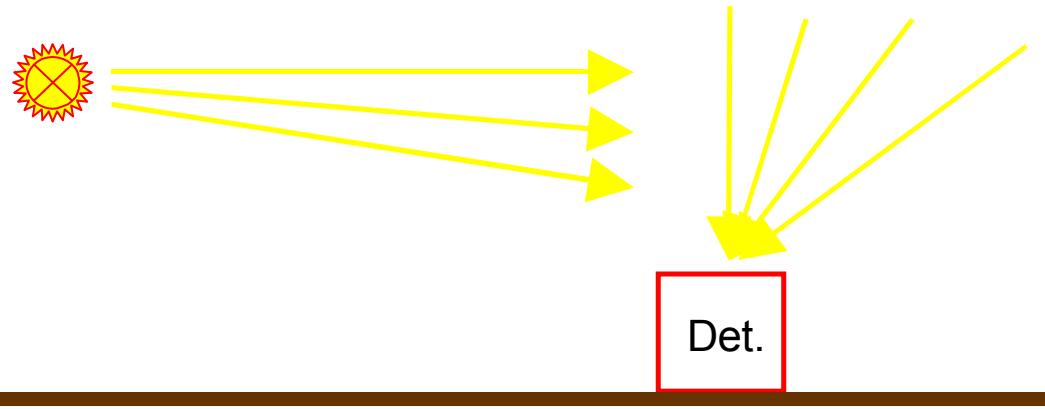
Passive DOAS Spectroscopy in the Atmosphere

Passive DOAS: Use **natural** light source (sun, moon, stars ...)

5) Zenith Scattered Light (ZSL-DOAS)



6) Multi Axis DOAS (MAX-DOAS)

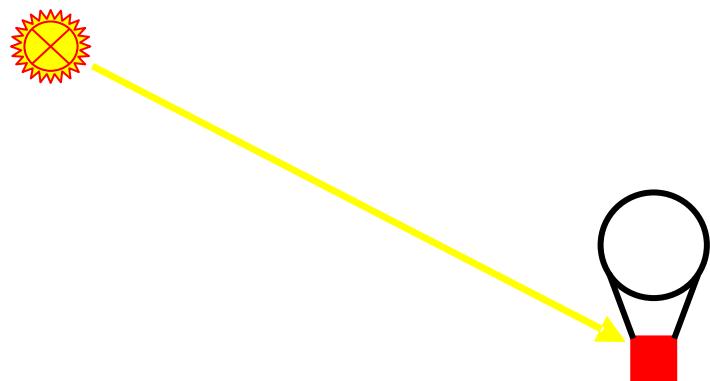


Passive DOAS Spectroscopy in the Atmosphere

7) Airborne Multi - Axis DOAS (AMAX-DOAS)



8) Balloon - Borne (direct sunlight) DOAS



LPMA/DOAS Gondola + Balloon



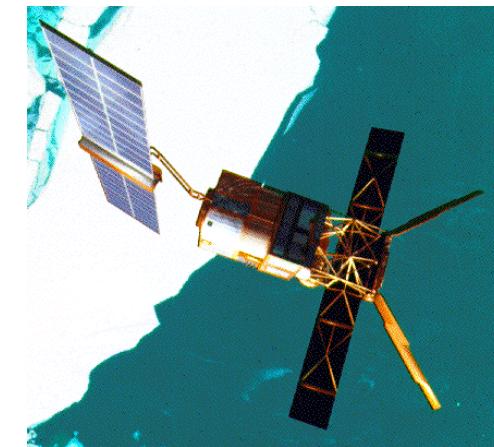
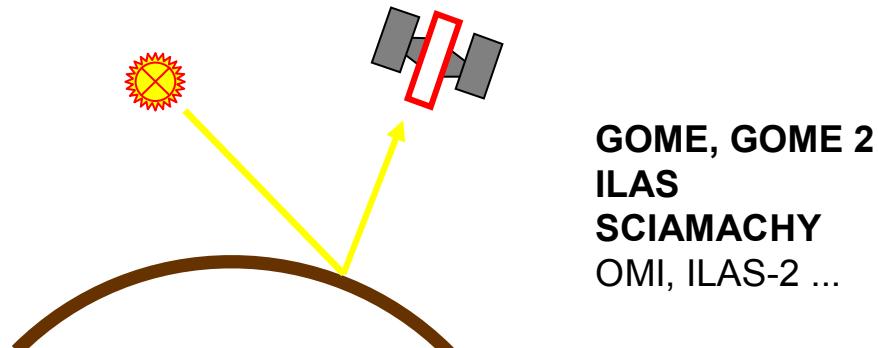
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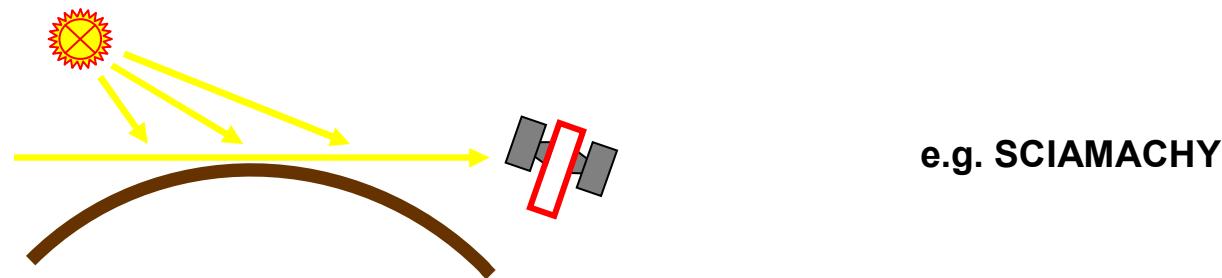


Passive DOAS Spectroscopy in the Atmosphere

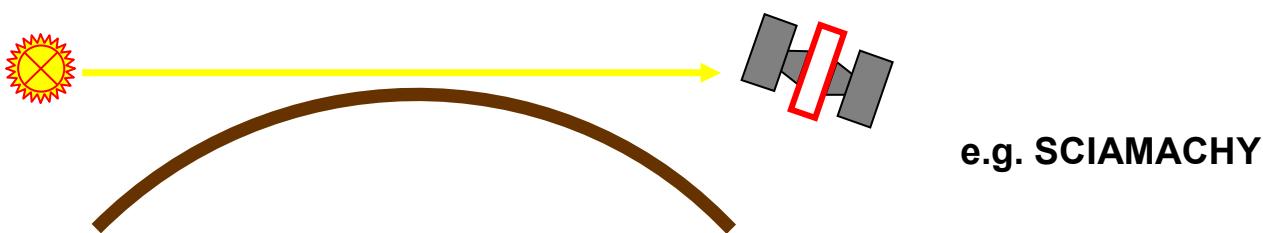
9) Satellite - Borne DOAS - Nadir Geometry



10) Satellite - Borne DOAS - Scattered Light Limb Geometry

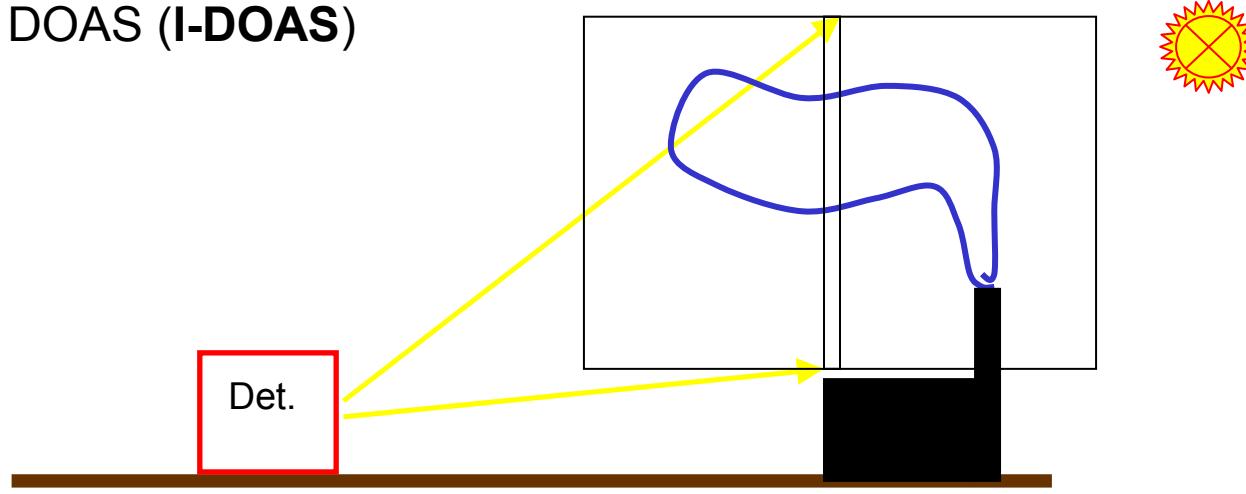


11) Satellite - Borne DOAS - Occultation

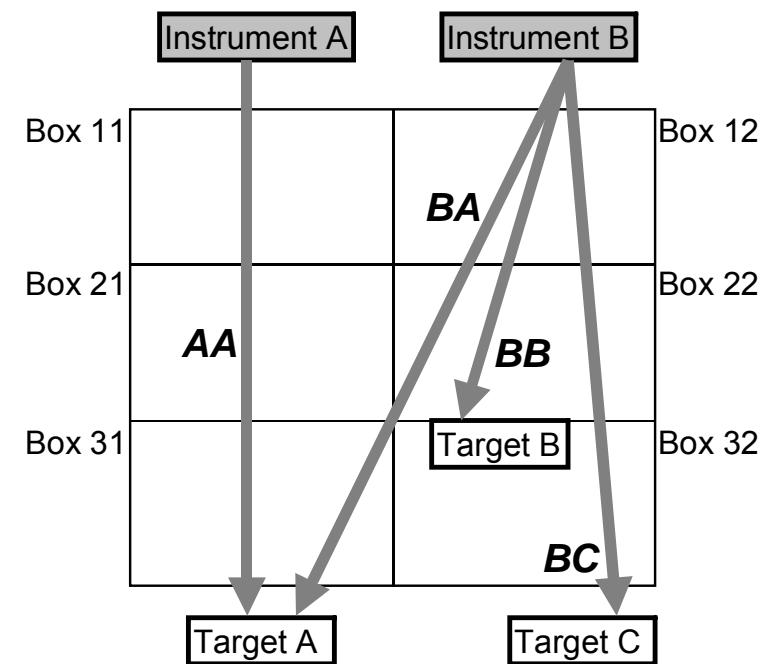
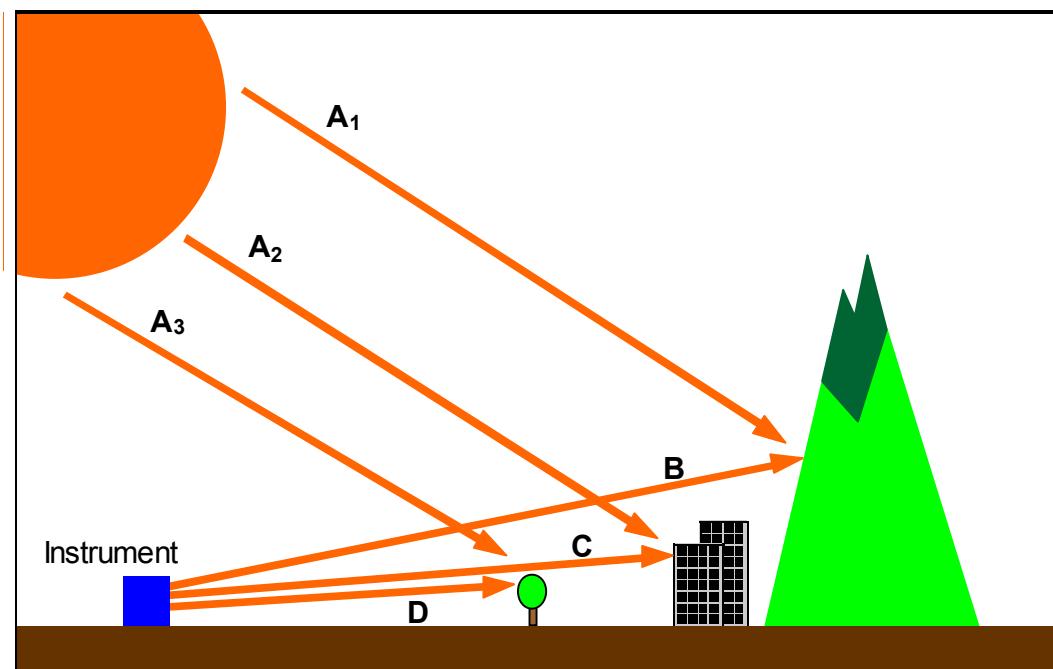


