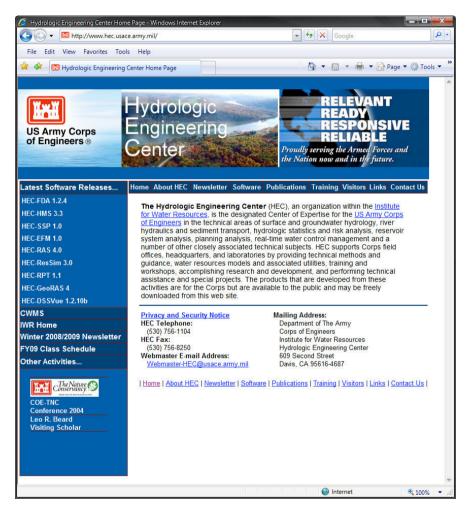
Where to get the Software and Documentation



http://www.hec.usace.army.mil/



HMS: Hydrologic Modeling System

SSP: Statistical Software Package

EFM: Ecological Functions Model

RAS: River Analysis System

RES-SIM: Reservoir Systems Simulation

DSS-Vue: Data Storage System Viewer

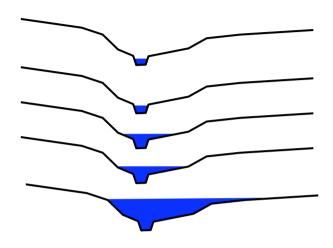


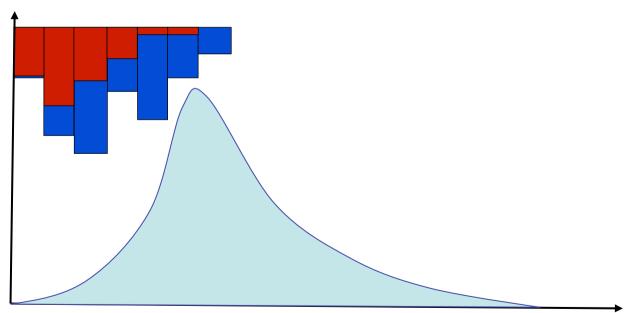


Flood, Fargo Moorhead (ND, 2001)



