



**The Abdus Salam  
International Centre for Theoretical Physics**



**2210-13**

**MedCLIVAR Workshop on: "Scenarios of Mediterranean Climate  
Change under Increased Radiative Active Gas Concentration and the  
Role of Aerosols**

*23 - 25 September 2010*

**Measurements of Mediterranean aerosol radiative forcing and the intense Saharan  
dust event of March 25-26, 2010, at Lampedusa**

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# Measurements of Mediterranean aerosol radiative forcing and the intense Saharan dust event of March 25-26, 2010, at Lampedusa

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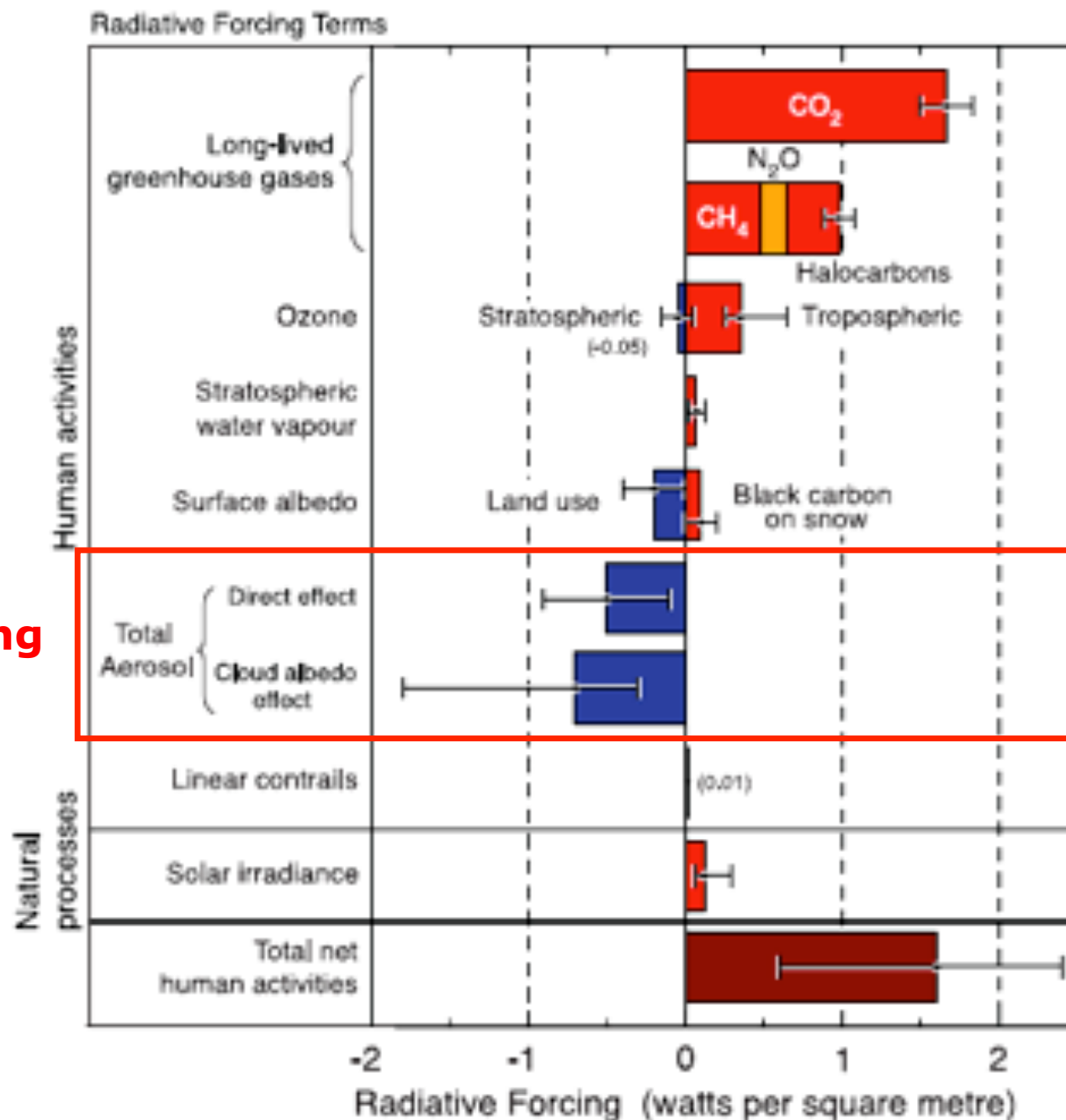
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*<sup>4</sup> ENEA/UTMEA-TER, Lampedusa, Italy*