Passive DOAS Spectroscopy in the Atmosphere

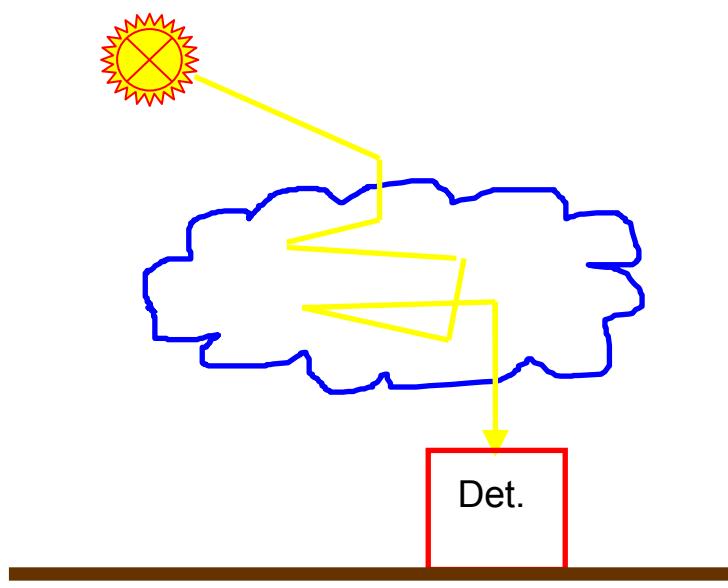
12) Imaging DOAS (I-DOAS)



13) Topographic Target Light Scattering - DOAS (ToTaL – DOAS)



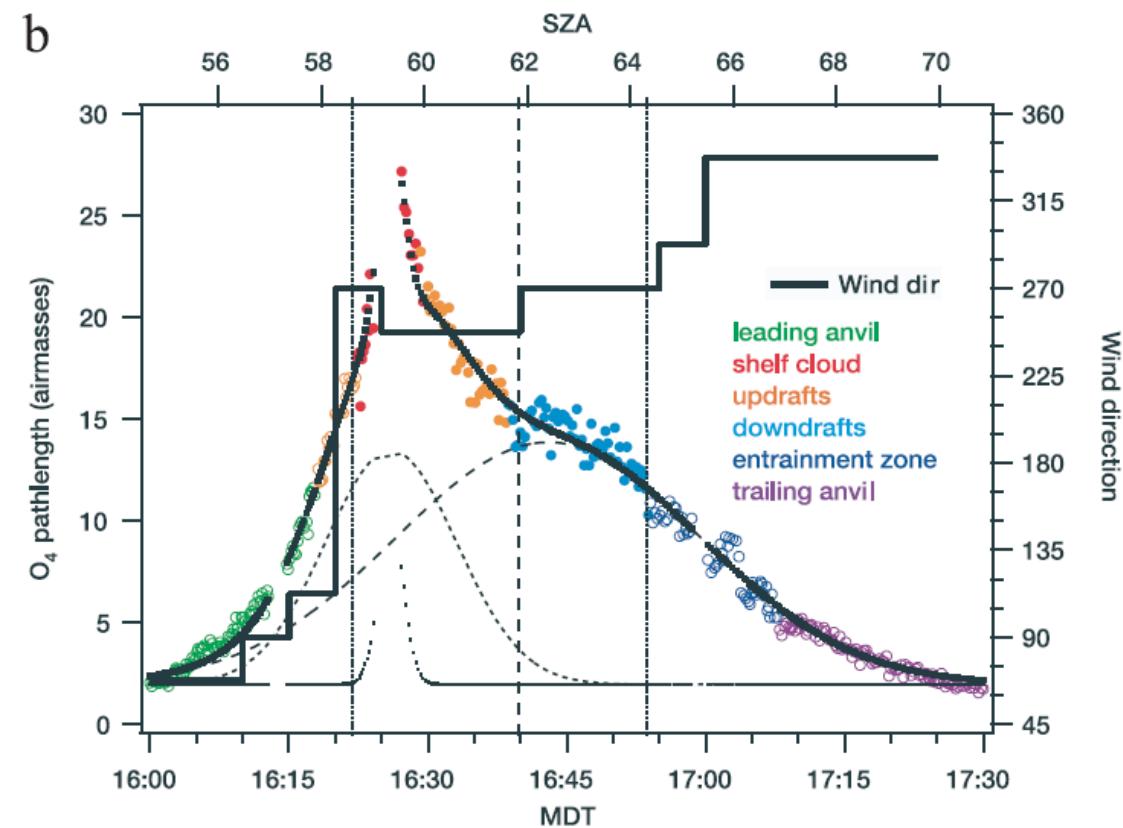
14) Determination of the Photon – Pathlength L (in Clouds) ‘inverse DOAS’



$$L = \frac{\ln I_0 / I}{\sigma \cdot c}$$



Langford, A.O. et al. 2004,
Spectroscopic measurements of NO₂
in a Colorado thunderstorm:
Determination of the mean production
by cloud-to-ground lightning flashes,
J. Geophys. Res. 109, D11304,
doi:10.1029/2003JD004158



Discoveries of Atmospheric Trace Gases by DOAS

OH (simultaneously with LIF?)

NO₃

HONO

O₄

OCIO strat.

CIO trop.

BrO strat. and trop.

