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#### **Advanced School on Direct and Inverse Problems of Seismology**

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Ray Theory (Overheads)

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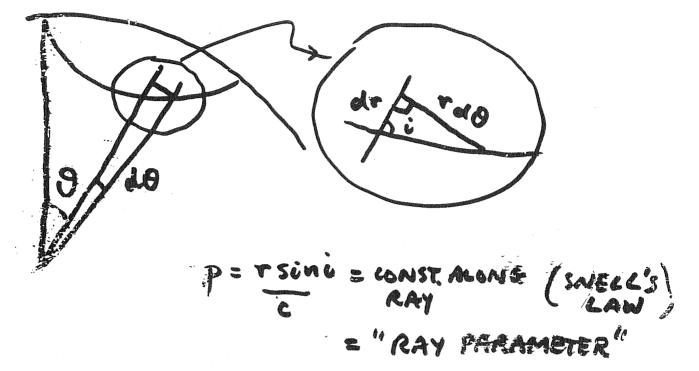
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### CLASSEAL RAY THEORY



C = WAVE SPEED = C(r)

i = ANGLE RAY MAKESLOCALLY

WITH VERTICAL

# EQUATION OF MOTION AND PLANE

#### WAVE SOLUTIONS:

ISOTROPIC HOOKE'S LAW

### PLANE WAVE SOLUTIONS:

CONSIDER WAVE TRAVELLING IN X- DIRECTION

#### P-WAVE

#### S- WAVE

# WHAT IS THE ENERGY FLUX?

WE NEED TO FIND THE RATE OF WORKING OF ONE SIDE OF A PLANE I'X-AXU ON THE OTHER.

LET n be a unit rector in x-direction

n = (1,0,0)

Traction = ti = tijni ds

where ds = element of area

Rate of working = tij njds ili (force x velocity)

=) ENERGY FLUX AVERAGED OVER A CYCLE

=  $\frac{1}{2} |U|^2 \omega k (2+2\mu)$ 

= 1/2 10/2 w2 pa

UNITS: ENERGY PER UNIT TIME PER UNIT AREA. S-WAVE SIMILARLY

 $T_{xy} = \mu ik V e^{i(\omega t - kx)}$ Energy  $\mu x = Re \{ \mu ik V e^{i(\omega t - kx)} \}$  $x Re \{ i \omega V e^{i(\omega t - kx)} \}$ 

ENERGY FLUX AVERAGED OVER A CYCLE

= 1/1/2 wkm

= 1/V/2 W2 pB

THE BASIC IDEH OF THE ASYMPTOTIC OR RAY
THEORIES IS THAT IN MEDIA IN WHICH THE
WAVE VELOCITIES AND DENSITY VARY SCOWLY
WAVES PROPAGATE IN MUCH THE SAME WAY
AS IN HOMOGENEOUS MEDIA.

CONSIDER A P-WAVE PROPAGATING IN THE X-DIRECTION IN A MEDIUM IN WHICH DENSITY AND P-WAVE SPEED ARE ALSO FUNCTIONS OF X. WAVE EQUATION:

is: 
$$\frac{9x}{9}\left(b\alpha_{5}\frac{9x}{9n}\right) = b\frac{9F_{5}}{95n}$$

$$\frac{9x}{9}\left(y+5h\right)\frac{9x}{9n} = b\frac{9F_{5}}{95n}$$

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SEEK AN APPROXIMATE SOLUTION OF THE FORM  $u(x,t) = U(x)e^{i\omega(t-\theta(x))}$ 

U=U(x), O=O(x) TO BE DETERMINED.

WIS CONSIDERED TO BE A LARGE PARAMETER

SUBSTITUTING, WE FIND

$$\frac{\partial u}{\partial x} = \left(-i\omega \frac{\partial \theta}{\partial x}U + \frac{\partial U}{\partial x}\right)e^{i\omega(t-\theta)}$$

$$\frac{\partial}{\partial x}\left(\rho\alpha^{2}\frac{\partial u}{\partial x}\right) = \left\{-\omega^{2}\rho\alpha^{2}U(\frac{\partial \theta}{\partial x})^{2} - i\omega \frac{\partial \theta}{\partial x}\frac{\partial U}{\partial x}\rho\alpha^{2}\right\}$$

$$-i\omega \frac{\partial}{\partial x}\left(\frac{\partial \theta}{\partial x}U\rho\alpha^{2}\right) + \dots\right\}e^{i\omega(t-\theta)}$$

$$-i\omega \frac{\partial}{\partial x}\left(\frac{\partial \theta}{\partial x}U\rho\alpha^{2}\right) + \dots\right\}e^{i\omega(t-\theta)}$$

where ... indicates terms of lower order in w

THUS, FROM w2 terms:

$$\left(\frac{\partial \theta}{\partial x}\right)^2 = \frac{1}{\alpha^2}$$

EQUATION FOR THE PHASE DGU
[EIKONAL
EQUATION]