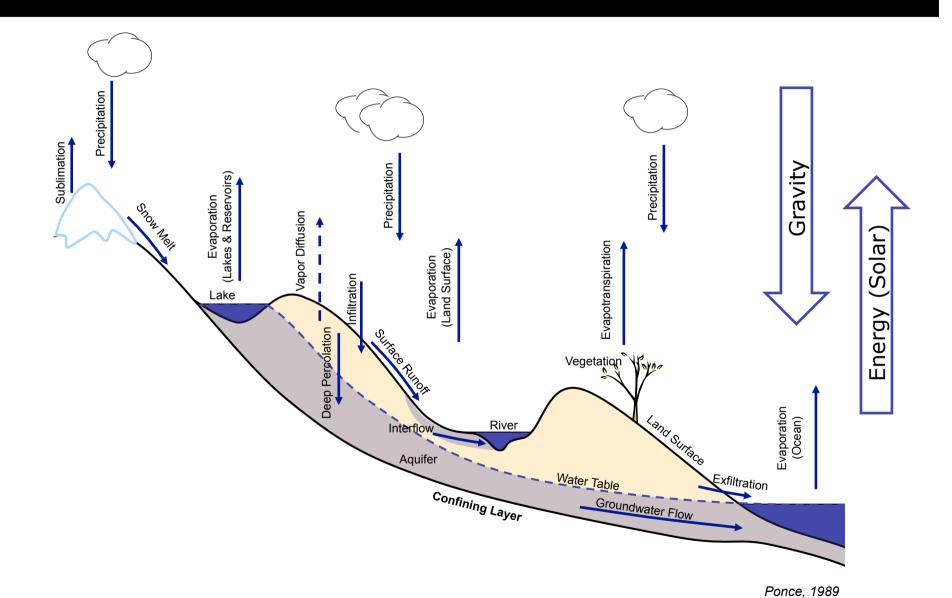










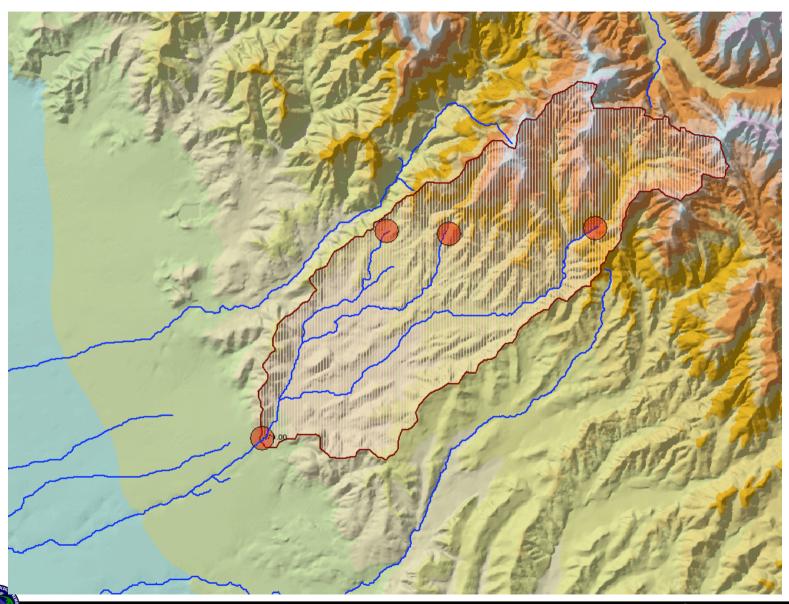






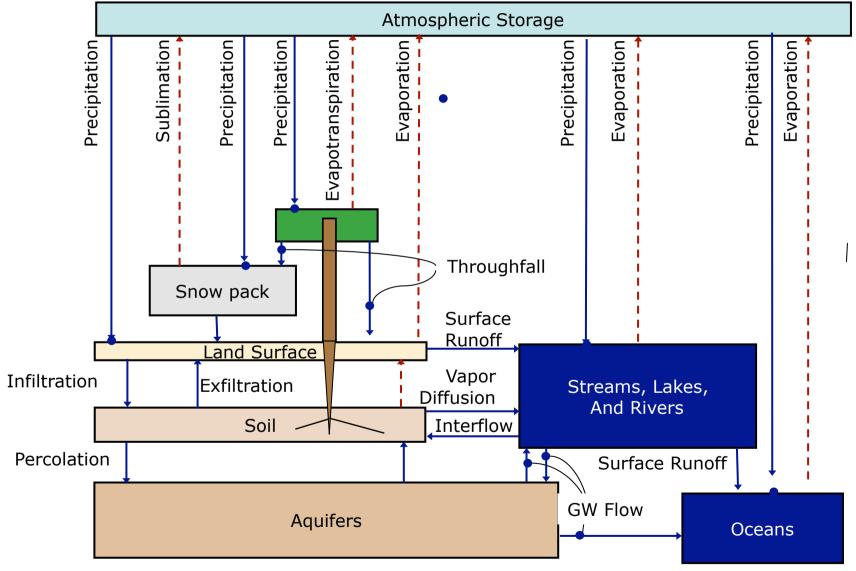
Trace The Water Drop





The Hydrologic Cycle: Compartments



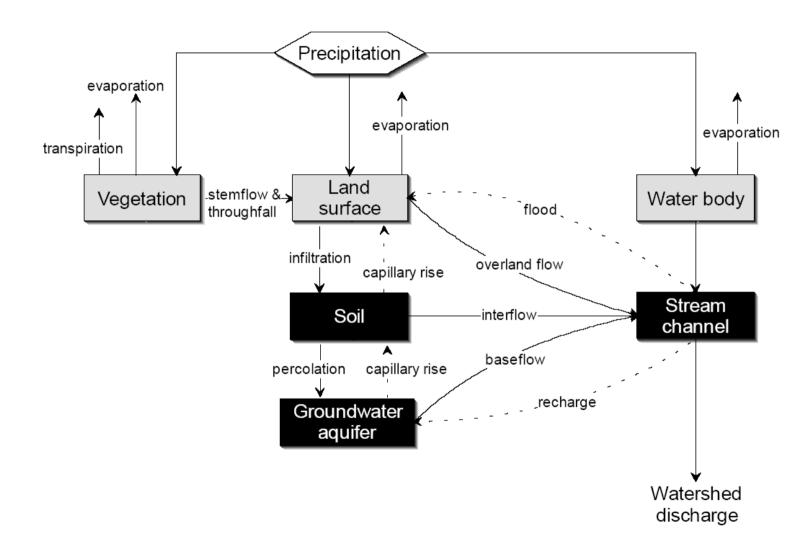






System representation



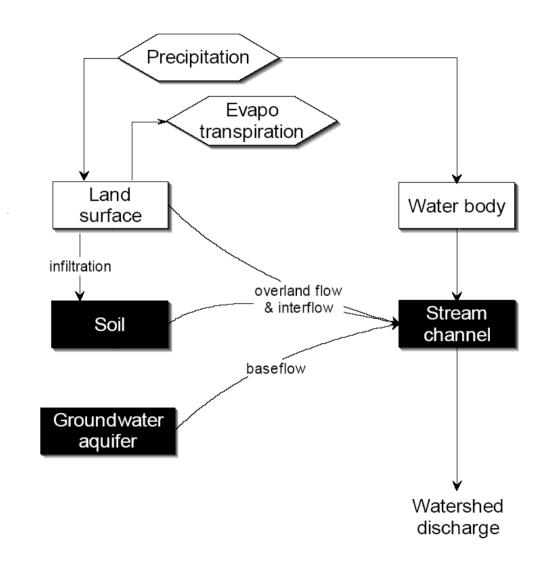






HEC's System representation

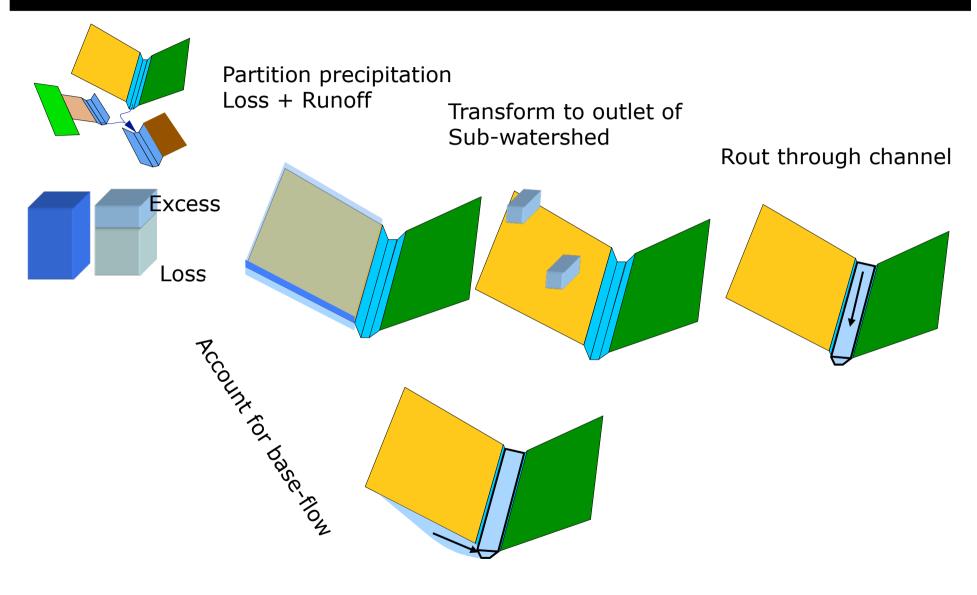










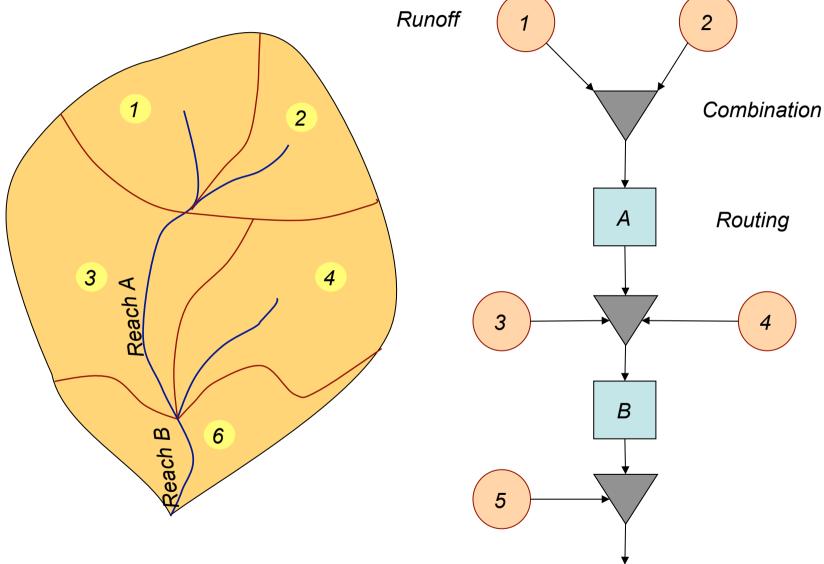






Watershed Elements



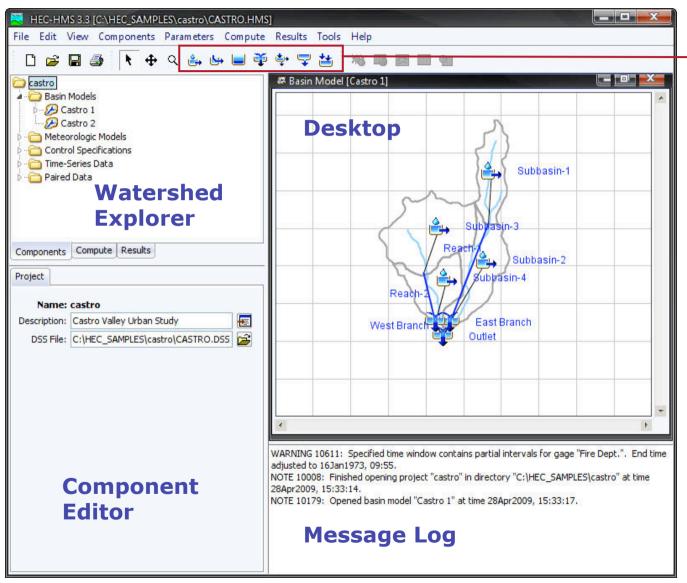






HEC-HMS Overview





Basin Model Components





Project Elements in the Watershed Explorer

castro

Basin Models

Castro 1

GageWts

🔐 Jan73

Paired Data

Castro 2

Meteorologic Models

Control Specifications

Precipitation Gages
Discharge Gages

Storage-Discharge Functions

Time-Series Data



Merteorologic Model

- Defines what type of model input to be used
- Connect model input (Time-series) to basin model
 - Methods of averaging input
 - Which gauge affects which sub-basin,

Control Specification

• Defines simulation time frame

Paired Data

- Provide data needed by reservoirs channels, and control structures
 - Stage Discharge
 - Storage discharge

Basin Models

- Provide physical watershed description
 - Hydrologic elements
 - Sub-basins
 - Reaches
 - Junctions (connectivity)
 - Reservoirs, sinks, sources and diversions
 - Define Selected calculation methods for each element
 - Define parameters for calculation methods

Time-Series Data

- Provide time dependent variables
 - Model input
 - precipitation,
 - temperature
 - solar radiation
 - Evapotranspiration
 - Source and sinks
 - Observation for calibration







Basin Model Components (Sub-basin)





Sub-basin:

- Represents the physical watershed
- Basic element in rainfall-runoff modeling and loss calculations
- Runoff is routed internally to the outlet of the sub-basin
- Baseflow (GW contribution) is added



Transform

Bounded recession

Constant monthly

Linear reservoir

Nonlinear Boussinesq

Recession

Loss

Deficit and constant rate (DC)

Exponential

Green and Ampt

Gridded DC

Gridded SCS CN

Gridded SMA

Initial and constant rate

SCS curve number (CN)

Smith Parlange

Soil moisture accounting (SMA)

Transform

Clark's UH

Kinematic wave

ModClark

SCS UH

Snyder's UH

User-specified s-graph

User-specified unit hydrograph (UH)





Basin Model Components (Reach)





Reach

- Represents channels and pipes
- Conveys streamflow downstream in the basin model
- inflow can come from any hydrologic element
- Outflow is computed by accounting for translation + attenuation of inflow hydrograph

Routing

Kinematic wave

Lag

Modified Puls

Muskingum

Muskingum-Cunge





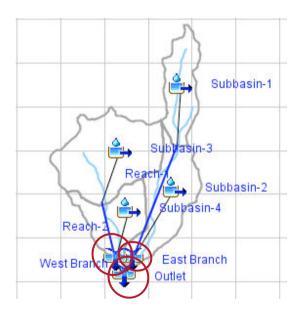


Basin Model Components (Junction)



Junction

- Represents confluences and other flow combination
- Connects upstream elements to downstream elements
- inflow can come from any hydrologic element
- Outflow is computed by summing up all inflows







Basin Model Components (Others)



Reservoir

- Model detention and attenuation
- Receives inflow from any/or many hydrologic elements upstream
- Outflow is computed from
 - Storage | outflow relationship
 - Elevation | storage | outflow
 - Elevation | area | outflow
- User can define outlet structure

Source

- Introduce external flow
- Has no inflow, but outflow is pre-defined

Diversion

- Models flow leaving main channel
- Receives inflow from any/or many hydrologic elements upstream
- Outflow consists of diverted and non-diverted flow, with diverted flow specified by user
- Both flows can be connected D/S

Sink

- Introduce external flow
- Has no inflow, but outflow is pre-defined





Sub-basin: Loss | Rainfall Runoff Relationship





Sub-basin:

- Represents the physical watershed
- Basic element in rainfall-runoff modeling and loss calculations
- Runoff is routed internally to the outlet of the sub-basin
- Baseflow (GW contribution) is added

Loss

Deficit and constant rate (DC)

Exponential

Green and Ampt

Gridded DC

Gridded SCS CN

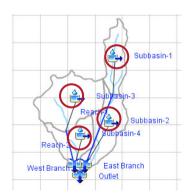
Gridded SMA

Initial and constant rate

SCS curve number (CN)

Smith Parlange

Soil moisture accounting (SMA)

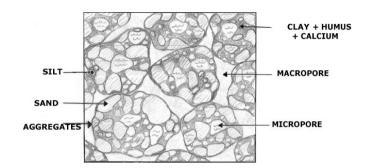






Soil Profile (Detailed)





Porosity: Voids in the soil profile

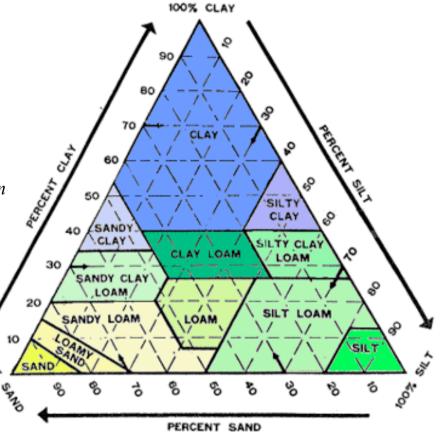
Field Capacity: Amount of water held against gravity

Wilting Point: amount of water below which plants can

not extract water from the soil

Hydraulic conductivity: Rate of water flow in soil

Values depend, among other factors, on soil texture



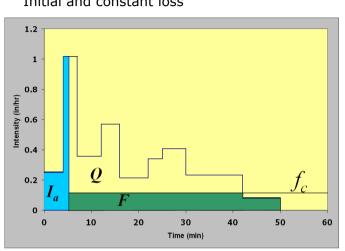


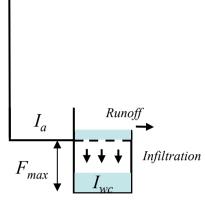


Constant Infiltration Rate Models

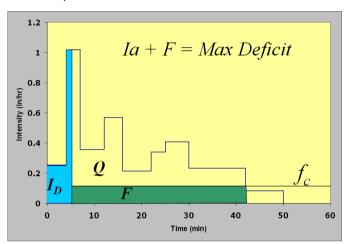


Initial and constant loss





Initial, constant and maximum loss



Soil group	Description	Range of loss rates (in/hr)
A	Deep sand, deep loess, aggregated silts	0.30-0.45
В	Shallow loess, sandy loam	0.15-0.30
С	Clay loams, shallow sandy loam, soils low in organic content, and soils usually high in clay	0.05-0.15
D	Soils that swell significantly when wet, heavy plastic clays, and certain saline soils	0.00-0.05

$$pe_t = \begin{cases} 0 & \text{if } \sum p_i < I_a \\ p_t - f_c & \text{if } \sum p_i > I_a \text{ and } p_t > f_c \\ 0 & \text{if } \sum p_i > I_a \text{ and } p_t < f_c \end{cases}$$

When Maximum Soil Storage is considered

$$p_{et} = p_t if F_t \ge F_{max}$$

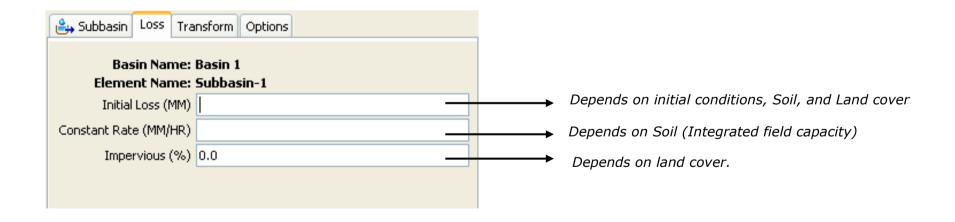




Parameters for Deficit Constant Model



Subbasin Loss Transform Options	
Element Name: Subbasin-1	Depends on initial conditions $(q_s - q_i)$
Initial Deficit (MM)	
Maximum Deficit (MM)	Depends on Soil (Integrated field capacity)
Constant Rate (MM/HR)	Depends on soil. (# Sat. Conductivity)
Impervious (%) 0.0	Depends on land cover.