***MedCLIVAR Workshop on Scenarios of Mediterranean Climate Change, Trieste 23-25 September 2010***

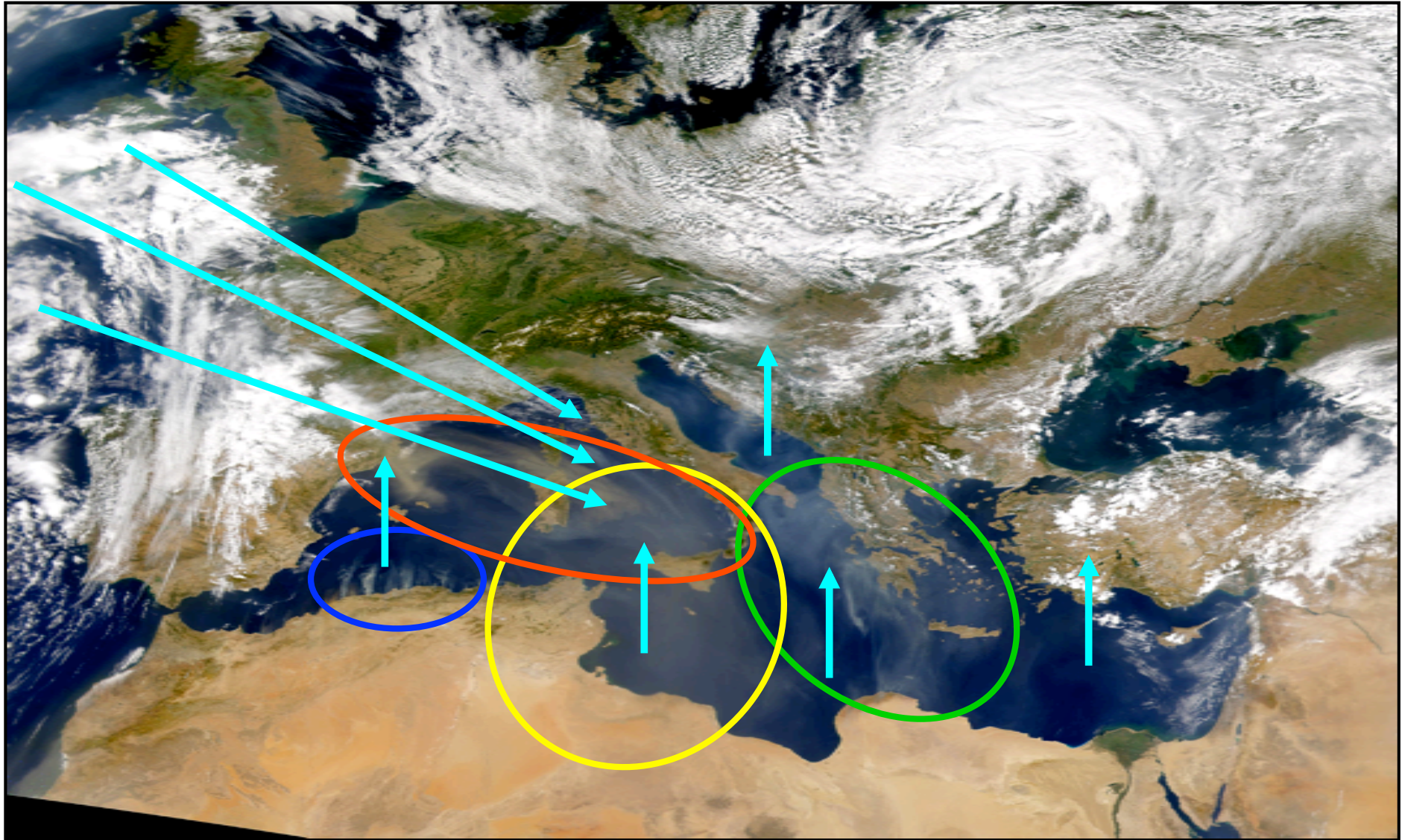


## Radiative forcing of climate