AND FROM WI TERMS

$$\frac{\partial \theta}{\partial x} \frac{\partial U}{\partial x} p x^2 + \frac{\partial}{\partial x} \left( \frac{\partial x}{\partial x} U p x^2 \right) = 0$$

ie 
$$\frac{\partial}{\partial x} \left( \frac{\partial \theta}{\partial x} U^2 \rho \alpha^2 \right) = 0$$

ie 
$$\frac{\partial}{\partial x} \left( U^2 \rho \mathcal{A} \right) = 0$$
 = CONSTANT ENERGY

3-D THEORY - WORKS SIMILARLY

(KARRL & KELLER, J. Aroust. Soc. Am., 31, 694, 1959)

SEEK ASOLUTION OF EQUIS. OF MOTION IN FORM
: w(t-0(x,y,+))

ui = Ui (=, y, E) e iw (t - 0(=, y, 2))

substitute into equation of motion, identify leading powers of  $\omega$  ( $\omega^2$ ).

DETAILS ARE COMPLEATED.

WE FIND THAT EITHER

$$\theta_{,i} \theta_{,i} = \frac{1}{2}$$
 with  $U_i \parallel \theta_{,i}$ 

OR

THUS WE GET TWO KMIDS OF SOLUTION, CORRESPONDING TO P-WAVES AND TO S-WAVES

THUS, IN BOTH CASES WE OBTAIN FOR

THE "TRAVEL TIME"  $O(\frac{\pi}{2})$  AN EQUATION

OF THE FORM

$$(\nabla \theta)^2 = \frac{1}{c^2}$$
 EQUATION

WHERE C= & FOR P-WAVES, OR C= B
FOR S-WAVES.

IMAYINE A PATH EVERYWHERE 11 70 70

WE HAVE  $0 = 0_0 + \int \frac{1}{2} ds$ 

HOW CAN WE DETERMINE SUCH PATHS?

DIFFERENTIATING THE EIKONAL EGAL

ie. 
$$2 \theta_{i} \theta_{ji} = \frac{\partial}{\partial x_{j}} \left(\frac{1}{c^{2}}\right)$$

but Ori is parallel to 5 ie

$$\theta_{ij} = \frac{1}{\epsilon} \frac{dx_i}{ds}$$

A

$$\frac{1}{\epsilon} \frac{\partial}{\partial s} = \frac{\partial}{\partial c_j} \left( \frac{1}{\epsilon} \right)$$

or 
$$d\theta_{ij} = \frac{c}{2} \frac{\partial}{\partial x_{i}} \left(\frac{1}{c^{2}}\right) = \frac{\partial}{\partial x_{j}} \left(\frac{1}{c}\right)$$

THUS, FROM A &B

$$\frac{dx_i}{ds} = c\theta, i$$

$$\frac{d\theta_i}{ds} = \frac{\partial}{\partial x_i} \left(\frac{1}{c}\right)$$

RAY-TRACING EQUATIONS ALTERNAVELY, WRITING

$$\frac{d}{ds} = \frac{1}{1} \frac{d\theta}{ds}$$

WE GET

$$\frac{dx_{i}}{d\theta} = \frac{c^{2}}{\omega}k_{i} = c\frac{\partial c}{\partial x}$$

$$\frac{dk_{i}}{d\theta} = \omega c\frac{\partial}{\partial x_{i}}\left(\frac{1}{c}\right) = -\frac{\omega}{c}\frac{\partial c}{\partial x_{i}} = -k\frac{\partial c}{\partial x_{i}}$$

$$\frac{\partial c}{\partial x_{i}} = c\frac{\partial c}{\partial x_{i}}$$

$$\frac{\dot{c}}{\dot{c}} = -k\frac{\partial c}{\partial x_{i}}$$

$$\frac{\partial c}{\partial x_{i}} = -k\frac{\partial c}{\partial x_{i}}$$

THESE REPRESENT THE MOTION OF A "PARTICLE" TRAVELLING AT THE LOCAL WAVE SPEED C, SUFFERING DEFLECTIONS FROM A STRAIGHT-LINE TRAJECTORY DUE TO VELOCITY GRADIENTS THAT ARE NOT !! TO THE PATH

A GENERAL WAY OF UNDERSTANDING
THE RAY EQUATIONS IS THROUGH THE
CONCEPT OF THE LOCAL DISPERSION
RELATION BY WHICH WE SHALL
MEAN THE RELATION BETWEEN
FREQUENCY W (=2TT/PERIOD)
AND WAVE-VECTOR R

[IR] = 2TT/WAVELENGTH).