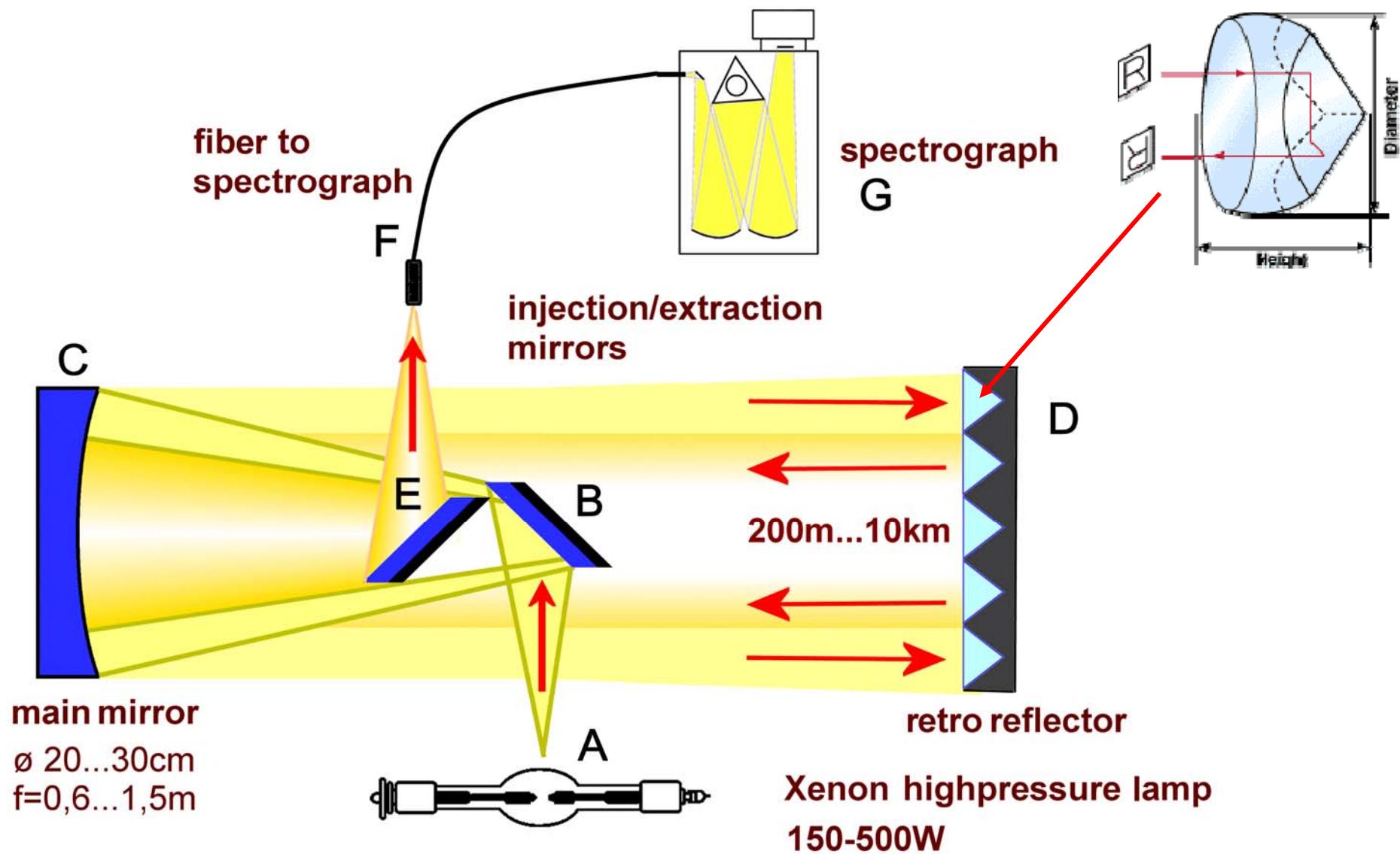
IO

OIO

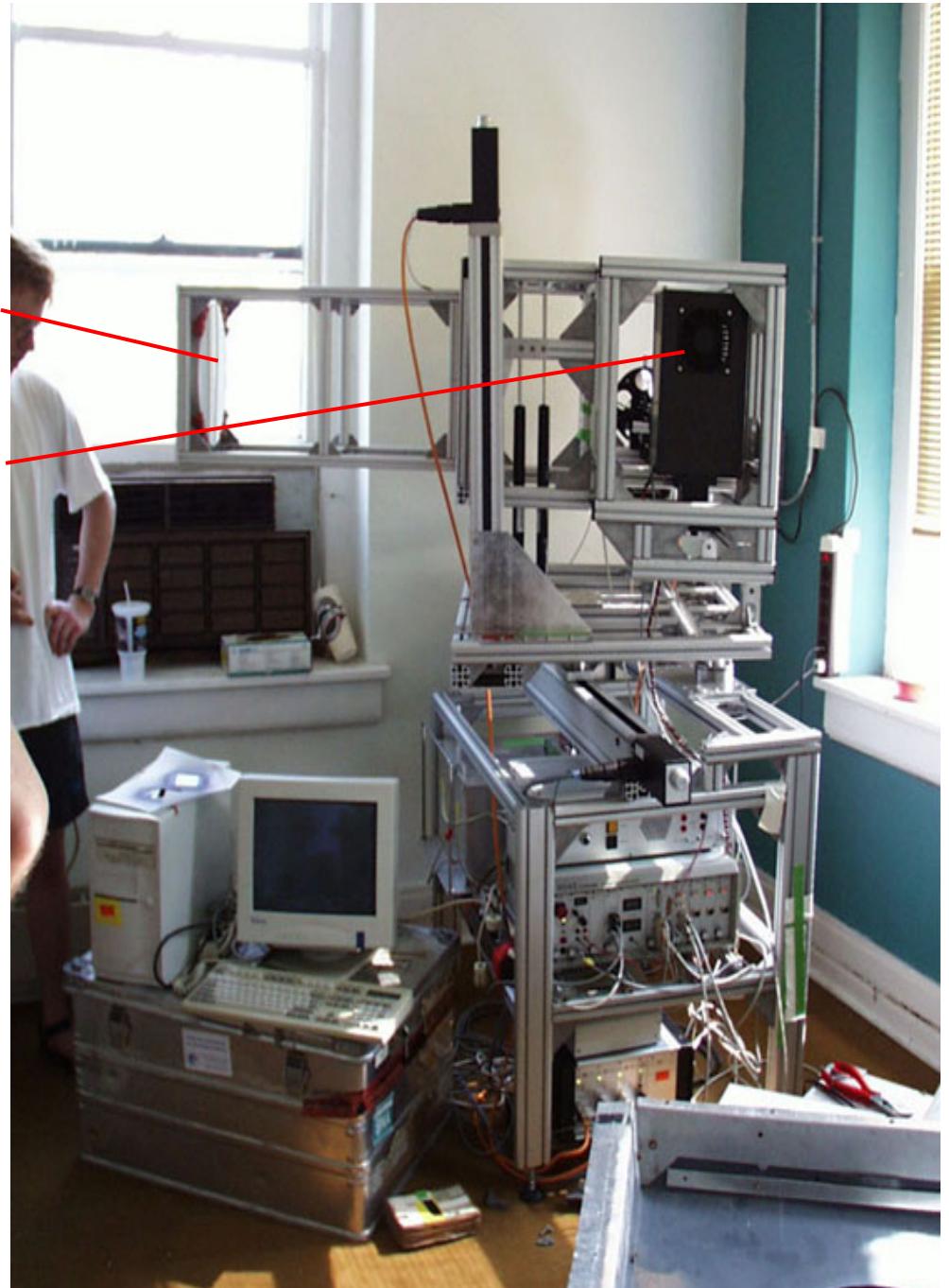
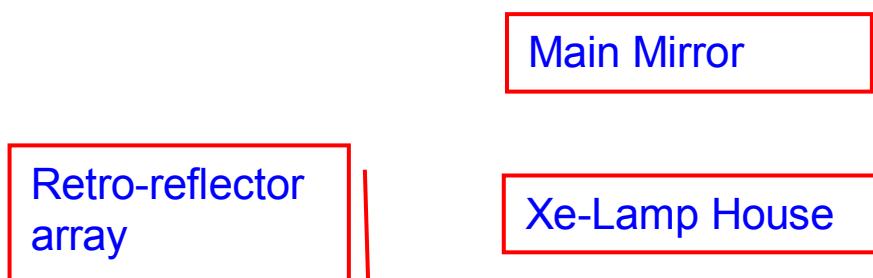
(H₂O)₂



Schematics of a „Coaxial“ Active DOAS Instrument



Components of an Active DOAS Instrument

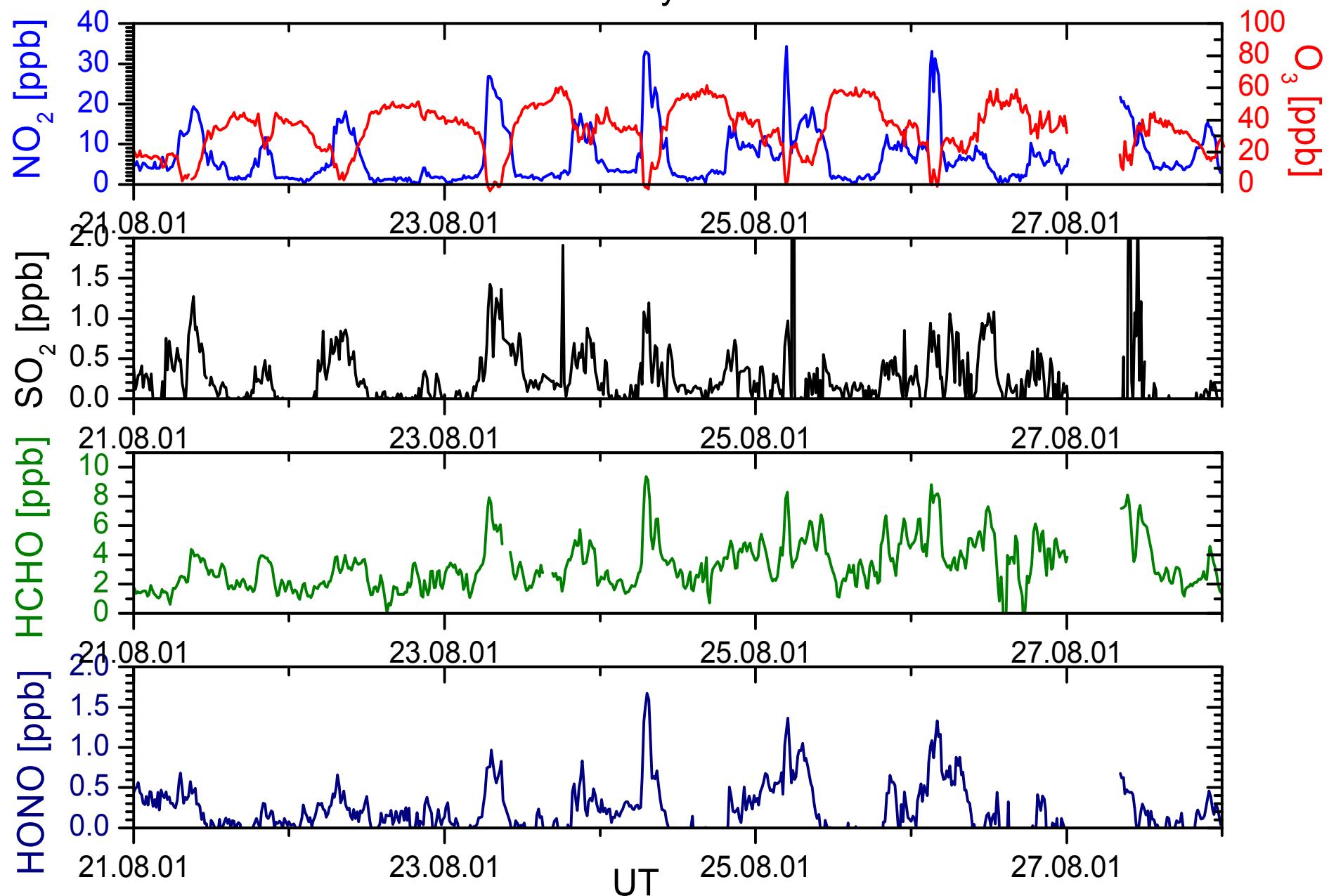


The Fiber-DOAS Instrument

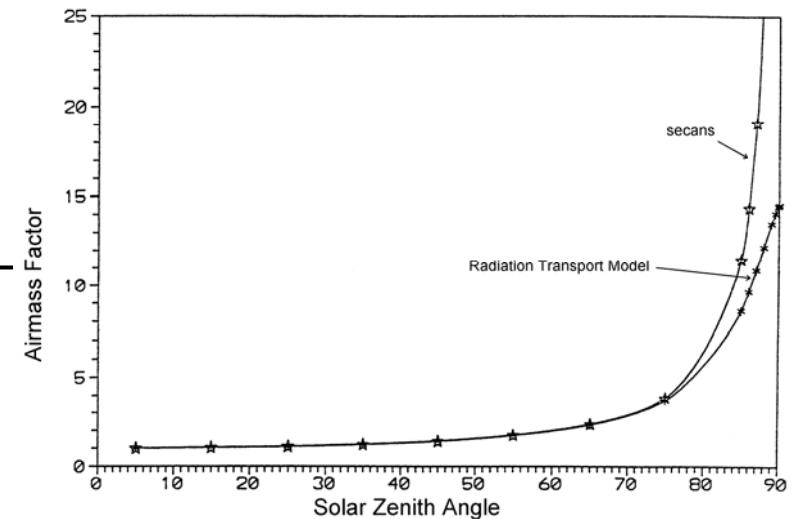
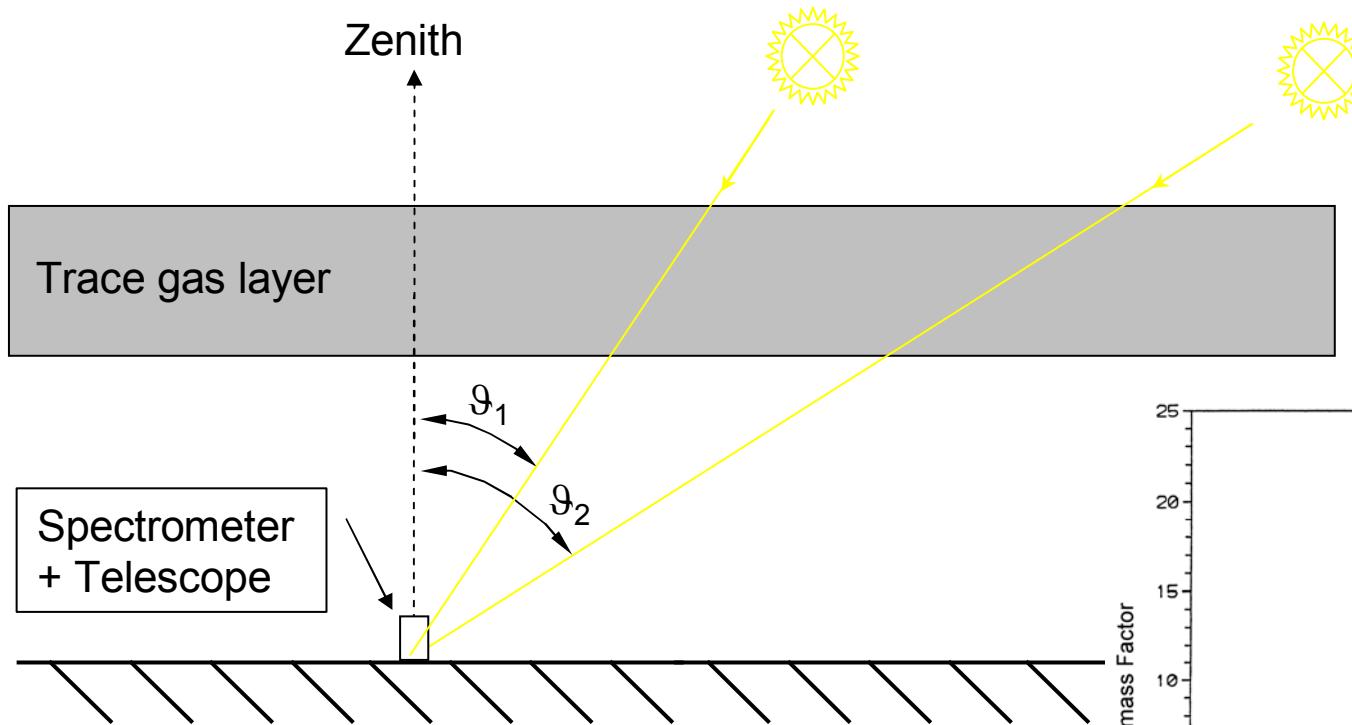
see Kern et al. 2008