Initial Value of Loss Rate Parameter



Soil group	Description	Range of loss rates (in/hr)
A	Deep sand, deep loess, aggregated silts	0.30-0.45
В	Shallow loess, sandy loam	0.15-0.30
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D	Soils that swell significantly when wet, heavy plastic clays, and certain saline soils	0.00-0.05

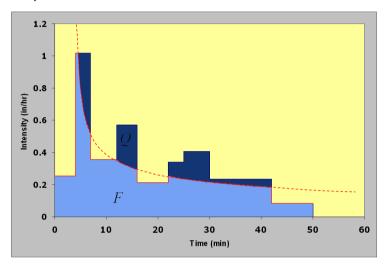


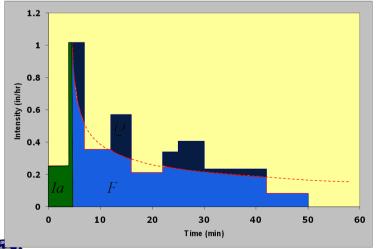


Infiltration Rate Decay Models



This type of models consider a more "realistic" infiltration approximation. Infiltration starts with the beginning of the storm at a very high rate, and it decays exponentially and almost asymptotically to a constant rate by the end of the storm. At any point, the infiltration rate depends on the amount of water that has already infiltrated.





$$pe_{t} = \begin{cases} 0 & \text{if } p_{t} \leq f_{t} \\ p - f_{t} & \text{if } p_{t} > f_{t} \end{cases}$$

$$f_{t} \propto \frac{1}{F_{t}} \Leftrightarrow \frac{1}{\sum_{t=0}^{t} f_{t}}$$

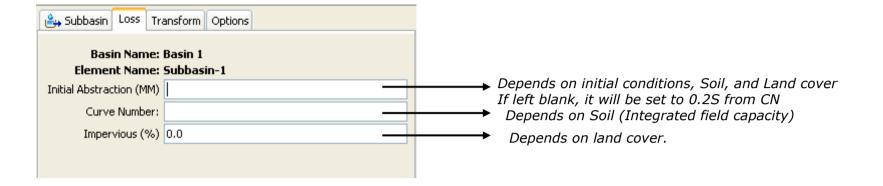
$$S = \begin{cases} \frac{1000 - 10 \ CN}{CN} & \text{(foot - pound system)} \\ \frac{25400 - 254 \ CN}{CN} & \text{(SI)} \end{cases}$$

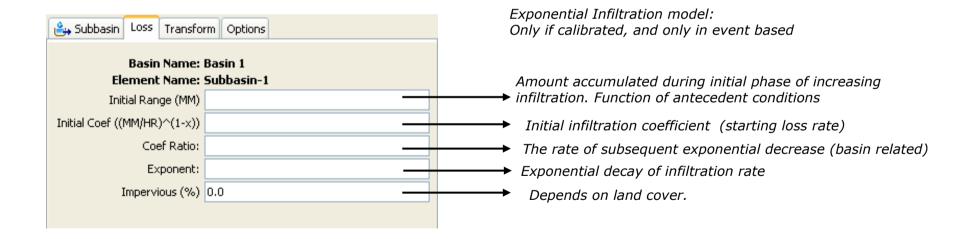
$$I_a = 0.2 S$$



CN method Parameters in HEC-HMS











Example Land Use Table for CN



SCS TR-55 Table 2-2a - Runoff curve numbers for urban areas¹

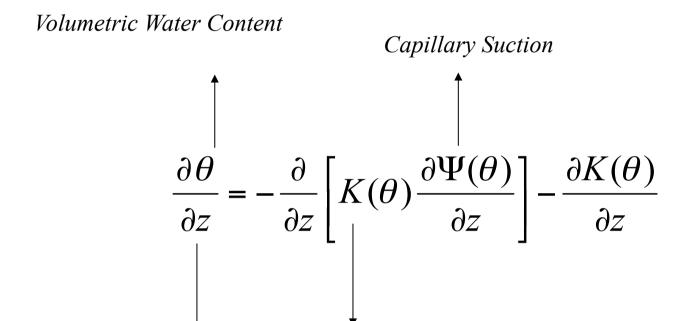
Cover description			Curve numbers for hydrologic soil group			
Cover type and hydrologic condition	Average percent impervious area ²	A	В	С	D	
Fully developed urban areas						
Open space (lawns, parks, golf courses, cemeteries,						
etc.) ³ :						
Poor condition (grass cover < 50%)		68	79	86	89	
Fair condition (grass cover 50% to 75%)		49	69	79	84	
Good condition (grass cover > 75%)		39	61	74	80	
Impervious areas:						
Paved parking lots, roofs, driveways, etc.						
(excluding right-of-way)		98	98	98	98	
Streets and roads:						
Paved; curbs and storm sewers (excluding		00	00	00		
right-of-way)		98	98	98	98	
Paved; open ditches (including right-of-way)		83	89	92	93	
Gravel (including right-of-way)		76	85	89	9:	
Dirt (including right-of-way)		72	82	87	89	
Western desert urban areas:		-		0.5		
Natural desert landscaping (pervious areas only) ⁴		63	77	85	88	
Artificial desert landscaping (impervious weed						
barrier, desert shrub with 1- to 2-inch sand		96	96	96	96	
or gravel mulch and basin borders)		70	,,,	,,,		
Commercial and business	85	89	02	94	95	
		81	92 88	91	93	
Industrial	72	81	88	91	93	
1/8 acre or less (town houses)	65	77	85	90	92	
1/4 acre	38	61	75	83	87	
1/3 acre	30	57	72	81	86	
1/2 acre	25	54	70	80	85	
1 acre	20	51	68	79	84	
2 acre	12	46	65	77	82	
Developing urban areas	12	10	05	,,	02	
Newly graded areas (pervious areas only,						
no vegetation) ⁵		77	86	91	94	





Semi-physically-based Infiltration





Hydraulic Conductivity

Distance from the Surface





Green Ampt Model



$$f_t = K \left[\frac{1 + (\phi - \theta_i) S_f}{F_t} \right]$$

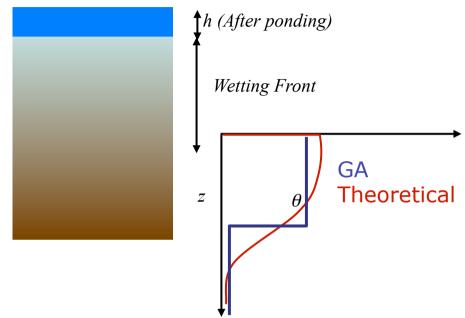
 $f_t = loss during period t;$

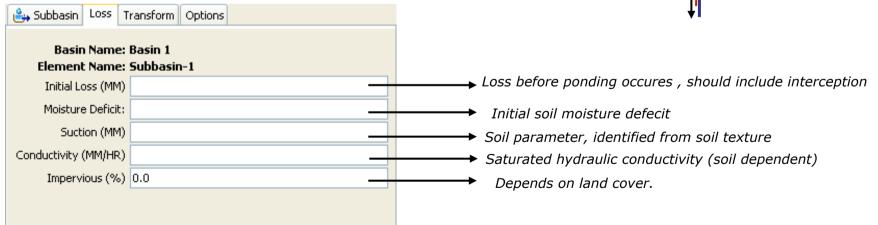
K = saturated hydraulic conductivity

 $(\phi - \theta_i) = volumetric moisture deficit;$

 S_f = wetting front suction;

 F_t = cumulative loss at time t.





Soil Classification and model parameters



Texture class	Porosity, ϕ (cm ³ /cm)	Hydraulic conductivity, θ _i , saturated (cm/hr)	Wetting front suction (cm)
Sand	0.437	21.00	10.6
Loamy sand	0.437	6.11	14.2
Sandy loam	0.453	2.59	22.2
Loam	0.463	1.32	31.5
Silt loam	0.501	0.68	40.4
Sandy clay loam	0.398	0.43	44.9
Clay loam	0.464	0.23	44.6
Silty clay loam	0.471	0.15	58.1
Sandy clay	0.430	0.12	63.6
Silty clay	0.479	0.09	64.7
Clay	0.475	0.06	71.4

Porosity = Volume of Voids/Volume of Solids





Or Using Soil Texture

TABLE 5.3.2 Water-Retention Properties Classified by Soil Texture

		Total porosity	Residual water content	Effective porosity	Bubbling pressure h_b	Pore-size distribution λ	Water retained at	Water retained at
Texture class	Sample size	ϕ , cm ³ /cm ³	θ_r , cm ³ /cm ³	ϕ_e , cm ³ /cm ³	Geometric,† mean, cm	Arithmetic mean	-33 kPa, cm ³ /cm ³	- 1500 kPa, cm ³ /cm ³
Sand	762	0.437* (0.374-0.500)	0.020 (0.001-0.039)	0.417 (0.354-0.480)	7.26 (1.36–38.74)	0.694 (0.298 – 1.090)	0.091 (0.018-0.164)	0.033 (0.007-0.059)
Loamy sand	338	0.437 (0.368-0.506)	0.035 (0.003-0.067)	0.401 (0.329-0.473)	8.69 (1.80-41.85)	0.553 (0.234-0.872)	0.125 (0.060-0.190)	0.055 (0.019-0.091)
Sandy loam	666	0.453 (0.351-0.555)	0.041 (-0.024-0.106)	0.412 (0.283-0.541)	14.66 (3.45–62.24)	0.378 (0.140-0.616)	0.207 (0.126-0.288)	0.095 (0.031-0.159)
Loam	383	0.463 (0.375-0.551)	0.027 (-0.020-0.074)	0.434 (0.334-0.534)	11.15 (1.63–76.40)	0.252 (0.086-0.418)	0.270 (0.195-0.345)	0.117 (0.069-0.165)
Silt loam	1206	0.501 (0.420-0.582)	0.015 (-0.028-0.058)	0.486 (0.394-0.578)	20.76 (3.58–120.4)	0.234 (0.105 – 0.363)	0.330 (0.258-0.402)	0.133 (0.078-0.188)
Sandy clay loam	498	0.398 (0.332-0.464)	0.068 (-0.001-0.137)	0.330 (0.235 – 0.425)	28.08 (5.57 – 141.5)	0.319 (0.079 – 0.559)	0.255 (0.186-0.324)	0.148 (0.085-0.211)
Clay loam	366	0.464 (0.409-0.519)	0.075 (-0.024-0.174)	0.390 (0.279-0.501)	25.89 (5.80-115.7)	0.242 (0.070-0.414)	0.318 (0.250-0.386)	0.197 (0.115-0.279)
Silty clay loam	689	0.471 (0.418-0.524)	0.040 (-0.038-0.118)	0.432 (0.347 – 0.517)	32.56 (6.68 – 158.7)	0.177 (0.039-0.315)	0.366 (0.304-0.428)	0.208 (0.138-0.278)
Sandy clay	45	0.430 (0.370-0.490)	0.109 (0.013-0.205)	0.321 (0.207 – 0.435)	29.17 (4.96–171.6)	0.223 (0.048-0.398)	0.339 (0.245-0.433)	0.239 (0.162-0.316)
Silty clay	127	0.479 (0.425-0.533)	0.056 (-0.024-0.136)	0.423 (0.334-0.512)	34.19 (7.04–166.2)	0.150 (0.040-0.260)	0.387 (0.332-0.442)	0.250
Clay	291	0.475 (0.427–0.523)	0.090 (-0.015-0.195)	0.385 (0.269-0.501)	37.30 (7.43 – 187.2)	0.165 (0.037-0.293)	0.396 (0.326-0.466)	0.272 (0.208-0.336)