between 1750 and 2005



**Aerosol radiative forcing**

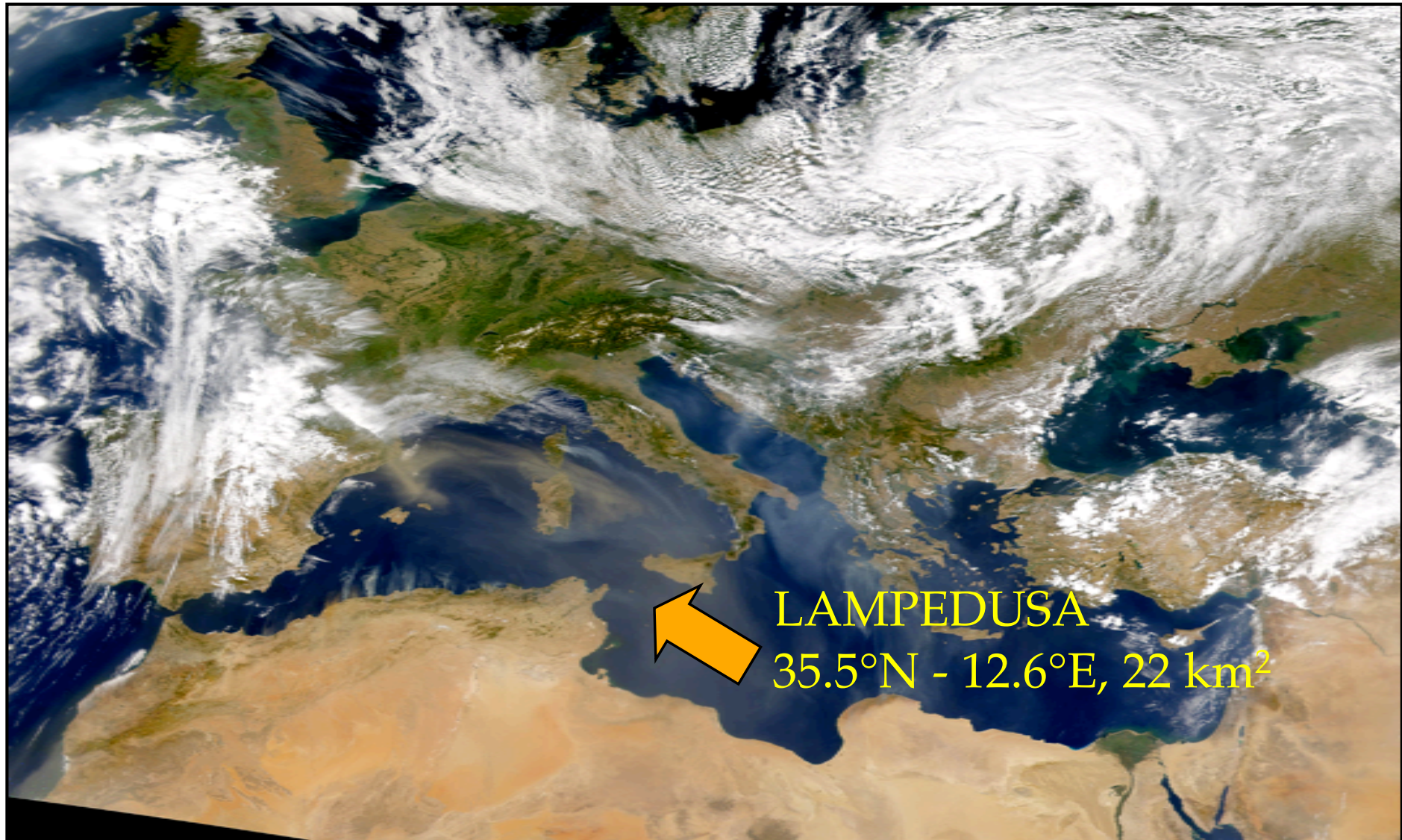
# Aerosols in the Mediterranean region



25/08/2005 (SeaWIFS Image)



