Stream Function and Mean Meridional circulation Variability

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Objective

 Understanding Basic features of General Circulation

What do we mean by circulation?

Outlines

- Introduction to General Atmospheric Circulation
- Hadley Cell
- Ferrel Cell





Meridional circulation (v,w)





Are not useful to see the circulation

Stream Function (Ψ)

The mass divergence in the meridional plane is zero (conservation of mass)

The stream funciton is a scalar field whose relationship to V is carefully selected to automatically satisfied the continuity $\begin{aligned}
\psi_p(\phi, p) &= \frac{2\pi\cos\phi}{g} \int_p^{p_s} [v] dp
\end{aligned}$ $\begin{aligned}
[v] &= \frac{g}{2\pi\cos\phi} \frac{\partial\psi}{\partial p} \\
[\omega] &= -\frac{g}{2\pi\cos\phi} \frac{\partial\psi}{\partial \phi}
\end{aligned}$

Properties of Stream Function:

- 1. Stream function always satisfies the continuity equation.
- 2. Stream function is constant for a stream line.

3. Stream function values vary from one stream line to another.

4. Stream function represents volumetric flux .i.e. volumetric flow rate per unit area. Units of stream function are kg/ms.

Zonal Average Views

Zonal Average of x = [x]x - $[x] = x^* =$ deviation from the zonal average u-[u] = u^{*} v-[v] = v^{*} t-[t] = t^{*}

Zonal mean Stream function NCEP Mass Stream Function DJF ERA40 Mass Stream Function DJF 100 \mathbb{N} 100 °o'r 200 --20 200 300 -300 -400 -20 400 500 -500 -600 · 600 · 700 · 700 0 800 -800 -20 -20 900 -900 --20 1000 | 905 1000 | 905 60S 30S 30N ΕQ 6ÓN 9ÓN 30S 30N 6ÓS ΕQ 6ÓN 90N

$$[\omega] = -\frac{g}{2\pi\cos\varphi} \frac{\partial\psi}{\partial\varphi}$$

$$[v] = \frac{g}{2\pi\cos\varphi} \frac{\partial\psi}{\partial p}$$



Thermally induced cell Indirect cell



Hadley Cell Standard deviation





Oort and Yienger 1996 J. of Climate

Hadley cell Standard deviation shows maximum variability however, Ferrel cell standard deviation shows minimum variability. Since geostrophic balance and hydrostatic balance mostly hold all the time in midlatitude.

Ferrel Cell (indirect cell) Eddy transport

Departure from mean zonal flow (subtropical jet stream) destroys thermal wind balance, which induces the meridioanl circulation to maintain thermal wind balance again

$$\frac{\partial^2 \overline{\chi}}{\partial y^2} + \frac{f_{\circ}^2}{N^2} \rho_{\circ} \frac{\partial}{\partial z} \left(\frac{1}{\rho_{\circ}} \frac{\partial \overline{\chi}}{\partial z} \right) = \frac{\rho_{\circ}}{N^2} \left[\frac{\partial}{\partial y} \left(\frac{\kappa \overline{J}}{H} - \frac{R}{H} \frac{\partial}{\partial y} \left(\overline{v'T'} \right) \right) \right]$$
$$-f_{\circ} \left(\frac{\partial^2}{\partial z \partial y} \left(\overline{u'v'} \right) - \frac{\partial \overline{X}}{\partial z} \right)$$

Diagnosis equation for stream function





References

1-J. Holton, An Introduction to Dynamical Meteorogology

2-D. Strauss, presentation

3-Orrt and Yienger, *observed interanual in the hadley circulation and its connection to* ENSO, J. of Climate, 1996

Thanks for your Attentions