^{*} First line is the mean value. Second line is ± one standard deviation about the mean.

† Antilog of the log mean.

Source: Reproduced from Ref. 80 by permission of ASCE.





Smith-Parlange Model



$$f_t = K[\frac{C_o}{KF_t} + 1]$$

$$f_{t} = K \frac{e^{F_{t}/C_{0}}}{e^{F_{t}/C_{0}} - 1}$$

$$C_o = [2S_f K(\phi - \theta_i)]$$

Element Name:	Subbasin-1	
Initial Content:		
Residual Content:		
Saturated Content:		
Bubbling Pressure (MM)		
Pore Distribution:		
Conductivity (MM/HR)		
Impervious (%)	0.0	
Temperature Gage:	None	F

 $f_t = loss during period t;$

K = saturated hydraulic conductivity

 $(\phi - \theta_i) = volume\ moisture\ deficit;$

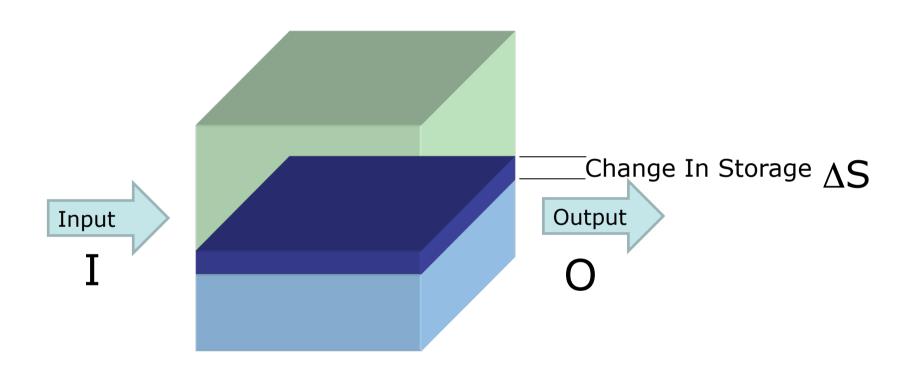
 S_f = wetting front suction;

 F_t = cumulative loss at time t.

 C_o = Sorptivity can also be estimated using other soil parameters

Conceptual Models





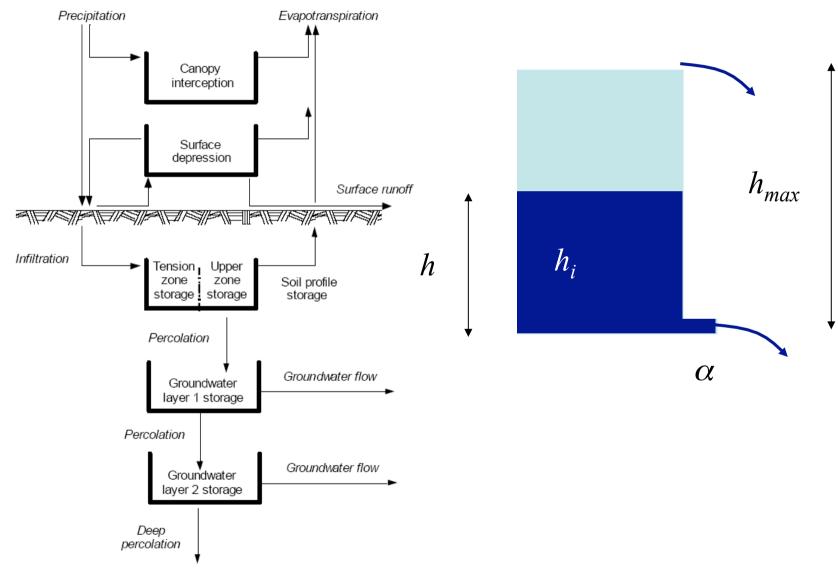
$$I - O = \Delta S$$





Conceptual Model (SMA)



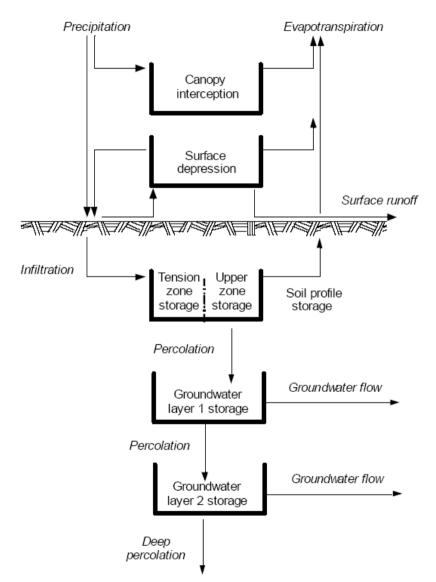


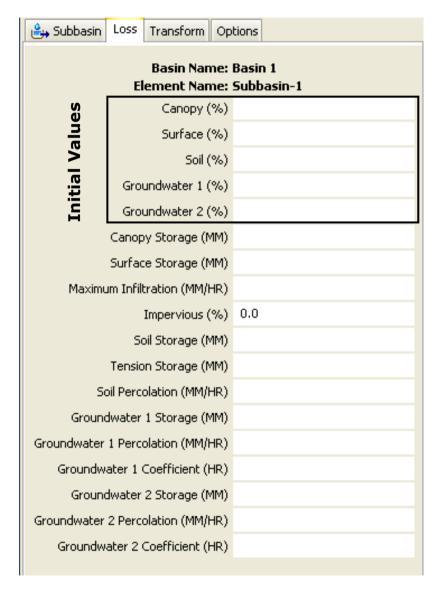




Conceptual Model (SMA)





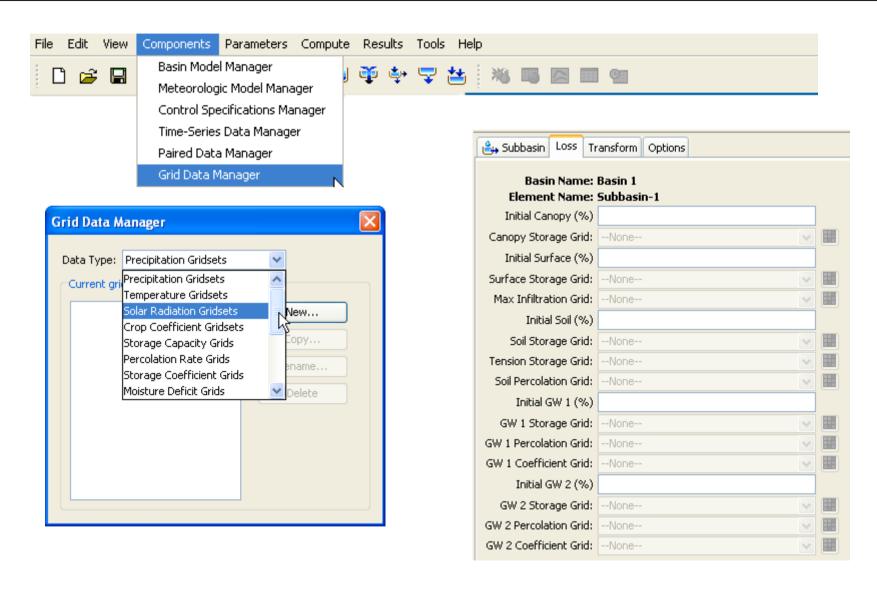






Gridded Models









Basin Model Components (Sub-basin)





Sub-basin:

- Represents the physical watershed
- Basic element in rainfall-runoff modeling and loss calculations
- Runoff is routed internally to the outlet of the sub-basin
- Baseflow (GW contribution) is added



Transform

Clark's UH

Kinematic wave

ModClark

SCS UH

Snyder's UH

User-specified s-graph

User-specified unit hydrograph (UH)



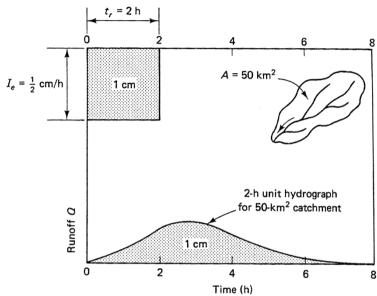


Unit Hydrograph Concept



Basic Definition

Consider a unit depth of excess generated uniformly (spatially and temporally) over a watershed with area A, during time duration $t_{\rm r}$. The resulting hydrograph is called the $t_{\rm r}$ Unit Hydrograph for the watershed. The intensity of the excess rain is given as $1/t_{\rm r}$



Main Assumptions

- The P_e is distributed uniformly spatially I_e is constant during each Dt.
- The direct runoff hydrograph resulting from a given increment of excess is independent
 of the time of occurrence of the excess and of the antecedent precipitation. This is the
 assumption of time-invariance.
- Precipitation excesses of equal duration are assumed to produce hydrographs with equivalent time bases regardless of the intensity of the precipitation.