# Aerosols in the Mediterranean region



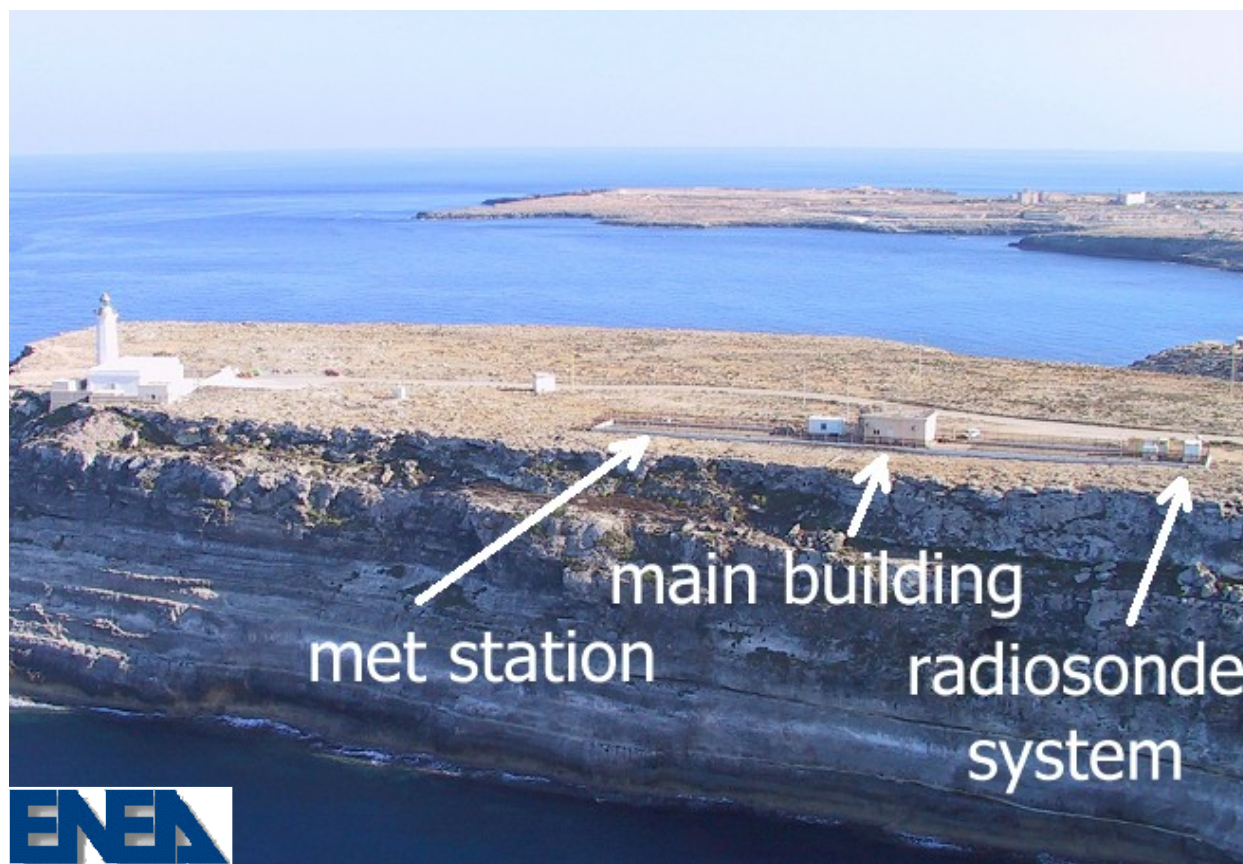
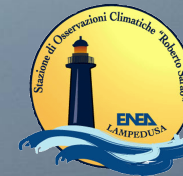
25/08/2005 (SeaWIFS Image)