THE WAVE VECTOR FOR A WAVE OF THE FORM ((WE-4(E)))
U e

CAN BE DEFINED AS

R; = 04

3xi

THE LOCAL DISPERSION RELATION IS THEN

GIVEN BY A FUNCTION W(Ri, xi),

SO THE PHASE Y(X) SATISFIES AN

EQUATION OF THE FORM

 $\omega = \omega(\frac{\partial Y}{\partial x_i}, x_i)$ 

THE METHOD OF CHARACTERISTICS (ESSENTIALLY

THE METHOD GIVEN ABOVE) THEN LEADS TO HAMILTON'S EQUATION

GIVEN A LOCAL DISPERSION RELATION

$$\omega = \omega(R_{i}, z_{i})$$

THE RAY EQUATIONS ARE

$$\dot{x}_{i} = \frac{\partial \omega}{\partial k_{i}}$$

$$\dot{k}_{i} = -\frac{\partial \omega}{\partial x_{i}}$$

cf. Hamilton's EQNS. FOR A
MECHANICAL SYSTEM:
GIVEN THE HAMILTONIAN
H (Pi, qi)
THE EVOLUTION OF THE SYSTEM
IS GOVERNED BY

H(P,q) X=> w(k, x

q: "GENERALISED "

Pi = "GENERALISED MOMENTA-" LET US USE THIS IDEA TO RE-DERIVE THE RAY EQUATIONS.

THE LOCAL DISPERSION RELATION IS DE

FOR BODY WAVES IN AN ISOTROPIC MEDIUM (c=d or c=b)

# : HAMILTON'S EQUATIONS GIVE

$$\dot{x}_{i} = c \, k_{i}$$

$$\dot{R} \qquad \qquad \begin{cases}
THE SAME \\
AS DERIVED \\
EARLIER
\end{cases}$$

$$\dot{\partial} x_{i} = -k \, \partial c$$

$$\dot{\partial} x_{i} = -k \, \partial c$$

wim k = (kiki) = 1 k /

FOR AN ANISOTROPIC MEDWM.

WE HAVE

(cijke UR, e), j +w2u; =0

=> -ikj cijke (-ike) uk +w² ui:0

ie (Cijke ke kj -w² Sin) Un = 0

THUS THE LOCAL DISPERSION RELATION

15 Det (Cijkekeki -wi Sin) = 0

THE DERIVATIVES DW DW CAN BE DOC: DK;

FOUND FROM STANDARD PERNRBATION
THEORY (RAYLEIGH'S PRINCIPLE)

WE FIND

Rm = - DW = - 1 DCijes Vi UR Ke kj

zm =  $\frac{\partial \omega}{\partial k_m} = \frac{1}{2\omega} \left( \text{Cijkm kj} + \text{Cimkeke} \right) \text{Jijk}$ 

Where Vi is a (local) unit eigenvector (CORRESPONDING TO THE WAVE OF INTEREST)

PANOTHER ELECANT PROPERTY OF HAMILTON'S EQUATIONS IS THAT THEY CAN BE WRITTEN DOWN IN ANY COORDINATE SYSTEM

SUPPOSE THAT WE WART TO DO 3-D
RAY TRACING IN SPHERICAL COURDINATES



(r, 0, 4)