Time Series of Pollutants in Heidelberg, 2001



Passive DOAS – Direct Radiation



„Slant“ Column Density S:

$$S = \int_0^{\infty} c(s) ds$$

Vertical Column Density V:

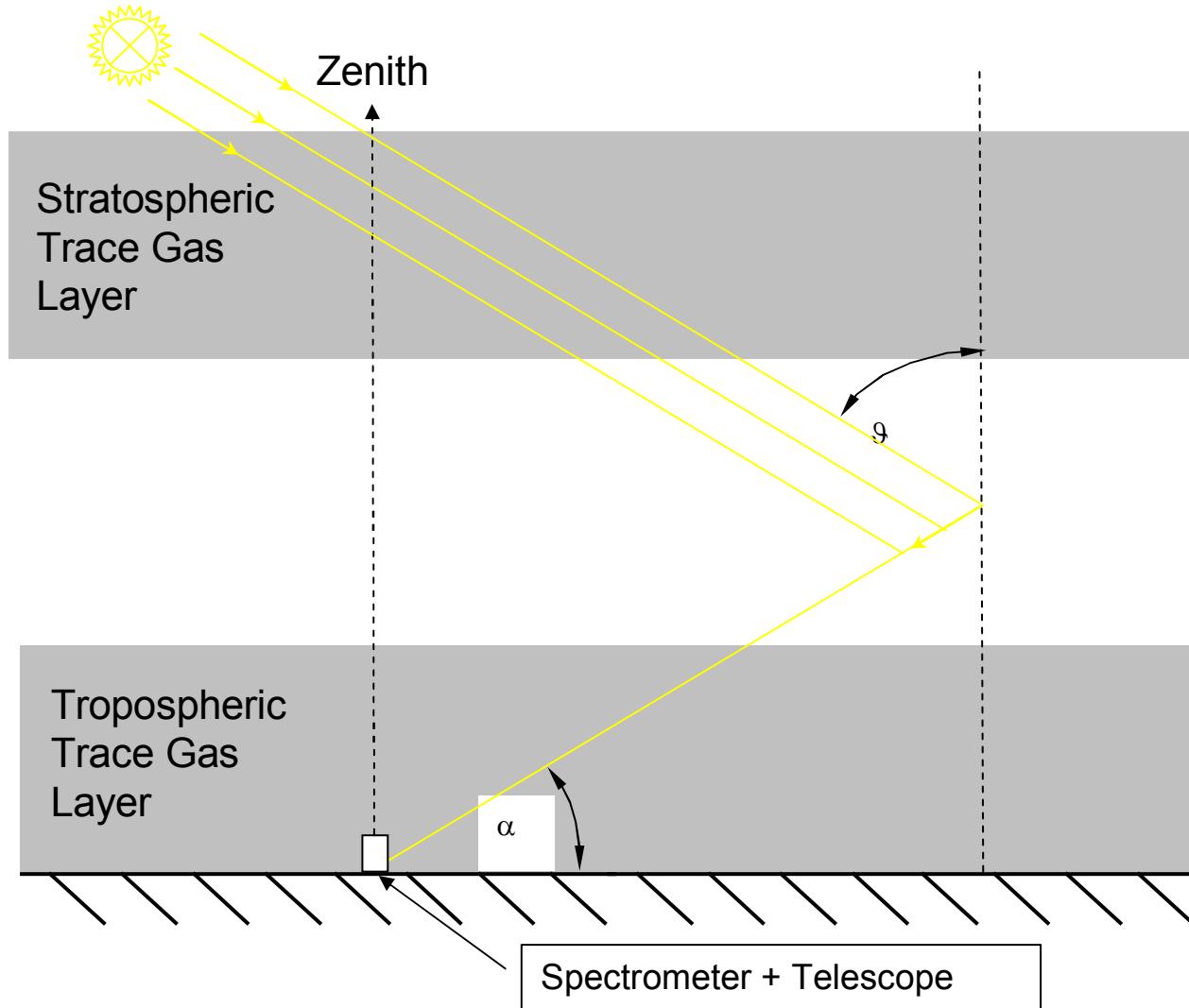
$$V = \int_0^{\infty} c(z) dz$$

Airmass Factor A:

$$A = \frac{S}{V}$$



Passive DOAS – Scattered Radiation Multi-AXis DOAS (MAX-DOAS)



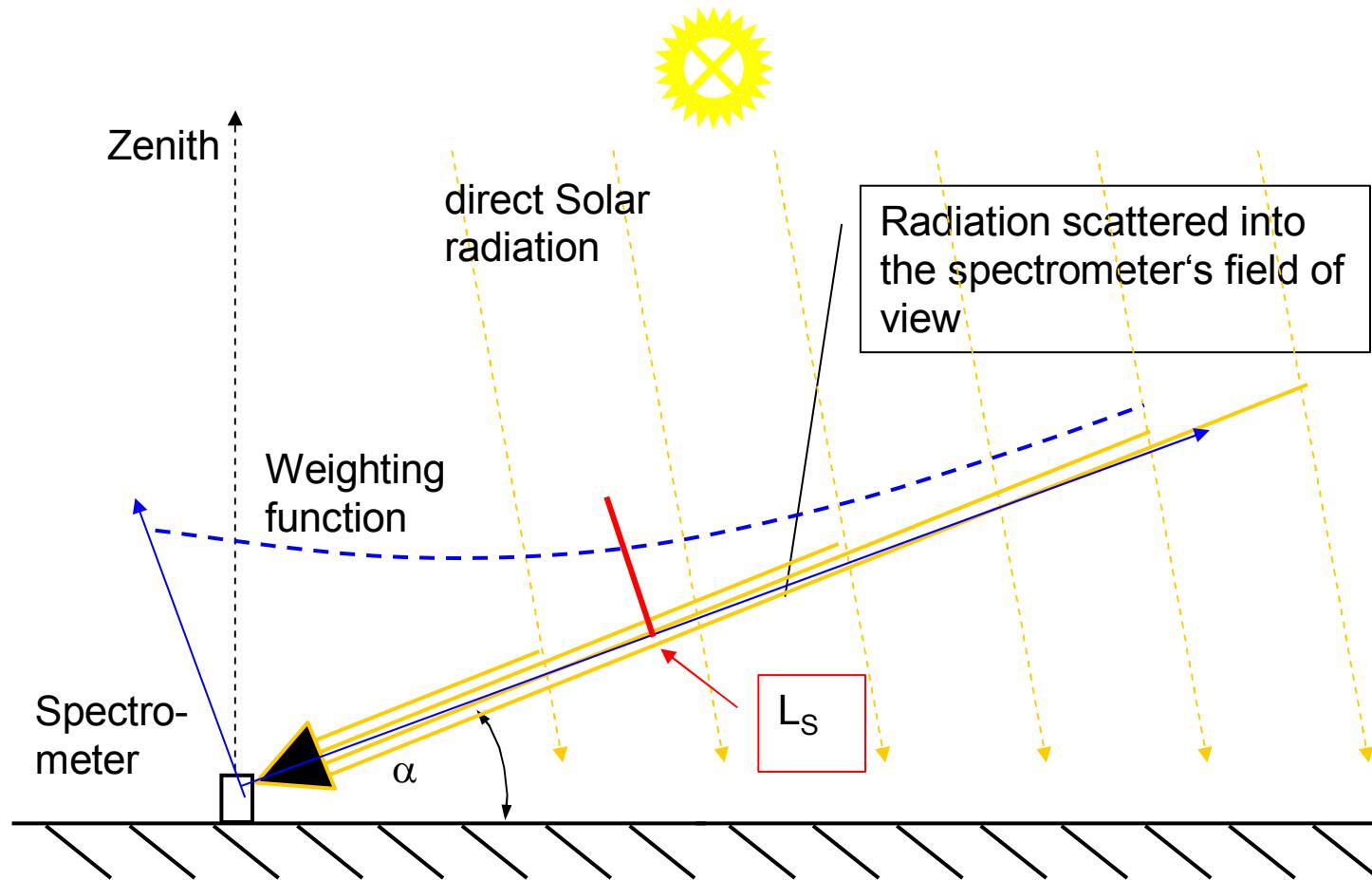
$$S = \int_0^s c(s') ds' = \int_0^s \frac{c(z)}{\cos(\vartheta)} dz = \frac{1}{\cos(\vartheta)} \cdot V$$

$$S = \int_0^s c(s') ds' = \int_0^s \frac{c(z)}{\sin(\alpha)} dz = \frac{1}{\sin(\alpha)} \cdot V$$