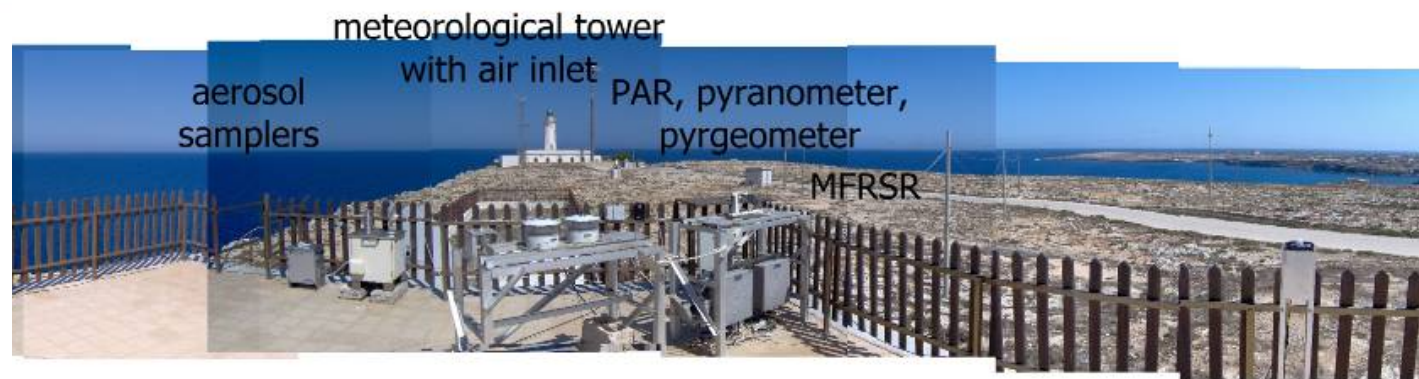
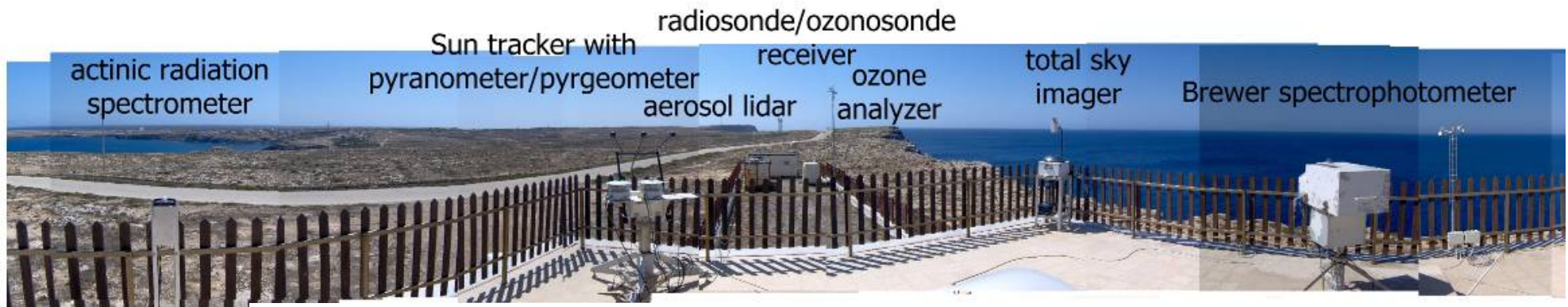
# ENEA Station for Climate Observation, since 1999



<http://www.palermo.enea.it/lampedusa/>







- Meteorological station [air pressure, temperature, humidity, wind direction and velocity, precipitation (Vaisala); solar irradiance (Kipp and Zonen)].
- Non-dispersive Infra-red (NDIR) analyzer [atmospheric CO<sub>2</sub> concentration (the system includes a Siemens 5E analyzer)].
- Gas chromatograph [atmospheric concentration of CH<sub>4</sub>, N<sub>2</sub>O, CFC-11 and CFC-12 (HP 6890)].
- Brewer MK III spectrophotometer [total ozone, spectral UV, aerosol optical depth].
- Aerosol lidar [together with University of Rome; aerosol backscattering and depolarization profiles].
- Visible Multi Filter Rotating Shadowband Radiometer [MFRSR; aerosol optical depth at several wavelengths, diffuse-to-direct irradiance ratio, column water vapor, aerosol single scattering albedo (Yankee Environmental Systems MFR-7)].
- PM-10 aerosol sampler [Tecora Skypost, daily chemical analyses performed at the University of Florence].
- Precision Spectral Pyranometer [downward shortwave irradiance (Eppley)].
- Precision Infrared Pyranometer [downward longwave irradiance (Eppley)].
- CGR4 [downward longwave irradiance (Kipp and Zonen)].
- Shaded Precision Spectral Pyranometer [diffuse downward shortwave irradiance (Eppley)].
- Photosynthetic radiation radiometer [downward photosynthetically active radiation].
- Actinic radiation spectrometer [actinic radiation spectra, photo dissociation rates (Metcon GmbH)].
- F-RAD UV filter radiometer [UV irradiance at 7 bands].
- Total sky imager [cloud cover (Yankee Environmental Systems TSI 440)].
- Water vapor Raman lidar [day/nighttime vertical profiles of water vapor, aerosol extinction (jointly with University of Rome)].
- Vaisala radio/ozonesonde [temperature, pressure, humidity, wind, ozone vertical profiles (Vaisala Digicora III)].
- Hat-Pro Microwave Radiometer [temperature and water vapor vertical profiles, integrated water vapour, liquid water content, cloud base height].
- Cimel sun photometer [aerosol optical depth and optical properties (jointly with University of Modena and Reggio Emilia)].
- SODAR [wind vertical profiles, three components, ERSE].
- ENEA gas sampling unit [weekly analyses of 15 different halogen compounds, made at ENEA, Rome].
- NOAA gas sampling unit [weekly analyses of CO<sub>2</sub>, CH<sub>4</sub>, SF<sub>6</sub>, CO, <sup>13</sup>C, H<sub>2</sub>, <sup>18</sup>O, made at NOAA].