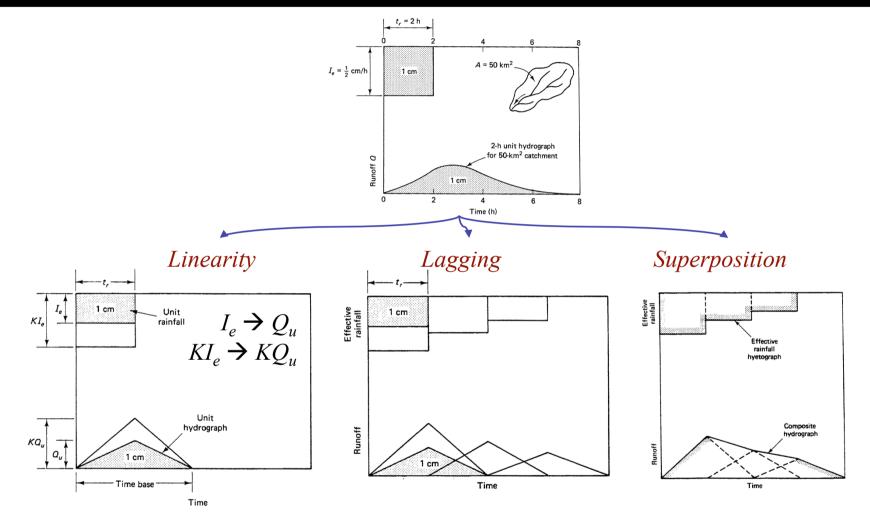
If these assumptions are valid, then we can assume a linear system and the following properties are valid





Unit Hydrograph: Its application





Allows us to identify hydrograph resulting from more/or less than one unit of excess rain

Allows us to identify hydrographs for each rain interval in a sequence of excess rains Allows us to combine hydrograph resulting from non-unit rain intervals in a sequence of excess rain. That is the event hydrograph





Discrete Convolution: The combination of all three properties



$$Q_n = \sum_{m=1}^{n \le M} P_m U_{n-m+1}$$

where Q_n = storm hydrograph ordinate at time nDt;

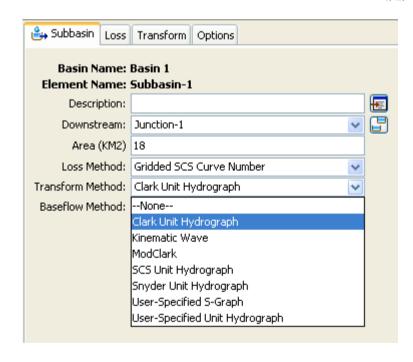
 P_m = rainfall excess depth in time interval mDt to $(m+1)\Delta t$;

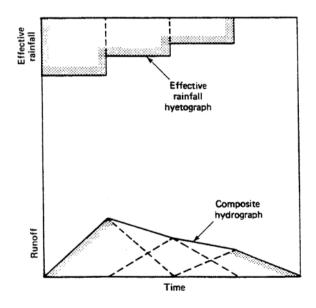
M = total number of discrete rainfall pulses; and

 $U_{n-m+1} = UH \text{ ordinate at time } (n-m+1)\Delta t$

 Q_n and P_m are expressed as flow rate and depth respectively,

 U_{n-m+1} (flow rate per unit depth).



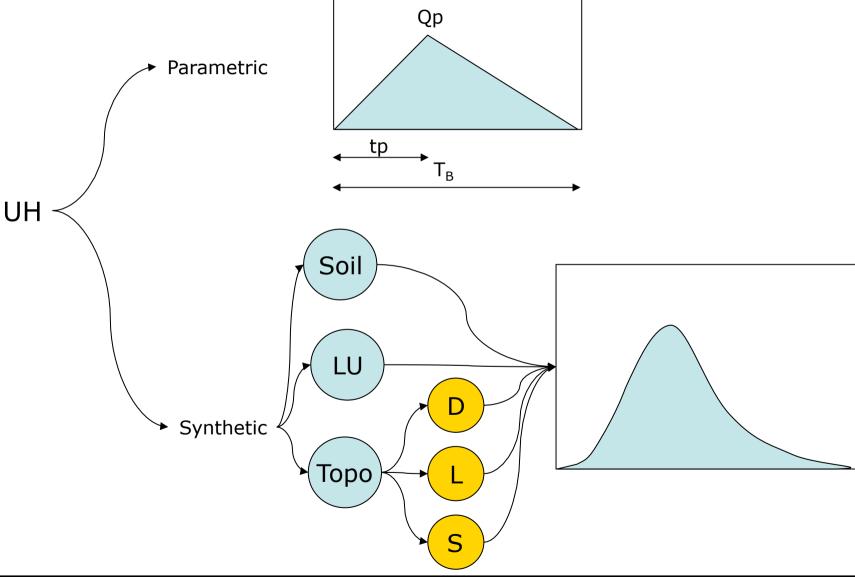






Types of Unit Hydrographs









Snyder UH: 1938 (Parameters)



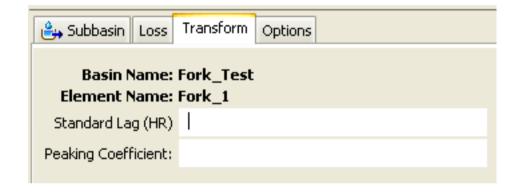
Snyder's observations \rightarrow

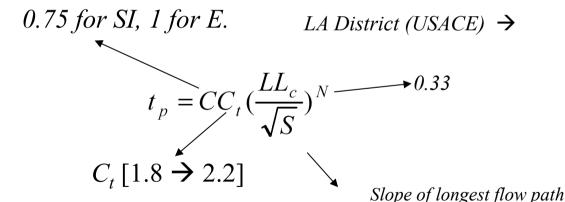
$$t_p = C_t (LL_c)^{0.3}$$

 $km: C_t[1.35 \to 1.65]$

$$t_p = C_t (LL_c)^{0.3}$$

 $mile: C_t [1.8 \rightarrow 2.2]$





Ct = basin coefficient;

L = length of the main stream from the outlet to the divide;

Lc = length along the main stream from the outlet to a point nearest the watershed centroid

SCS UH: Dimensionless UH (Concept)



$$U_{P} = C \frac{A}{T_{P}} - Watershed Area$$

$$Duration of Pe$$

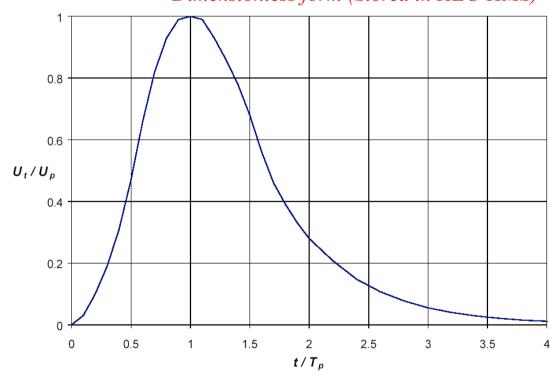
$$T_{p} = \frac{\Delta t}{2} + t_{lag} - Lag Time$$

$$Pe Centroid to Qp$$

Dimensionless form (Stored in HEC-HMS)

$$t_{lag} = 0.6 t_c$$

Time of Concentration







SCS UH: Parameters (Parameters -First Approach)



$$t_{lag} = 0.6 t_c$$

$$t_{c} = t_{sheet} + t_{shallow} + t_{channel}$$

$$t_{sheet} = \frac{0.007(NL)^{0.8}}{(P_{2})^{0.5}S^{0.4}} \qquad V = \begin{cases} 16.1345\sqrt{S} & \text{for unpaved surface} \\ 20.3282\sqrt{S} & \text{for paved surface} \end{cases}$$

$$t_{shallow} = \frac{L - 200}{V_{shallow}}$$

 P_2 : 2 yr/24 hr rainfall depth

N: Overland Roughness Coefficient

L: Longest Flow Path

S: Hydraulic slope, approximate with slope

$$t_{channel} = \frac{L}{V}$$

$$CR^{2/3}S^{1/2}$$





Overland flow roughness coefficient

Table 6-1. Overland-flow roughness coefficients for sheet-flow modeling (USACE, 1998)

Surface description	N
Smooth surfaces (concrete, asphalt, gravel, or bare soil)	0.011
Fallow (no residue)	0.05
Cultivated soils:	
Residue cover ≤ 20%	0.06
Residue cover > 20%	0.17
Grass:	
Short grass prairie	0.15
Dense grasses, including species such as weeping love grass, bluegrass, buffalo grass, blue grass, and native grass mixtures	0.24
Bermudagrass	0.41
Range	0.13
Woods ¹	
Light underbrush	0.40
Dense underbrush	0.80

Notes:





 $^{^{1}}$ When selecting N, consider cover to a height of about 0.1 ft. This is the only part of the plant cover that will obstruct sheet flow.

SCS UH: Parameters (small catchments < 8 km²)



$$t_{lag} = 0.6 t_c$$

$$t_{lag} = \frac{L^8 (2540 - 22.86CN)^{0.7}}{14104CN^{0.7}Y^{0.5}}$$

$$t_{lag} = \frac{L^{0.8} (1000 - 9CN)^{0.7}}{1900CN^{0.7}Y^{0.5}}$$

L: Longest flow path (meters)

Y: Average slope (meter/meter)

CN: Basin curve number

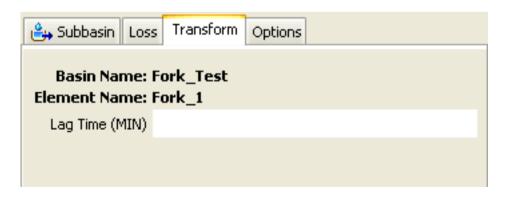
 t_{lag} : lag time in hours

L: Longest flow path (ft)

Y: Average slope (%)

CN: Basin curve number

 t_{lag} : lag time in hours







Clark's UH

R

Translation or movement of the excess from its origin throughout the drainage to the watershed outlet

Attenuation or reduction of the magnitude of the discharge as the excess is stored throughout the watershed.

$$\frac{dS}{dt} = I_t - O_t$$

$$S_t = RO_t$$

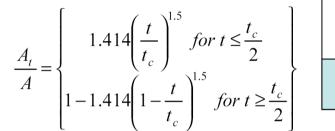
$$O_t = C_A I_t + C_B O_{t-1}$$

Routing coefficients

$$C_A = \frac{\Delta t}{R + 0.5\Delta t}$$

$$C_B = 1 - C_A$$

$$\overline{O}_t = \frac{O_{t-1} + O_t}{2}$$



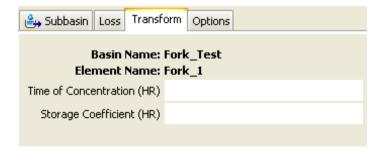
where

At = cumulative watershed area contributing at time t;

A = total watershed area

 $tc = time\ of\ concentration\ of\ watershed =$

In HEC-HMS ← *Calibration or using SCS approach*





Limitations (Data)



Availability of information for calibration or parameter estimation.