We have  $R_r = \frac{24}{57}$ ,  $R_\theta = \frac{24}{59}$ ,  $R_{\phi} = \frac{24}{59}$  and  $R_{\phi} = (R_r^2 + \frac{1}{72}R_{\phi}^2 + \frac{1}{725m^2\theta}R_{\phi}^2)^{\frac{1}{8}}$ 

with the usual dispersion relation

$$\omega = c(r, 0, \phi)k$$

WE OBTAIN RAY-TRACING EQUATIONS:

$$\dot{k}_{\theta} = -\frac{\partial c}{\partial \theta} k + \frac{\cot \theta}{k r^2 \sin^2 \theta} k_{\phi}^2$$

$$\dot{k}_{\phi} = -3c$$

TO MAKE CONTACT WITH CLASSICAL RAY THEORY IN THE SPHERICAL EARTH LET US NOW SIMPLIFY THESE FOR THE CASE C: C(V) TAKE SOURCE AT 0=0, K&=0

$$\dot{\tau} = krc$$

$$\dot{k}r = -\frac{\partial c}{\partial r}k + \frac{1}{kr^0}k^2$$

$$\dot{\theta} = \frac{1}{r^2}\frac{k\theta c}{R}$$

$$\dot{k}\theta = 0$$

$$\dot{k}\psi = 0$$

$$\omega = c(k_r^2 + \frac{1}{2}k_0^2)^2 = const$$

ie 
$$P_{\tau} = \left(\frac{1}{e^2} - \frac{P^2}{\tau^2}\right)^{\frac{1}{2}}$$
  $\left(P \equiv P_{\theta}\right)^{\frac{1}{2}}$   $= \text{"RAY}$ 

PARAMETER

# THUS WE OBTAIN THE CLASSICAL RAY INTEGRALS

$$E = \int \frac{1}{2} \left( 1 - \frac{1}{2} \frac{1}{2} \right)^{-1/2} dr$$

$$O (=0) = \int \frac{1}{2} \left( 1 - \frac{1}{2} \frac{1}{2} \right)^{-1/2} dr$$

# AMPLITUDES AND WAVEFORMS

BECAUSE RAY THEORY (FOR BODY WAVES)
IS FREQUENCY - INDEPENDENT, IT
PREDICTS THAT WAVES PROPAGATE
WITHOUT ANY CHANGE TO THE WAVEFORM

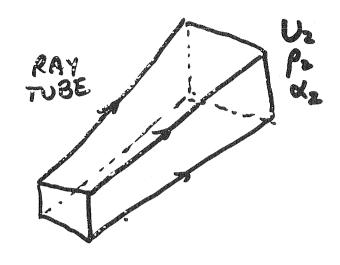
( JUST AS IN A HOMOGENEOUS MEDIUM)

THE ASYMPTOTIC THEORY CAN BE USED TO DERIVE WAVE AMPLITUDES (BY INVESTIGATING THE TERMS & \omega)
THE DERIVATION WILL NOT BE GIVEN HERE (SEE LITERATURE) THE RESULT IS

THAT ENERGY FLUX IN A RAY TUBE IS CONSTANT

RECALLING THAT

THIS MEANS THAT RAY AMPLITUDES VARY
INVERSELY AS NOW AND ALSO AS 1/VA'
WHERE A IS HA CROSS-SECTIONAL AREA
OF THE RAY TUBE.



$$U_1^2 P_1 \alpha_1 A_1 = U_2^2 P_2 \alpha_2 A_2$$
ie  $U_2 = U_1 \sqrt{\frac{P_1 \alpha_1 A_1}{P_2 \alpha_2 A_2}}$ 

(SIMILARLY FOR S-WAVES WITH B INSTEAD
OF &)

TRAVEL TIME IS STATIONARY WITH RESPECT
TO PERTURBATIONS OF THE PATH

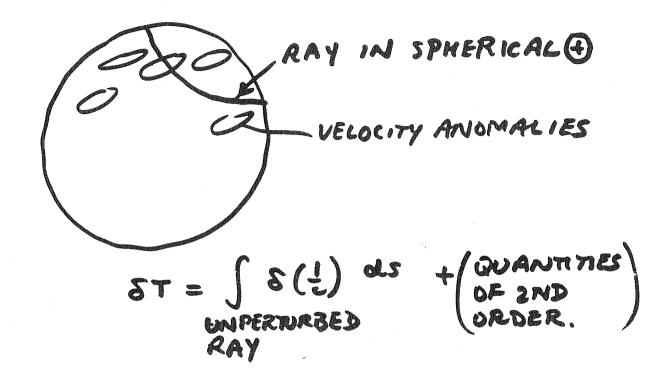
T' = "TRAVEL TIME CALCULATED ALONG THE NON-RAY "

$$= \int \frac{1}{c(x+\epsilon)} ds'$$

$$= \int V(\frac{1}{c}) \cdot \epsilon ds + \int \frac{1}{c} \frac{ds'}{ds} ds + O(\epsilon^{2})$$
But  $\frac{ds'}{ds} = \left\{ \frac{d}{ds} (x_{i} + \epsilon_{i}) \right\} \frac{d}{ds} (x_{i} + \epsilon_{i}) \right\}^{\frac{1}{2}}$ 