a) MFRSR Radiometers  
b) Cimel photometer (<http://aeronet.gsfc.nasa.gov/>)

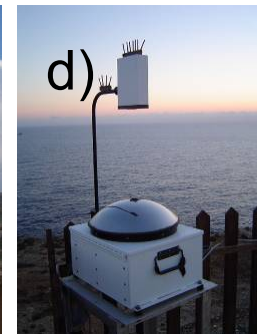
## Aerosol optical properties

c) Solar tracker with PIR (infrared) and PSP (solar) radiometers  
e) CGR4 pyrgeometer (infrared)  
f) CMP21 pyranometer (solar)  
g) PIR and PSP radiometers

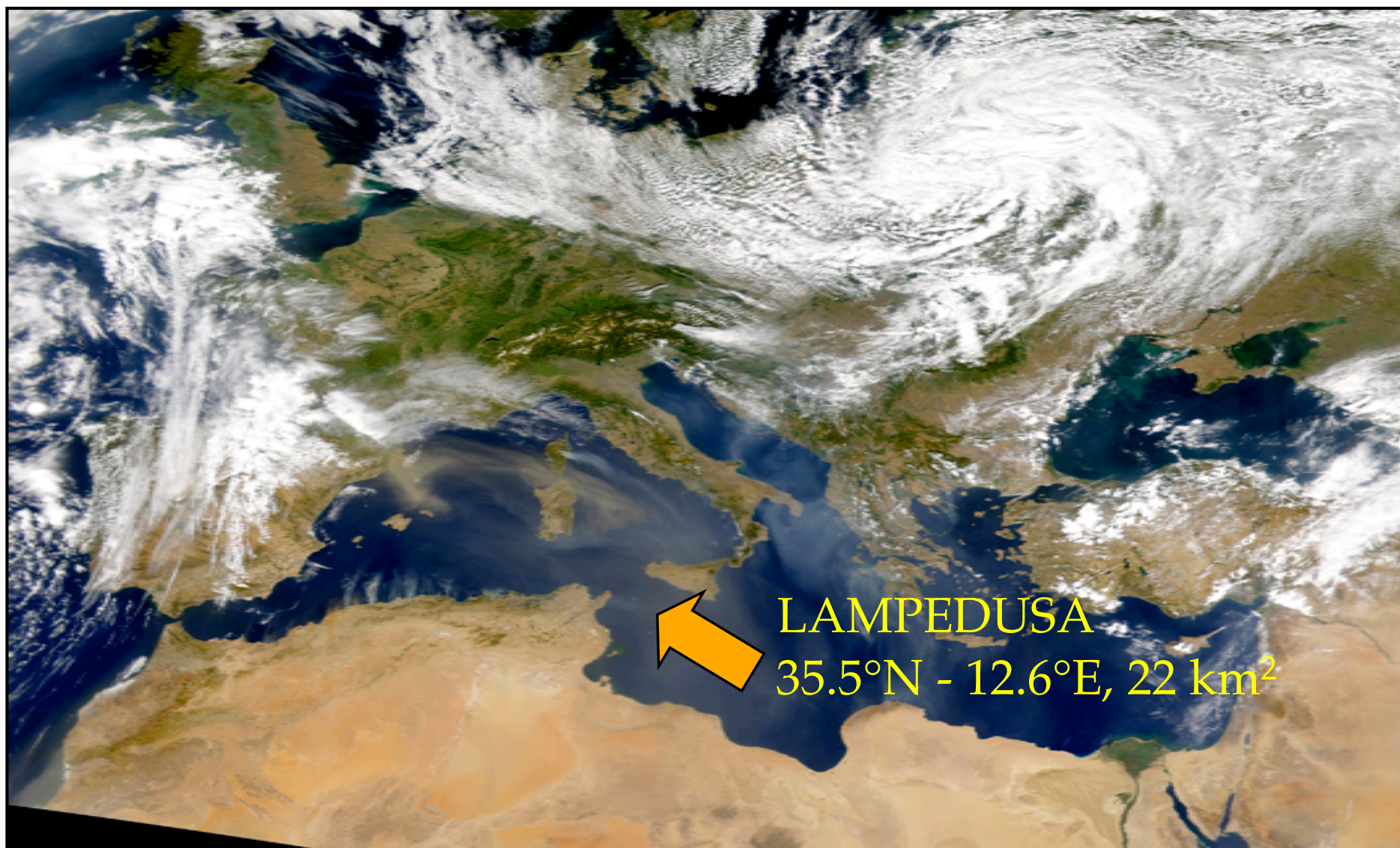
## Radiative fluxes at the surface (global/diffuse sw and lw)