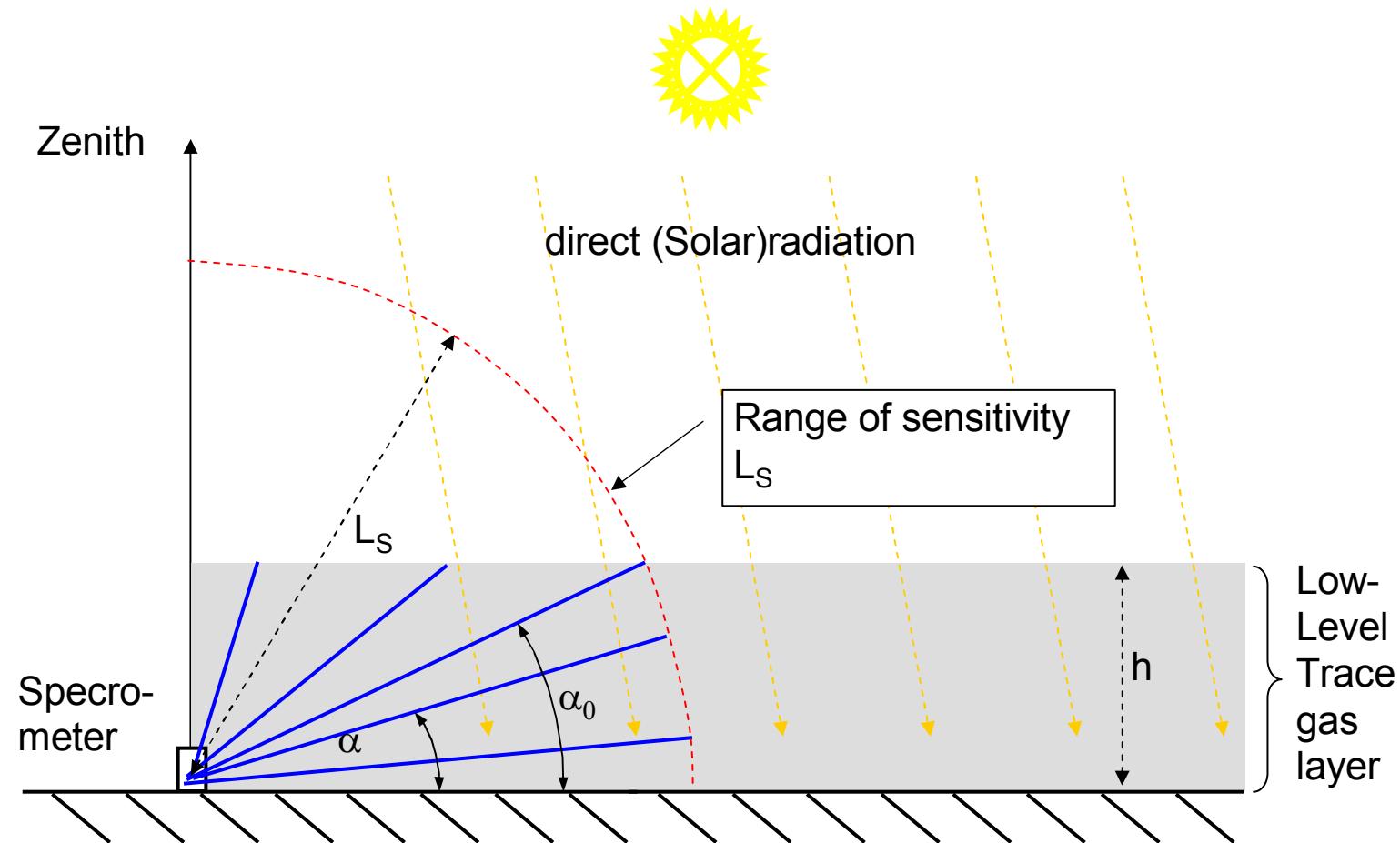
Remember
Demonstration by
Prof. Wagner ...



Passive DOAS – Scattered Radiation



MAX-DOAS: How to Determine Vertical Profiles

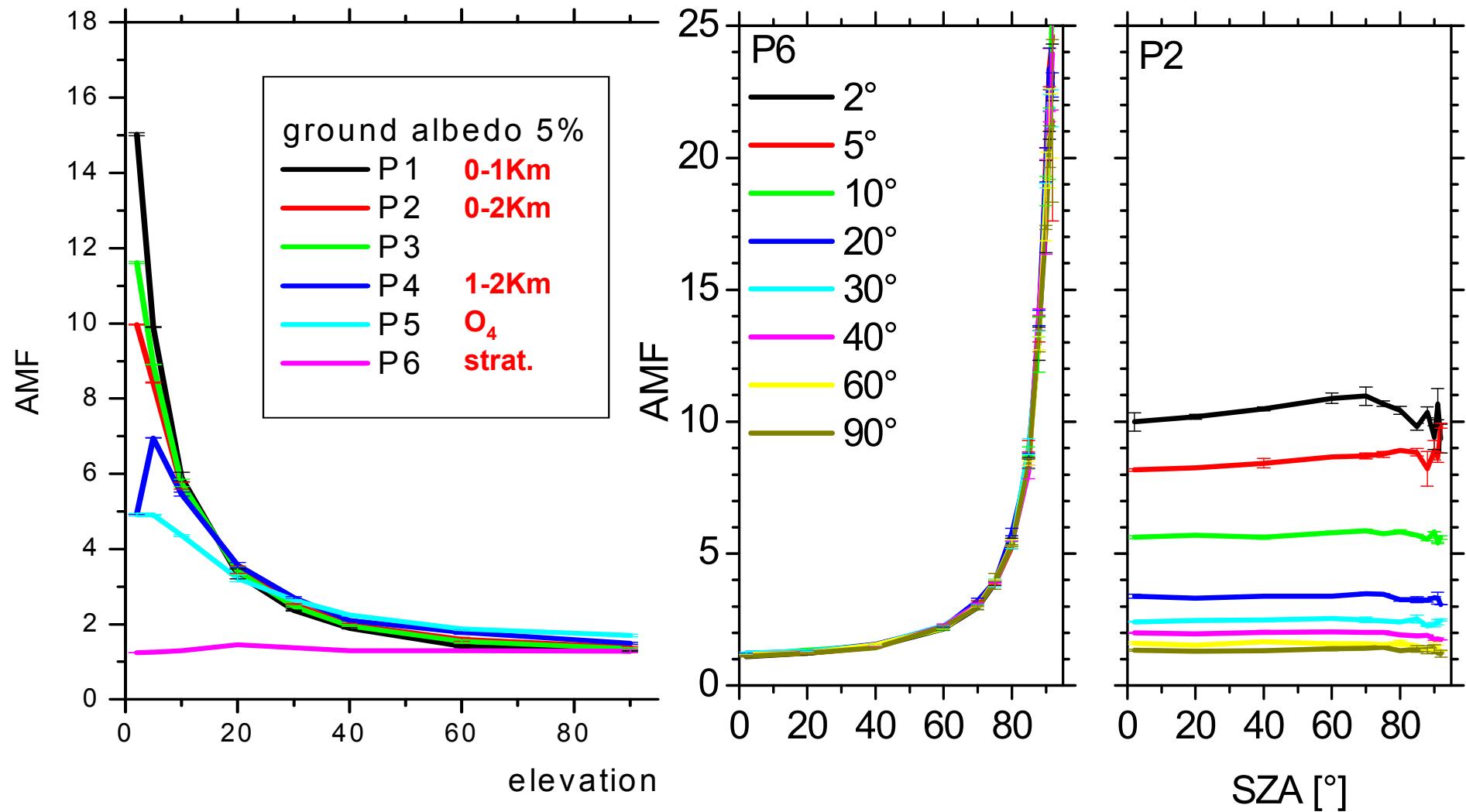


$$h = L_s \cdot \sin(\alpha_0)$$

$$c_j = \frac{S_j}{h}$$



AMF as a Function of Observation-Elevation and Solar-Zenith Angles



P1-P4: *Tropospheric Profiles*

P5: O_4

P6: *Strat. (25Km, 10Km FWHM)*

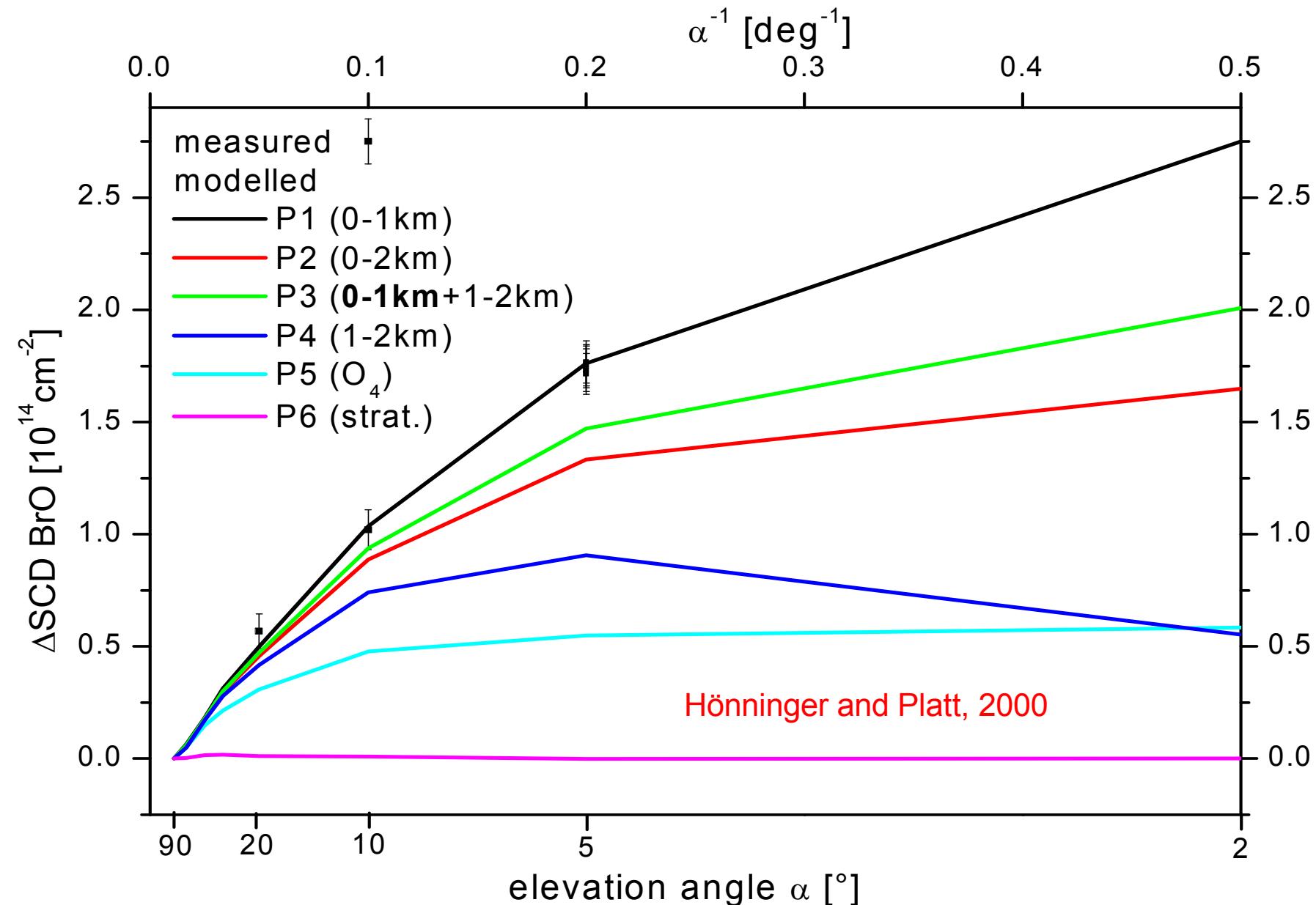
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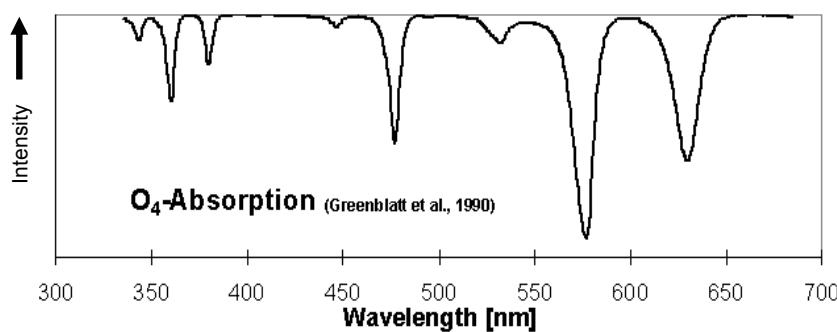
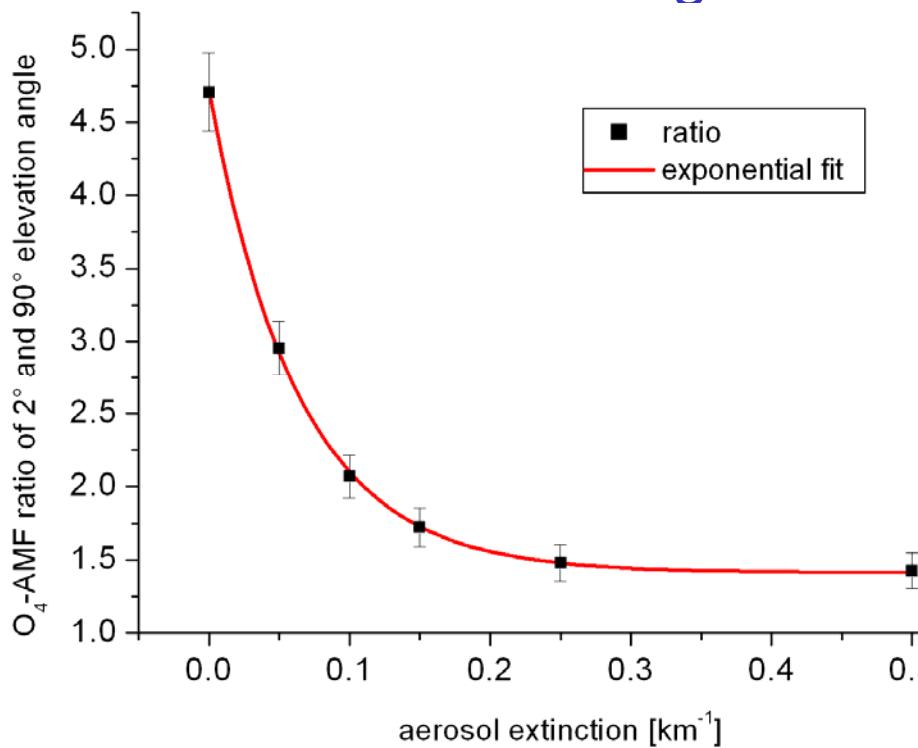
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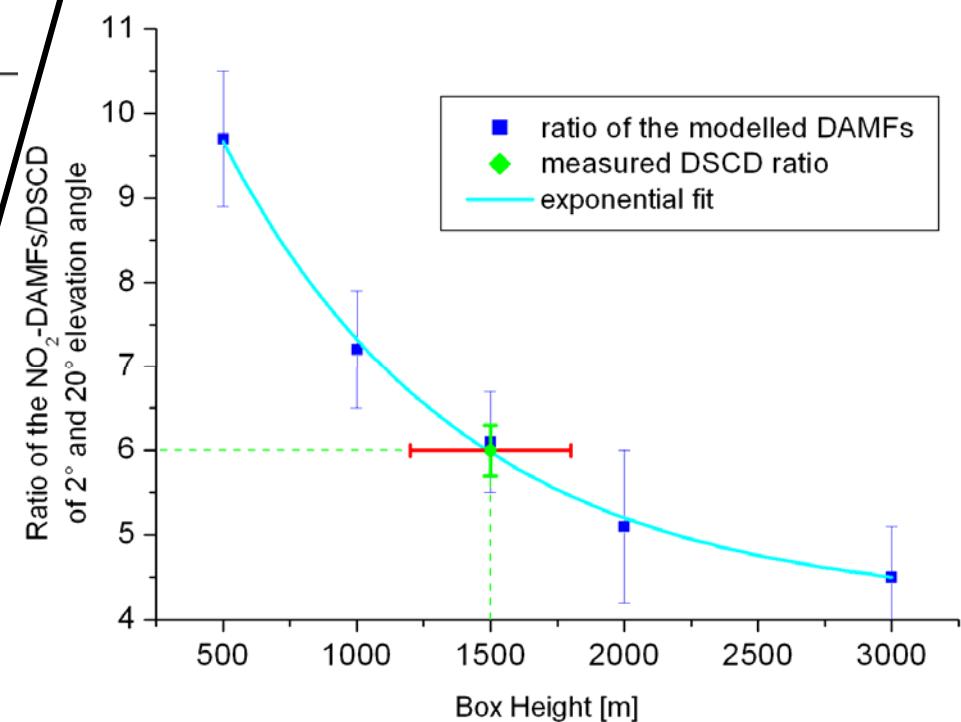
Vertical Resolution of MAX-DOAS → ‘Poor Man’s LIDAR’
ΔSCD’s Measured on May 4, 2000 (15:15 – 15:40 UT), Alert 2000 PSE