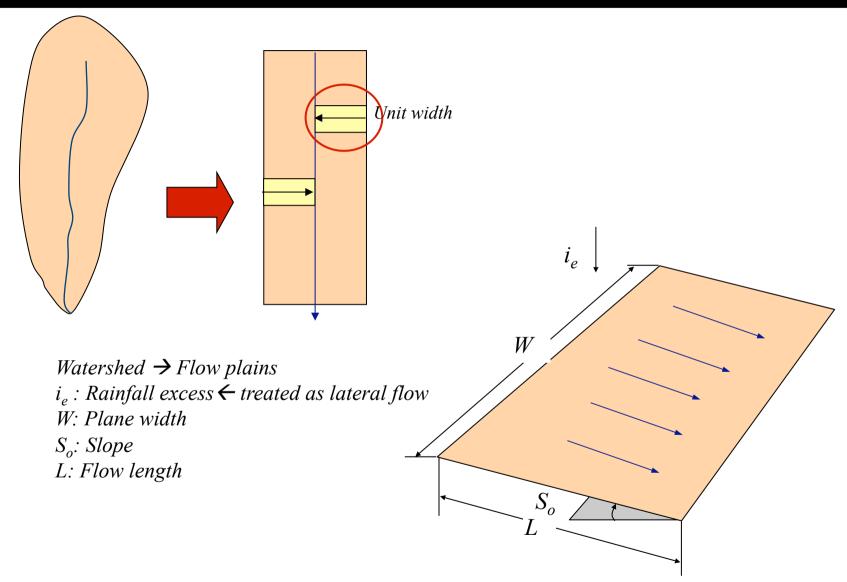
- Parametric UH models requires model parameters.
- Empirical parameter predictors
- Optimal source of these parameters is calibration,
- If calibration in an urban watershed is not available, then use the kinematic-wave model





Physically Based Models of Overland Flow Theory









Approximations



$$S_f = S_0 - \frac{\partial y}{\partial x} - \frac{V}{g} \frac{\partial V}{\partial x} - \frac{1}{g} \frac{\partial V}{\partial t}$$

$$S_f = S_0 - \frac{\partial y}{\partial x} - \frac{V}{g} \frac{\partial V}{\partial x}$$
 Quasi Steady State Dynamic Wave \longrightarrow HEC-RAS Steady flow

$$S_f = S_0 - \frac{\partial y}{\partial x} - \frac{V}{g} \frac{\partial V}{\partial x} - \frac{1}{g} \frac{\partial V}{\partial t}$$

$$St. \ Venant / Dynamic \ Wave \longrightarrow Unsteady / Gradually \ varied \ Flow \ Many \ numerical \ models$$





Kinematic Wave: Combine Simplifications



In shallow flow over a plain

$$\frac{\partial Q}{\partial x} + \frac{\partial A}{\partial t} = q$$

$$Q = \alpha A^{m}$$
Combine

$$\frac{\partial A}{\partial t} + \alpha m A^{(m-1)} \frac{\partial A}{\partial x} = q$$

In HEC-HMS (for shallow overland flow) on a unit width of a wide rectangular channel

$$\alpha = 1.486S^{1/2} / N$$

and

$$m = \frac{5}{3}$$

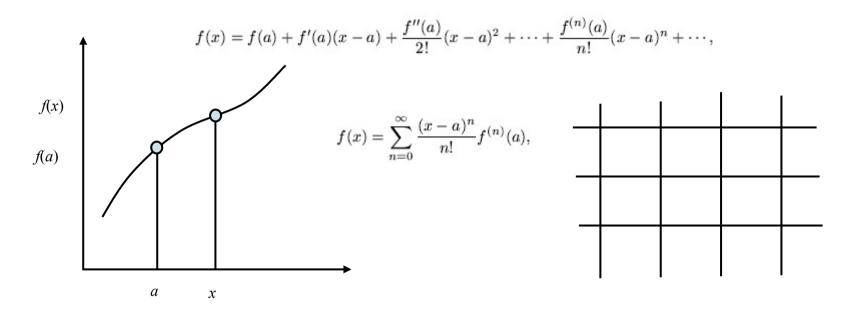




Basic Concept of Finite Difference Method



Taylor Series



Needs boundary conditions

Needs method to approximate $f^{(n)} \rightarrow How$ *many neighboring points will you consider*

Central Implicit

Explicit



Overland flow roughness coefficient

Table 6-1. Overland-flow roughness coefficients for sheet-flow modeling (USACE, 1998)

Surface description	N
Smooth surfaces (concrete, asphalt, gravel, or bare soil)	0.011
Fallow (no residue)	0.05
Cultivated soils:	
Residue cover ≤ 20%	0.06
Residue cover > 20%	0.17
Grass:	
Short grass prairie	0.15
Dense grasses, including species such as weeping love grass, bluegrass, buffalo grass, blue grass, and native grass mixtures	0.24
Bermudagrass	0.41
Range	0.13
Woods ¹	
Light underbrush	0.40
Dense underbrush	0.80

Notes:





 $^{^{1}}$ When selecting N, consider cover to a height of about 0.1 ft. This is the only part of the plant cover that will obstruct sheet flow.

HEC-HMS Implementation for Sub-basins



 $\Delta x/\Delta t \approx c \rightarrow Accurate$ and stable solution

 $c = average \ kinematic$ -wave speed over a distance increment Δx

Types of routing elements for KW (see pp 71 for details)

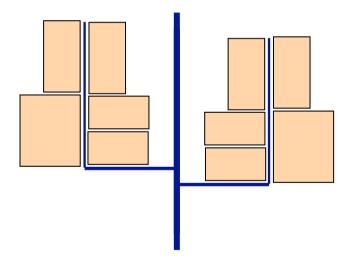


Table 6-3. Information needs for kinematic wave modeling

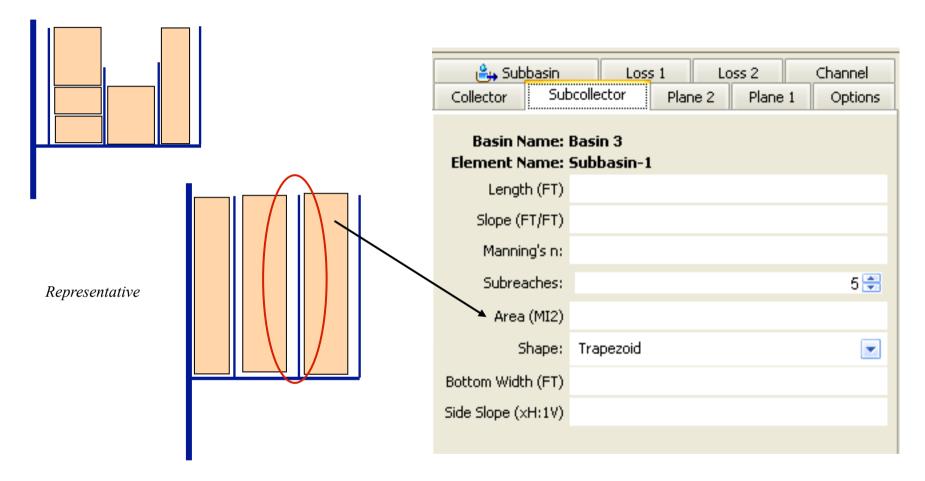
Overland flow planes	Collectors and subcollectors	Main channel
Typical length	Area drained by channel	Channel length
Representative slope Overland-flow roughness	Representative channel length	Description of channel shape
coefficient Area represented by plane	Description of channel shape	Principle dimensions of channel cross section
Loss model parameters (see Chapter 5)	Principle dimensions of representative channel cross section	Channel slope Representative Manning's roughness coefficient
	Representative channel slope	Identification of upstream inflow hydrograph (if any)
	Representative Manning's roughness coefficient	





Sub-Collectors





Small feeder pipes or channels,

Principle dimension < 18 inches,

They might service area < 10 acres.

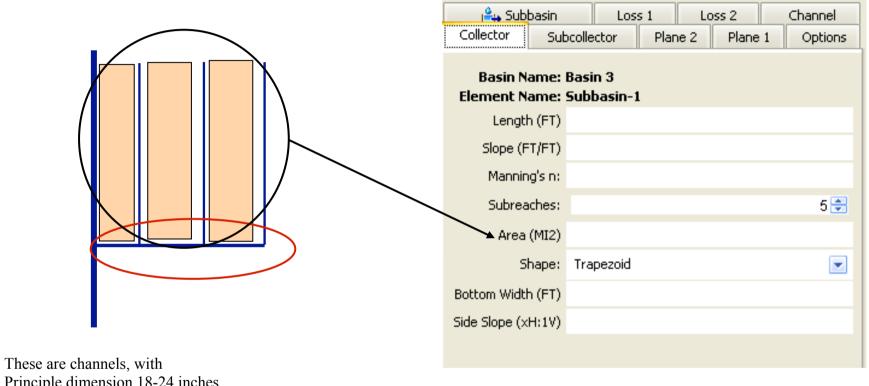
Flow is assumed to enter the channel uniformly along its length.





Collectors





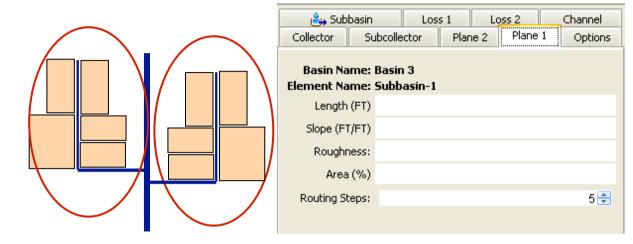
Principle dimension 18-24 inches, Collect flows from sub-collectors Convey flow to main channel. Flow enters latterally

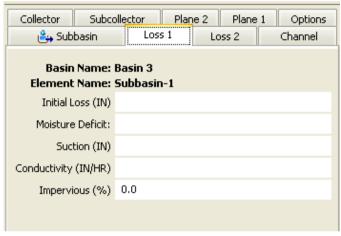


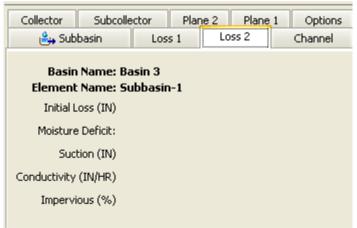


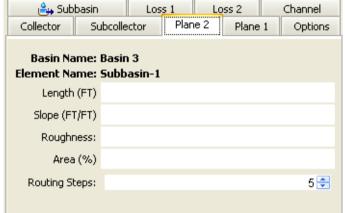
Planes









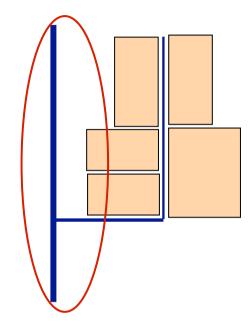




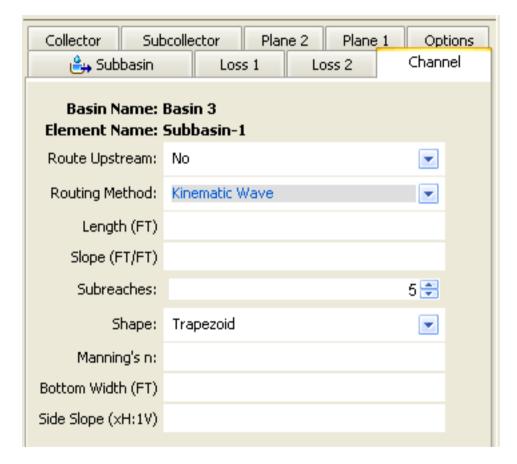


Channel





Conveys flow from upstream subwatersheds Convey flows that enter from the collector channels or overland flow planes