d) Total Sky Imager (TSI)  
h) Meteorological tower

## Meteorology and sky conditions



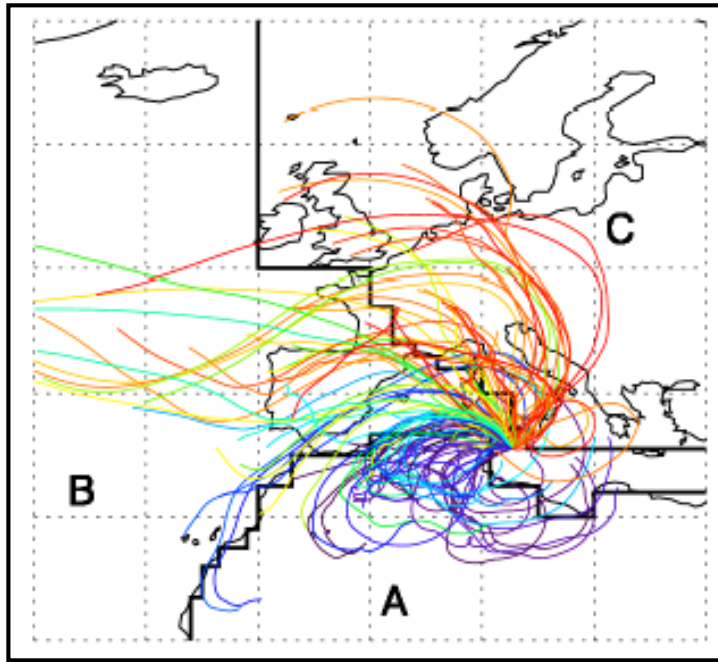




LAMPEDUSA  
35.5°N - 12.6°E, 22 km<sup>2</sup>

25/08/2005 (SeaWiFS Image)

# Different aerosol types at Lampedusa



*Pace et al., ACP 2006*

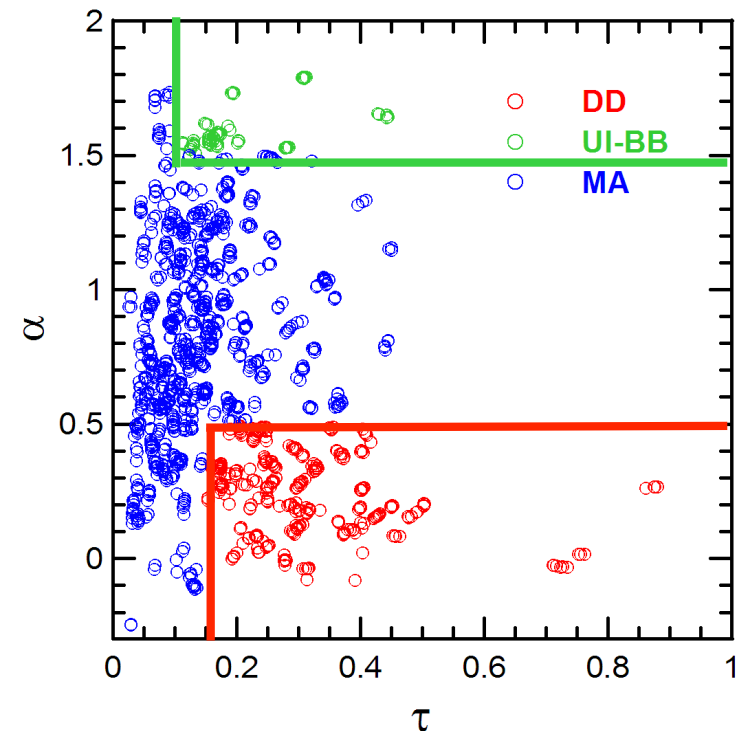
Backward airmass trajectories (HYSPLIT)  
+  
Aerosol optical properties