$$= 1 + \frac{dx}{ds} \cdot \frac{d\epsilon}{ds}$$

# APRICATION IN TOMOGRAPHY

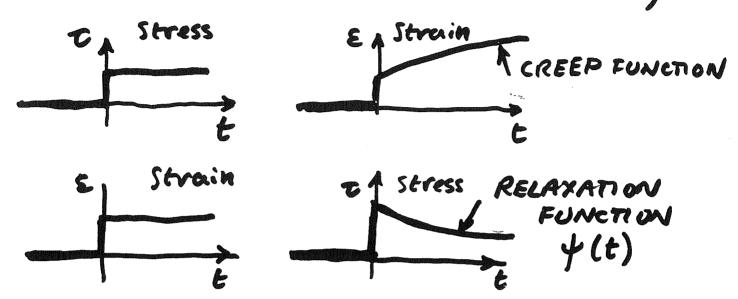


NOTE THAT IT IS NOT TRUE

THAT THE PERTURBATION OF THE RAY
PATH IS 2ND ORDER.

# ATTENUATION AND PHYSICAL DISPERSION OF SEISMIC WAVES

(RECALL DR. YANOUSKAYA'S LECTURES & NOTES)



FOR ASMUSOIDAL STIEAR DISTURBANCE

W = Ue iw E

$$\tau(t) = \mu(\omega) \epsilon(t)$$

WHERE  $\mu(w)$  is complex and

# FREQUENCY DEPENDENT

FSIMILARLY FOR COMPRESSION

T(t) = K(W)E(t)

WRITING F(W) = \frac{1}{2\pi} \begin{picture}(c) & \text{4(t)} & \end{picture}

WRITING \frac{1}{2\pi} \begin{picture}(c) & \text{4(t)} & \end{picture}

IT IS EASY TO SEE THAT

$$\mu(\omega) = i\omega \Psi(\omega)$$

IT IS CONVENTIONAL TO DEFINE

BUT OFTEN MURE CONVENIENT TO USE

Writing 
$$\frac{1}{\sqrt{s}} = \sqrt{\frac{2}{\mu_1(\omega)}} = s_1 - is_2$$

THUS THE EXPRESSION FOR A PLANE WAVE TRAVELLING IN THE X-DIRECTION IS OF THE FORM

$$u \sim V_0 e^{i\omega(t-x/v_s)}$$

$$= V_0 e^{-\omega x s_2} e^{i\omega(t-xs_i)}$$

with 52 = Re(3). 29/

DECAY IN ONE WAVELENGTH

$$exp\{-\omega \frac{2\pi}{\omega s_i} \frac{1}{2}q_m s_i\} = exp(-\pi q_m)$$

AMPLITUDE DECAY FOR S-WAVE

Qp IS ALSO SOMETIMES DENOTED BY

QB (= Q FOR S-WAVES)

CORRESPONINGLY

AMPLITUDE DECAY FOR P-WAVE

-TT/Q2

PER CYCLE

where Qx = -2Re(1/Up)
Im(1/Up)

OF COURSE WE HAVE

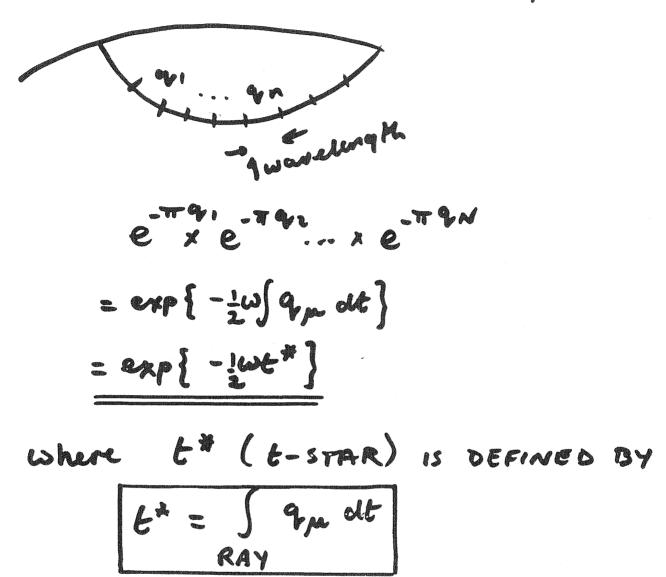
Up = \( \frac{K + 4/3 \text{ M}}{\text{ P}} \)

 $K = (Re K)(1 + iq_K)$  etc.