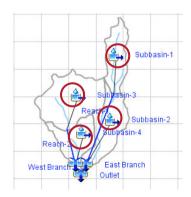
Basin Model Components (Sub-basin)





Sub-basin:

- Represents the physical watershed
- Basic element in rainfall-runoff modeling and loss calculations
- Runoff is routed internally to the outlet of the sub-basin
- Baseflow (GW contribution) is added



Transform

Bounded recession

Constant monthly

Linear reservoir

Nonlinear Boussinesq

Recession





Constant Monthly



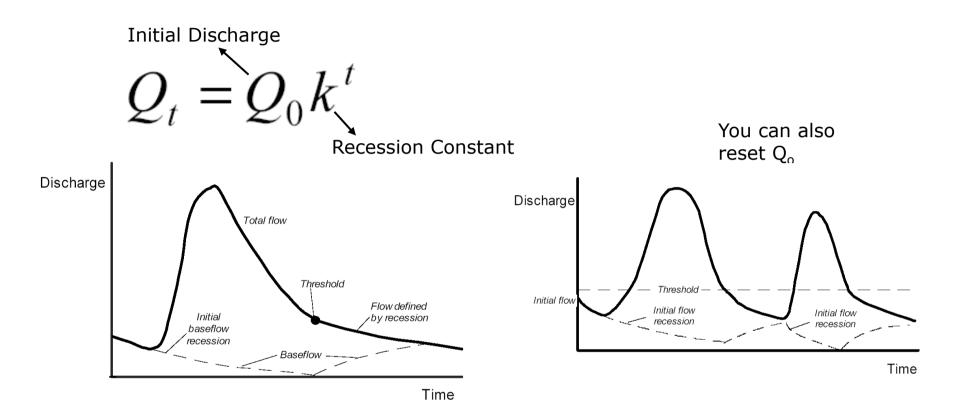
🔐 Subbasin 🛮 Loss	; Baseflow	Options	
Basin Name: Basin 1 Element Name: Subbasin-1			
January (CFS)			
February (CFS)			
March (CFS)			
April (CFS)			
May (CFS)			
June (CFS)			
July (CFS)			
August (CFS)			
September (CFS)			
October (CFS)			
November (CFS)			
December (CFS)			





Recession Models



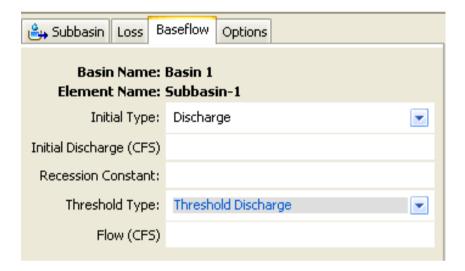


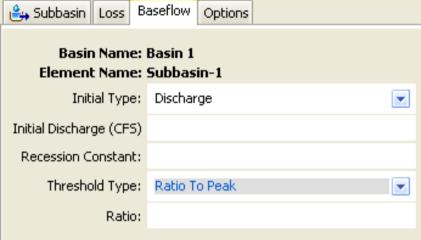




Recession Model





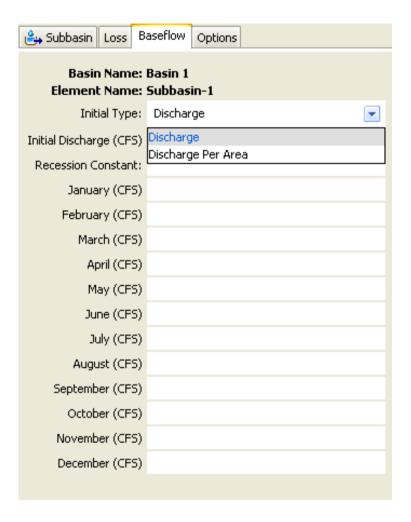






Bounded Recession





Bounded recession is similar to recession method. The difference is mainly in selecting temporally varying threshold.

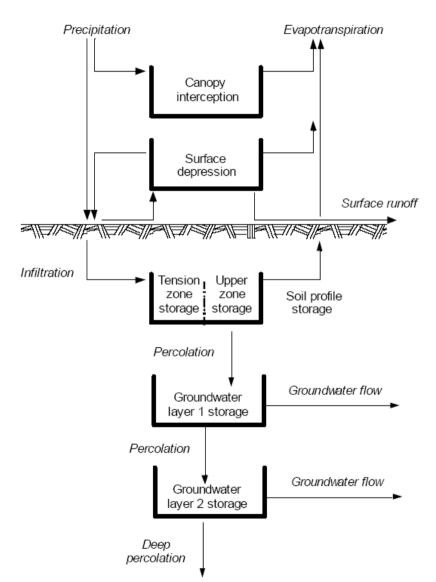
You can also identify the type of initial flow.

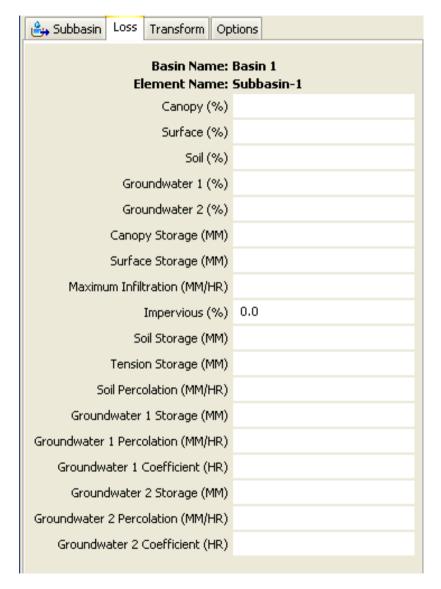




Linear Reservoir (Remember SAC-SMA)





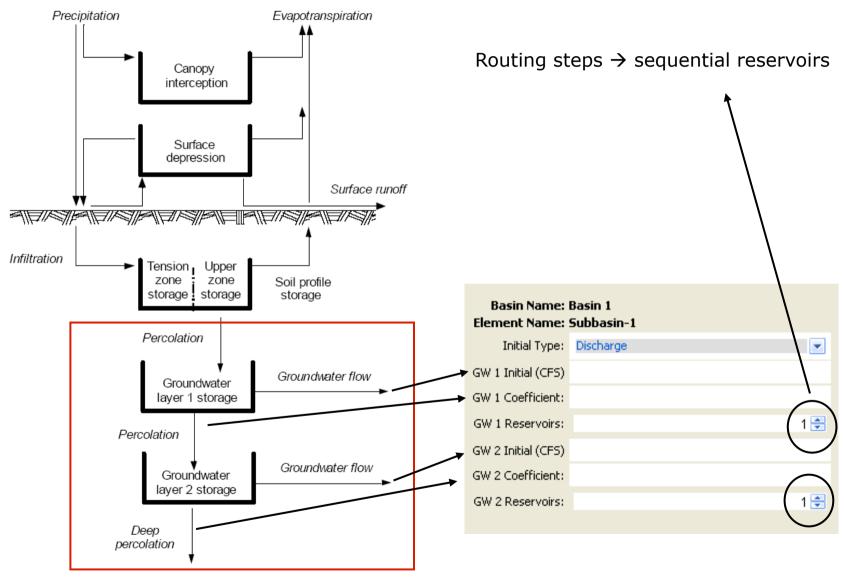






Linear Reservoir



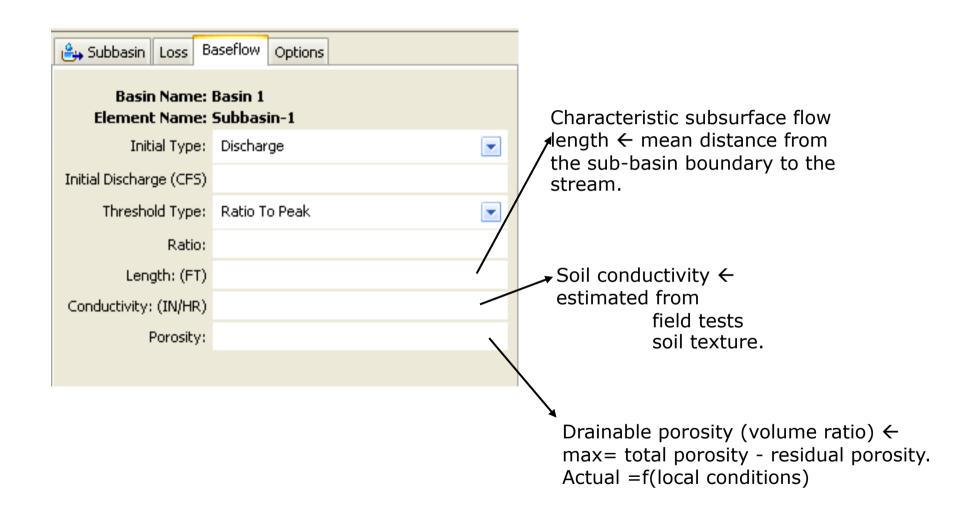






Non-Linear Boussinesq









Basin Model Components (Reach)





Reach

- Represents channels and pipes
- Conveys streamflow downstream in the basin model
- inflow can come from any hydrologic element
- Outflow is computed by accounting for translation + attenuation of inflow hydrograph

Routing

Kinematic wave

Lag

Modified Puls

Muskingum

Muskingum-Cunge







Channel Routing Requirements



Description of the Channel

- Width
- Bed-slope
- Cross-section shape

Energy loss model parameters

- Physically based: manning equation
- Others: parametric

Initial conditions

Flow or stage d/s. Example (use base-flow as estimate)

Boundary conditions

Upstream inflow (determined by the direct runoff model)