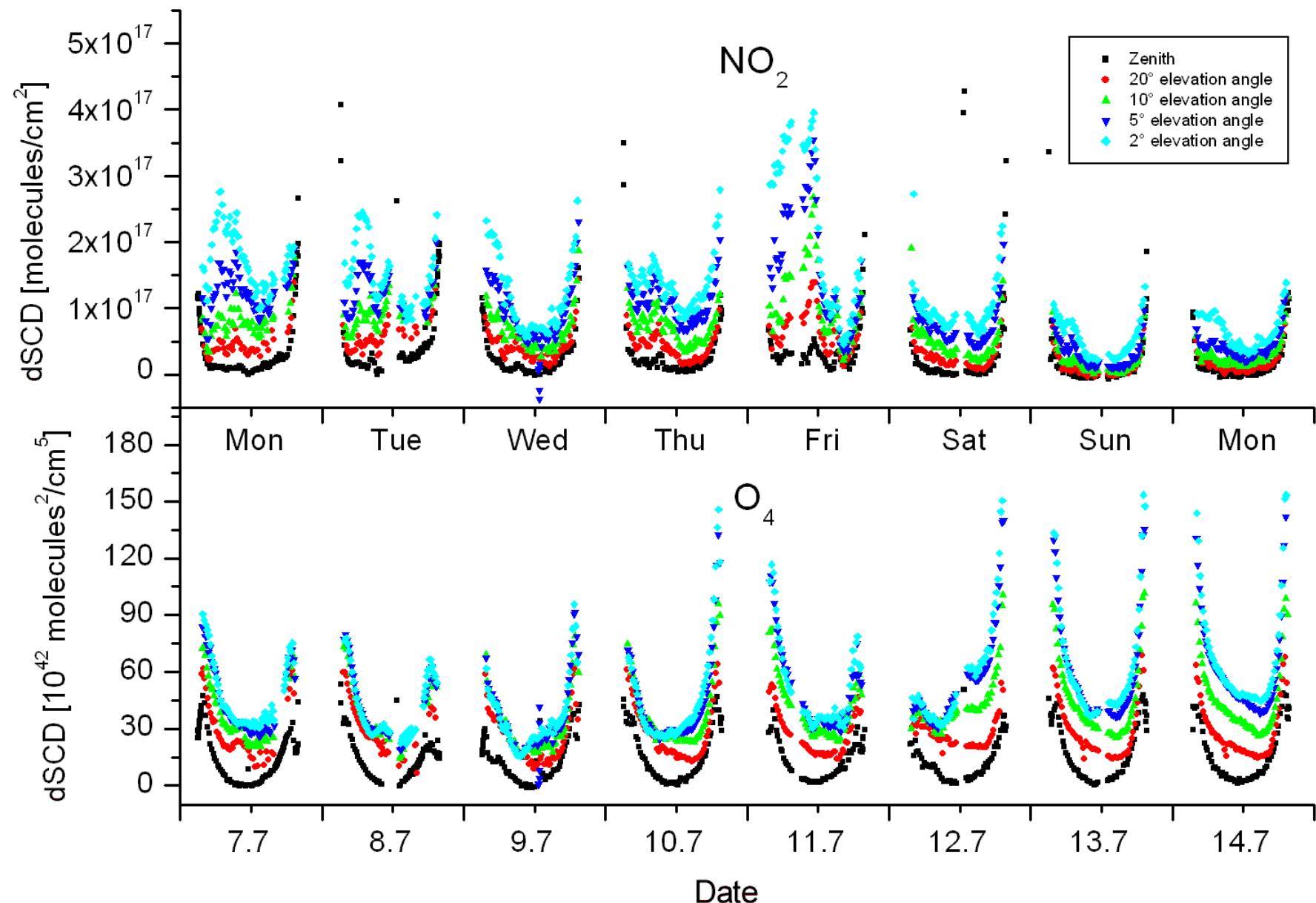
Ratios of the modelled O₄ and NO₂- Diff. AMF's at different elevation angles → Determine Layer Height



Ratios of the modelled NO₂- Diff. AMF's at 2° and 20° elevation angle --> Determine Layer Height