SO IT IS EASY TO FIND EXPRESSIONS FOR Qu IN TERMS OF Qx, Qp. IN PARTICULAR

WE OBTAIN

# WITHIN RAY THEORY THIS LEADS TO AN ADDITIONAL AMPLITUDE DECAY



FREQUENCY WAVES

#### PHYSICAL DISPERSION

WE SAW THAT

μ(ω) = iω ψ(ω)

WHERE  $\Psi(\omega) = F.T.$  OF RELAXATION FUNCTION  $\Psi(t)$ 

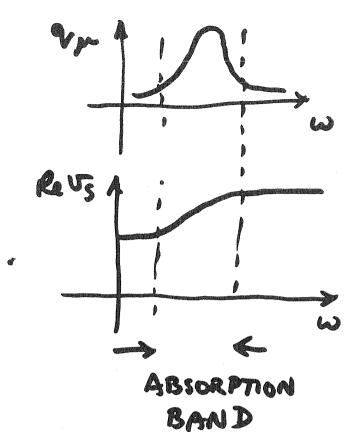
WAVE VELOCITY ( \sum\_{\beta} ) IS RELATED TO Re(\mu) AND DAMPING TO IM (\mu).
BUT SINCE \( \mu(\omega) \) IS THE TRANSFORM OF A SINGLE REAL (CAUSAL) FUNCTION Re(\mu) AND \( \mu(\omega) \) ARE RELATED.

EG. FOR THE STANDARD LINEAR SOLID (SEE "WAVE PROPAGATION" NOTES FROM DR. YANOUSKAYA)

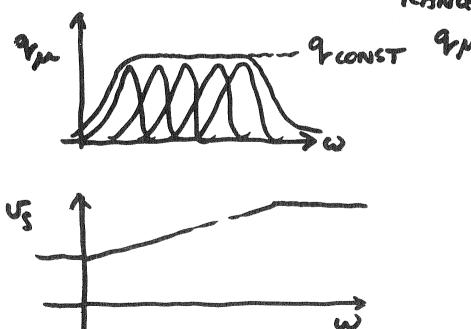
$$\begin{array}{ccc}
\tau + T_{z} \dot{\tau} &= \mu_{0} (\varepsilon + T_{\varepsilon} \dot{\varepsilon}) \\
\Rightarrow & \mu(\omega) &= \mu_{0} (1 + i\omega T_{\varepsilon}) \\
\hline
1 + i\omega T_{z}
\end{array}$$

This can be used to find both quew)

AND  $U_S(\omega) = Re\sqrt{M(\omega)}$ 



THUS VS INCREASES THROUGH THE
ABSORPTION BAND. FOR MANY ABSORPTION
BANDS US INCREASES THROUGHOUT THE
RANGE OF CONSTANT



QUANTITYELY IT CAN BE SHOWN THAT APPROXIMATELY, AND WITHIN THE BAND OF CONSTITUT

OR (INTEGRATING) FOR WI, WE WITHIN THE BRAND

$$\ln \frac{U_s(\omega_z)}{U_s(\omega_1)} = \frac{1}{\pi} q_{\mu} \ln \left(\frac{\omega_z}{\omega_1}\right)$$

THESE LEAD TO A RELATIONSHIP BETWEEN
THE DELAY OF AWAVE OF GIVEN FREQUENCY
AND 2T'S DECAY. THE PHENOMENON IS KNOWN AS
PHYSICAL DISPERSION

SEE LIU, ANDERSON, KANAMORI, GJ. 1976
AND REFERENCES CITED THEREIN]

WE CAN ALSO WRITE FOR THE COMPLEX VELOCITY

WHERE UD IS THE (REAL) VELOUTY AT REFERENCE FREQUENCY WO

CONSEQUENTLY THE EFFECT ON THE SIGNAL IS REPRESENTED BY

$$exp\{-\frac{1}{2}\omega t^*(1-\frac{2i}{\pi}\ln\frac{\omega}{\omega_0})\}$$

THIS REPRESENTS (APPROXIMATELY,
AND) ASSUMING THAT THE ENTIRE
SIGNAL IS WITHIN THE CONSTANT QUE
BAND) THE TOTAL AFFECT OF
ATTENUATION ON THE SIGNAL.