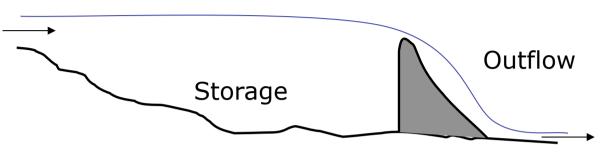




Reservoir Storage Concept

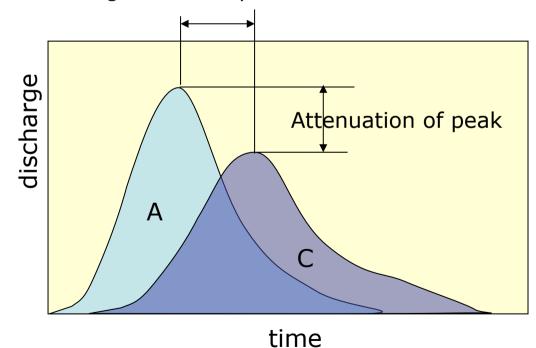


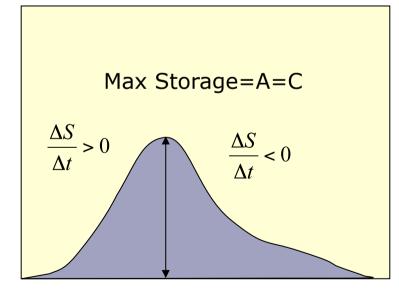




$$I - O = \frac{\Delta S}{\Delta t}$$

Lag of time to peak







Channel Routing Models in HEC-HMS



Reach Routing Options				
Basin Name: Fork_Test Element Name: Reach-1				
Description:		₩		
Downstream:	Junction-2 ▼			
Routing Method:	Lag			
Loss/Gain Method:	None Kinematic Wave Lag Modified Puls Muskingum			
	Muskingum-Cunge Straddle Stagger			





Modified Pulse Model: Simple Storage Routing



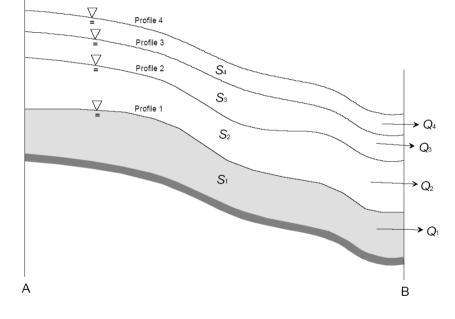
$$\frac{\partial Q}{\partial x} + \frac{\partial A}{\partial t} = 0$$

$$\bar{I}_t - \overline{O}_t = \frac{\Delta S}{\Delta t}$$

$$\left(\frac{S_t}{\Delta t} + \frac{O_t}{2}\right) = \left(\frac{I_{t-1} + I_t}{2}\right) + \left(\frac{S_{t-1}}{\Delta t} - \frac{O_{t-1}}{2}\right)$$

Unknown

$$\frac{1}{2}(I_1 + I_2) - \frac{1}{2}(Q_1 + Q_2) = \frac{S_2 - S_1}{\Delta t}$$



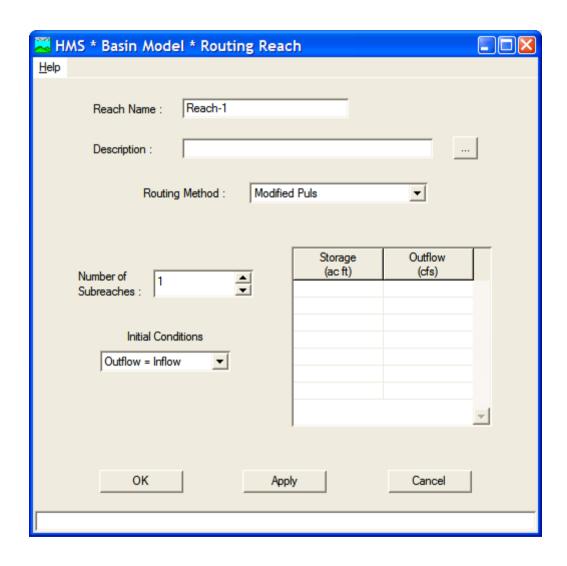
Solution requires Discharge-Storage (Q|S) relationship

HEC-HMS Solves equation recursively using trial and error



Modified Pulse





Notice Storage Outflow requirement





Muskingum Model: The starting point



$$\left(\frac{S_t}{\Delta t} + \frac{O_t}{2}\right) = \left(\frac{I_{t-1} + I_t}{2}\right) + \left(\frac{S_{t-1}}{\Delta t} - \frac{O_{t-1}}{2}\right)$$

Proposed S/I-O relationship

$$S = \frac{b \left[xI^{m/n} + (1-x)O^{m/n} \right]}{a^{m/n}}$$

 $(I,O) \propto ay^n \ (Manning)$

$$S_{reach} \propto by^n$$

a,b,m,n are constants

Assume m/n = 1 and b/a = K

$$S = K [xI + (1 - x)O]$$





Muskingum Model (S/I-O) relationship

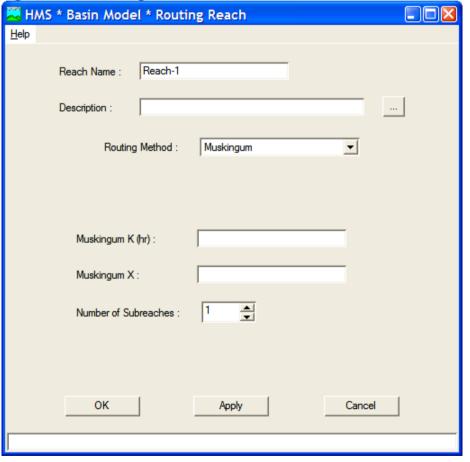


$$S_t = K \left[x I_t + (1 - x) O_t \right]$$

K: Average travel time of flood wave through the routing reach

 $x I_t + (1-x) O_t$ is a weighted discharge. When S is controlled by downstream conditions, thus storage and outflow are highly correlated and x = 0.0 and S =KO (linear reservoir model)

If x = 0.5, equal weight is given to inflow and outflow, and the result is a uniformly progressive wave that does not attenuate as it moves through the reach.

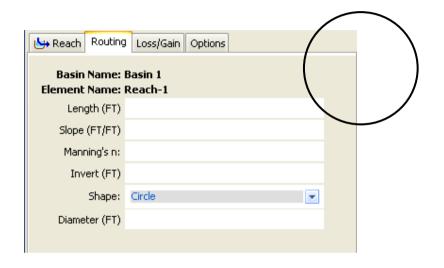


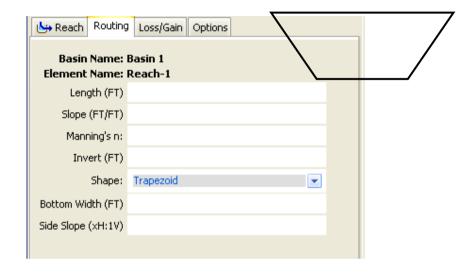


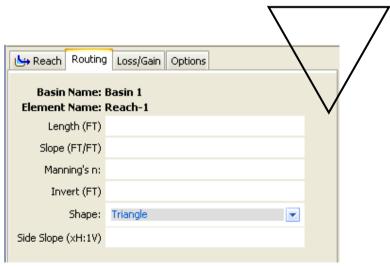


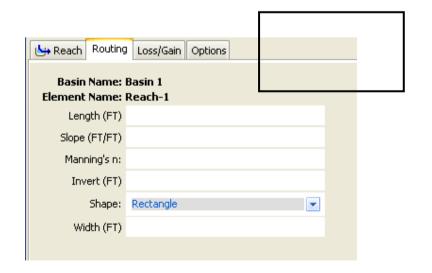
Muskingum Kung Standard









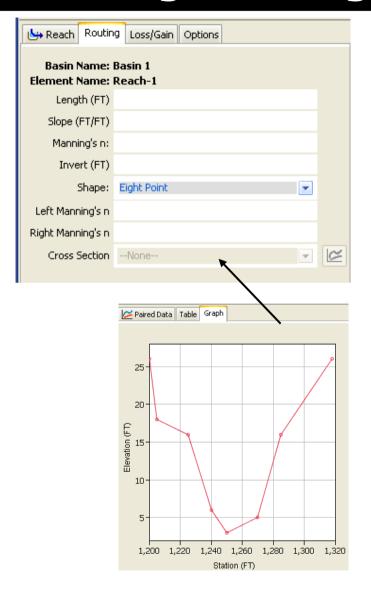


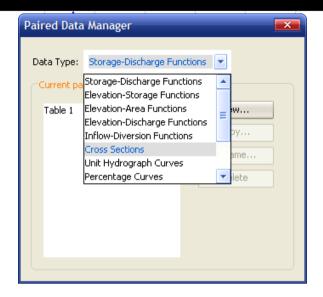




Muskingum Kung 8 Points







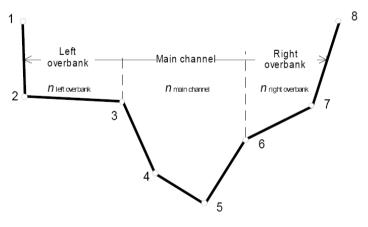


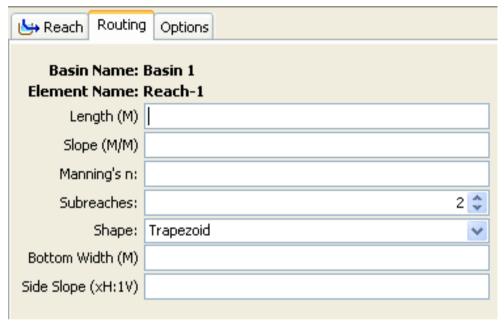
Figure 8-5. Format for describing channel geometry with 8 points





Kinematic Wave





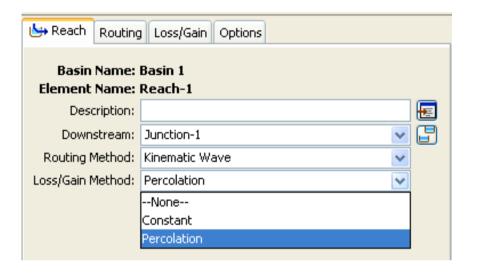
Trapezoid	~
Circle	
Deep Rectangle	
Rectangle	
Trapezoid	
Triangle	

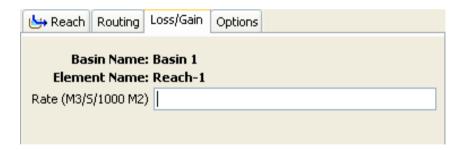


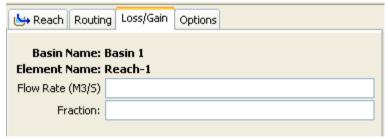


Kinematic Wave









Requires inundated area Not compatible with all methods Requires discharge/Elevation/Area relationships After constant rate is subtracted Remaining flow is further reduced by 1-fraction