Slant column densities of NO_2 and O_4 in Heidelberg, July 2003



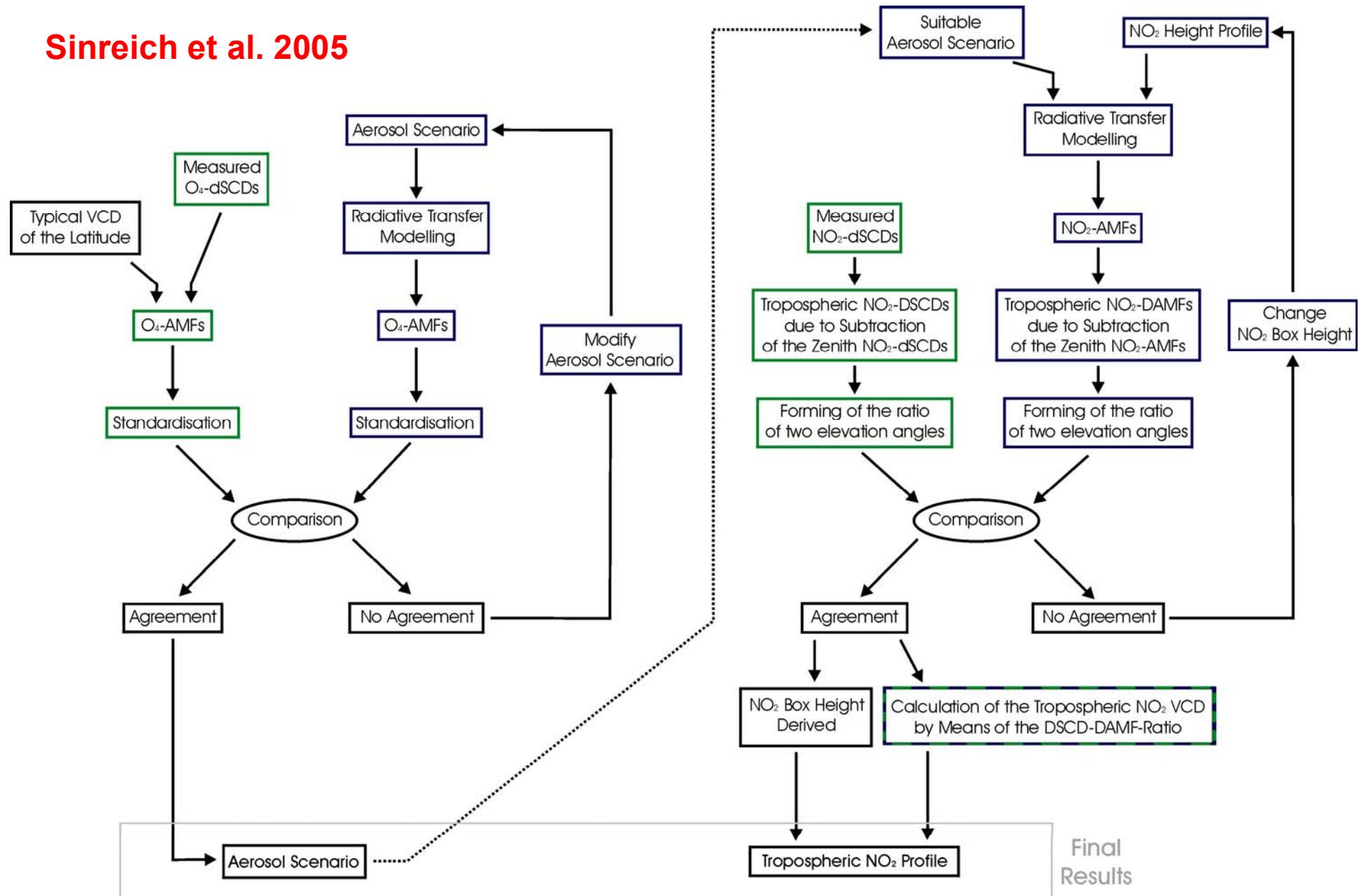
Reference spectrum of July 6 of at noon

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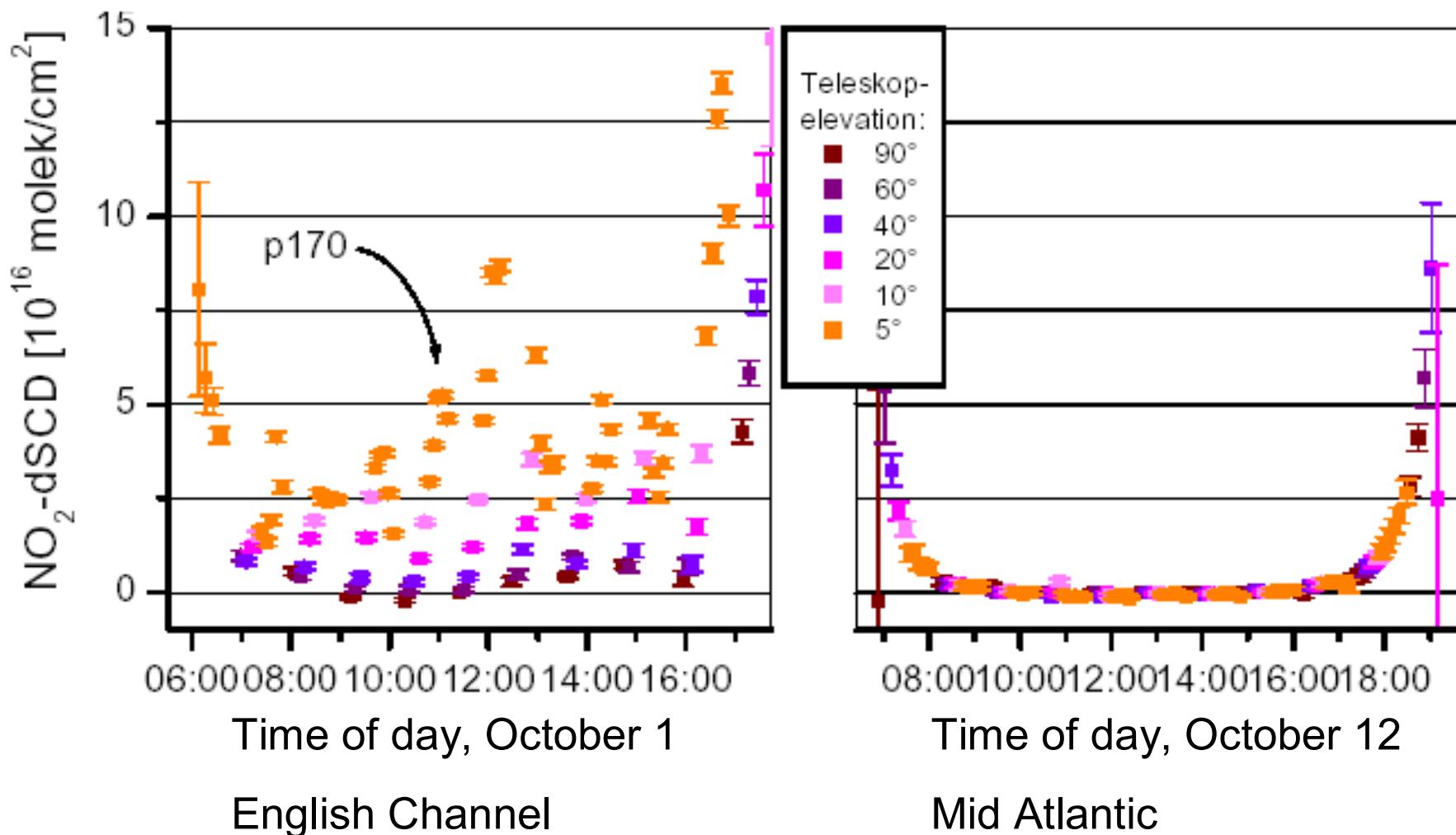


Flow Chart of MAX-DOAS Trace Gas + Aerosol Evaluation Procedure

Sinreich et al. 2005



MAX-DOAS Measurements of NO₂ from Polarstern 2000



Leser et al. GRL 2003
doi:10.1029/2002GL015811

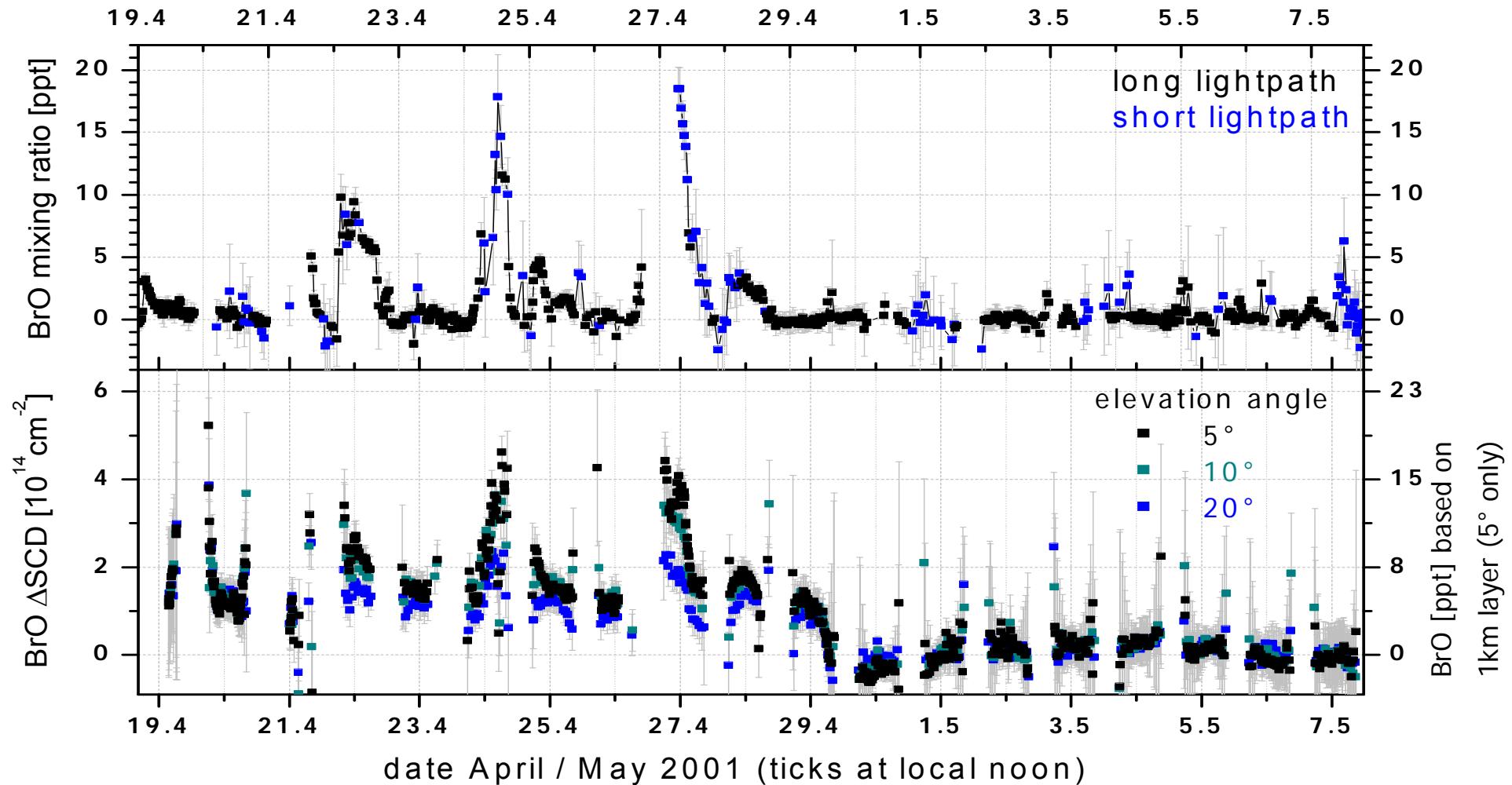


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Hudson Bay 2001: Comparison of BrO Measurements by active LP-DOAS and passive MAX-DOAS



Hönniger and Platt, Atmos. Environ. 36, 2481-2489, 2002



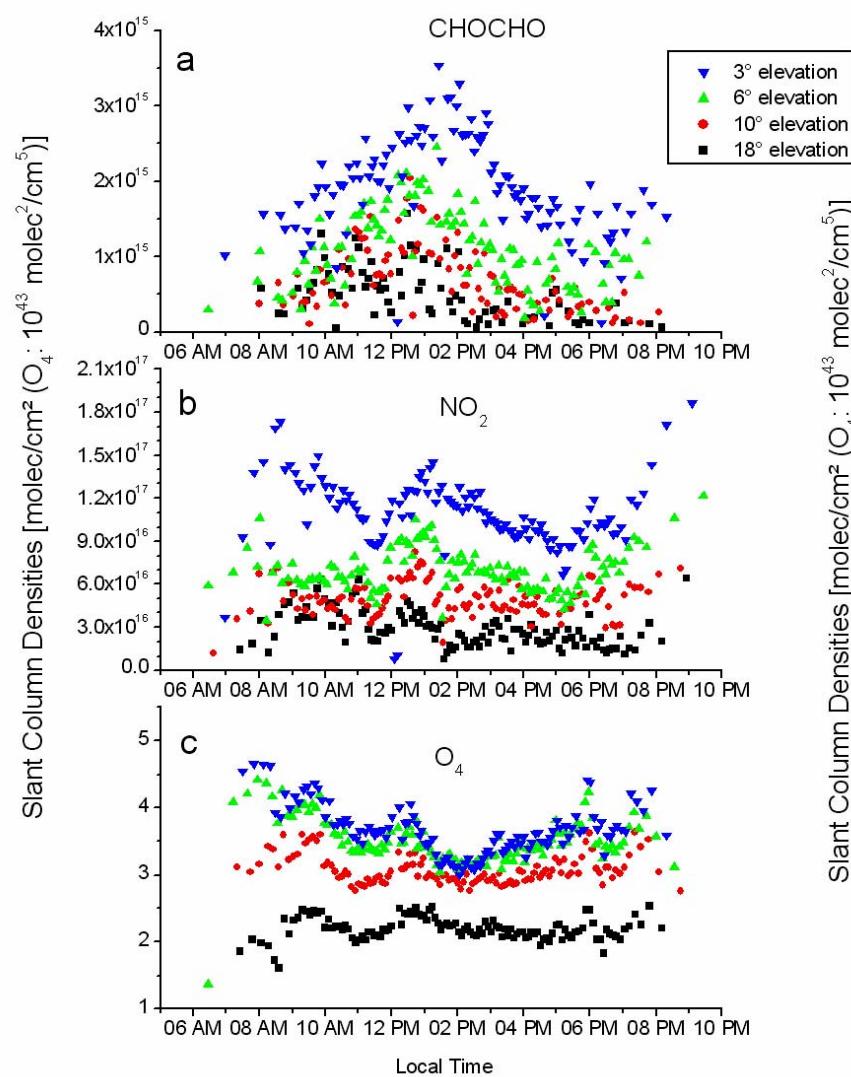
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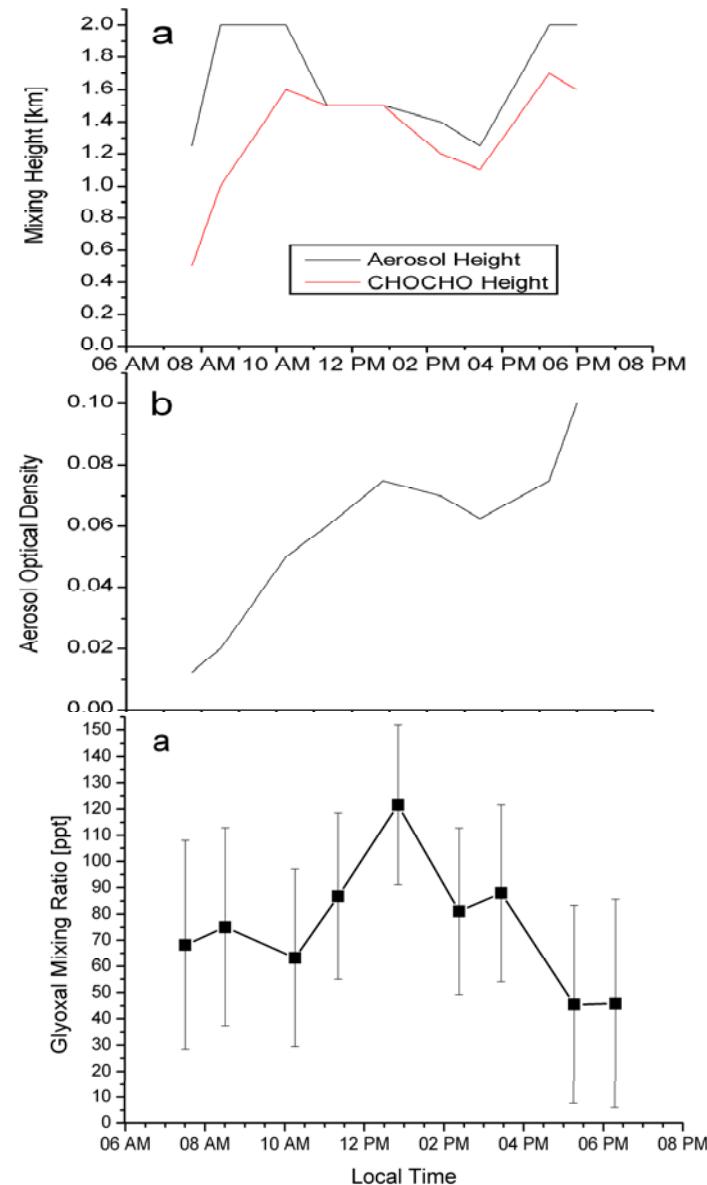


MAX-DOAS Measurements of NO₂ and Glyoxal (CHOCHO) at Cambridge (Mass.), July 26, 2004

Observed Slant Column Densities



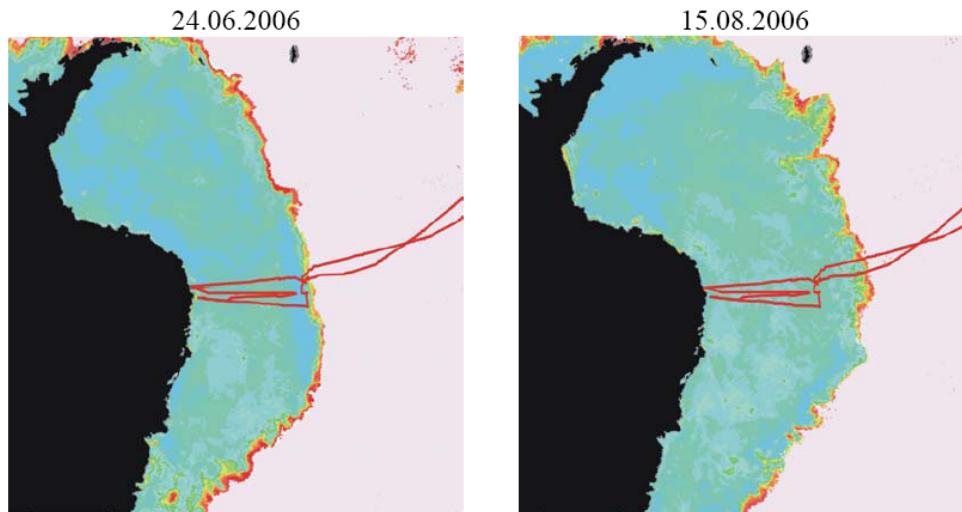
Derived Mixed-Layer Height and CHOCHO VMR



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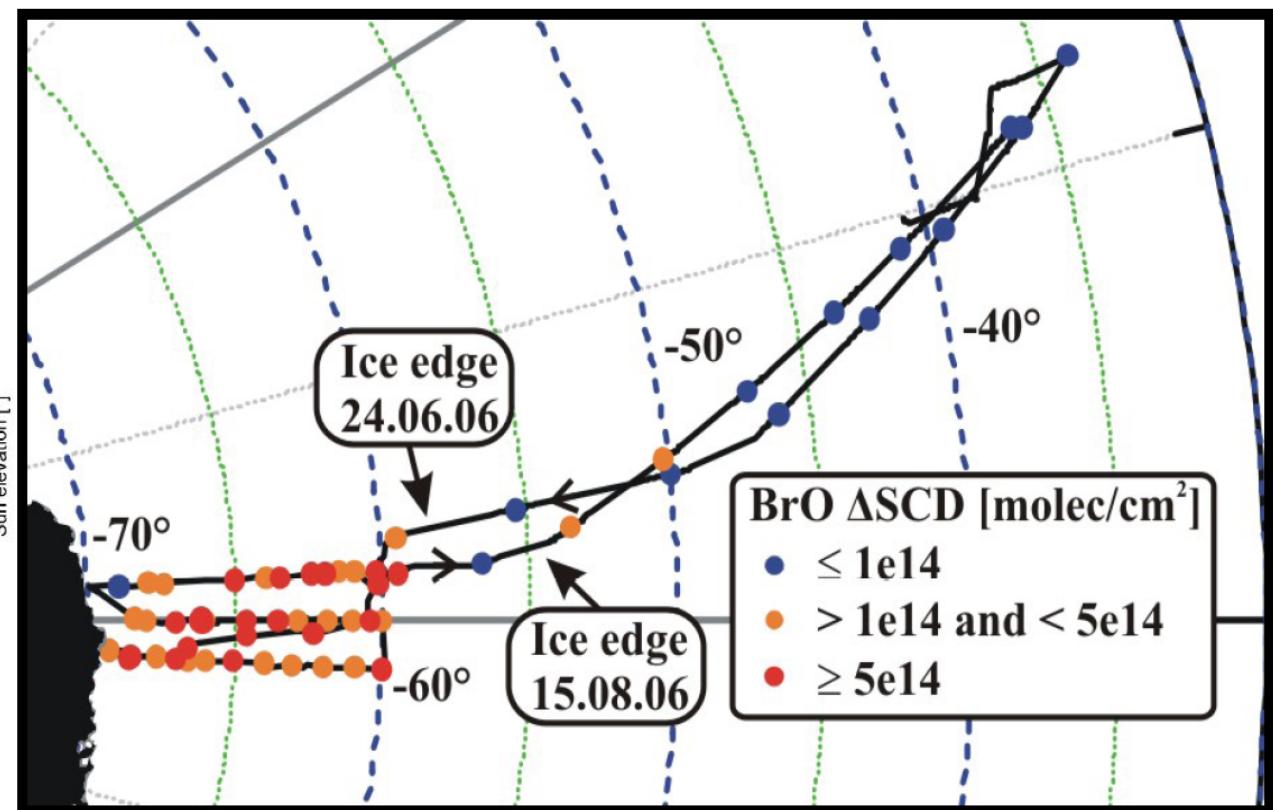
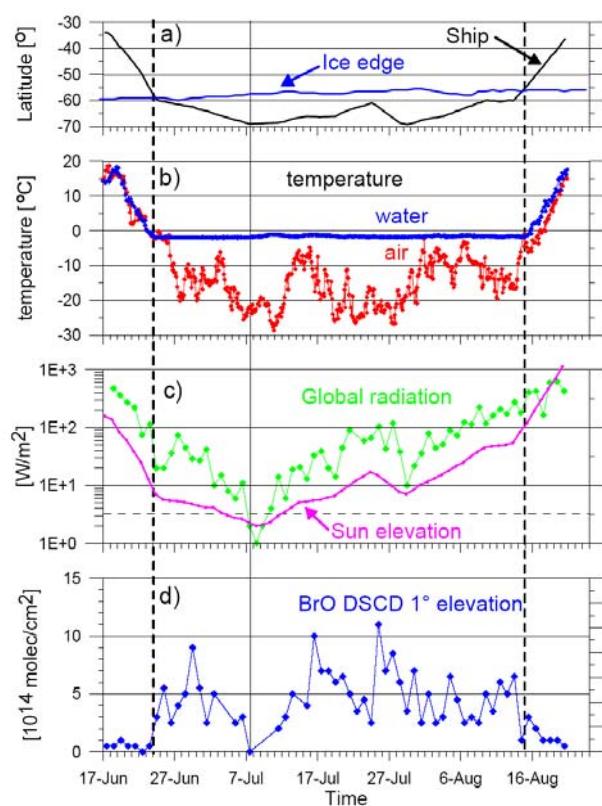
Sinreich et al., ACPD 2006

Antarctic BrO, MAX-DOAS Observations fom RV Polarstern, 2006



T. Wagner, O. Ibrahim, R. Sinreich, U. Frieß, R. von Glasow, U. Platt, ACP, 2007

BrO slant column densities
at 1° elevation



Summary

- UV-Vis Spectroscopy is a basically simple yet very powerful technique for probing the atmosphere
- For many trace gases DOAS allows sub – ppt detection limits
- DOAS is inherently calibrated
- During recent decades many variants of the DOAS technique were developed to solve a large variety of measurement problems ranging from fundamental research to pollution monitoring and satellite applications