**DD**: desert dust  
**UI-BB**: urban/industrial-biomass burning  
**MA**: mixed aerosols (marine and mixed)

2004 - 2007

Aerosol type	$\tau$ (495.7 nm)	$\alpha$
<b>DD</b>	$0.31 \pm 0.13$	$0.25 \pm 0.15$
<b>UI-BB</b>	$0.21 \pm 0.09$	$1.60 \pm 0.09$
<b>MA</b>	$0.14 \pm 0.08$	$0.85 \pm 0.37$





# Aerosol Radiative Forcing (RF) at Lampedusa

## **2004-2007**

*Shortwave* RF at TOA (top of atmosphere) - surface - atmosphere

DD – desert dust

UI-BB – urban/industrial-biomass burning aerosols

MA – mixed aerosols (marine + DD + UI-BB)

surface data + satellite data

## **2010: the most intense desert dust event ever observed at Lampedusa since 1999**

*Shortwave and Longwave* RF at TOA - surface - atmosphere

DD – desert dust

surface data + satellite data + radiative transfer model

# How to derive the aerosol Radiative Forcing (RF)?

At a given atmospheric level  
At a given spectral band

$$RF = F_{net} - (F_{net})_{\text{aerosol-free}}$$

radiative transfer model

Forcing Efficiency  $\longrightarrow$   $FE = \frac{d(F_{net})}{d\tau}$

[Satheesh and Ramanathan, NATURE 2000]

Shortwave FE

$$FE_S = \frac{d[(1 - A_S) \cdot I_{SW\downarrow}]}{d\tau}$$

$$FE_{TOA} = \frac{d[-I_{TOA\uparrow}]}{d\tau}$$

$$FE_{ATM} = FE_{TOA} - FE_S$$

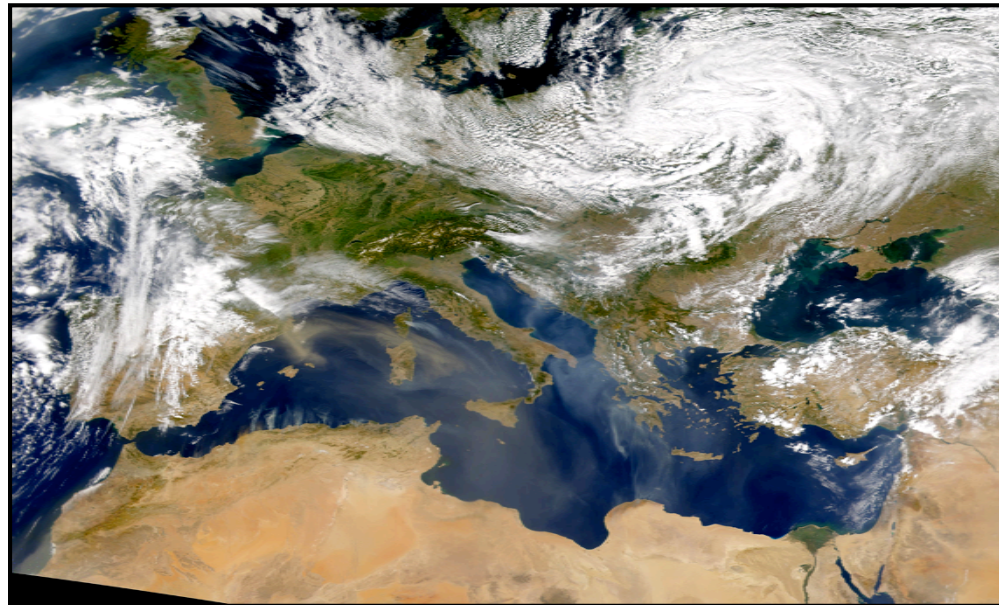
$\tau$  = aerosol optical depth  
 $A_S$  = surface shortwave albedo  
 $I_{SW\downarrow}$  = surface downward shortwave irradiance  
 $I_{TOA\uparrow}$  = TOA upward shortwave irradiance

only observational data

needs a large dataset



**DIRECT SHORTWAVE RF FOR DIFFERENT AEROSOL TYPES  
(DD, UI-BB, MA) OBSERVED IN THE PERIOD 2004-2007**



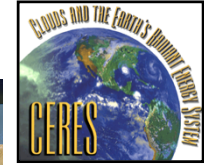
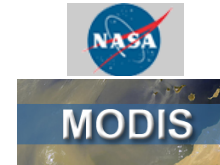
# DIRECT SHORTWAVE RF FOR DIFFERENT AEROSOL TYPES (DD, UI-BB, MA) OBSERVED IN THE PERIOD 2004-2007



$FE_S$

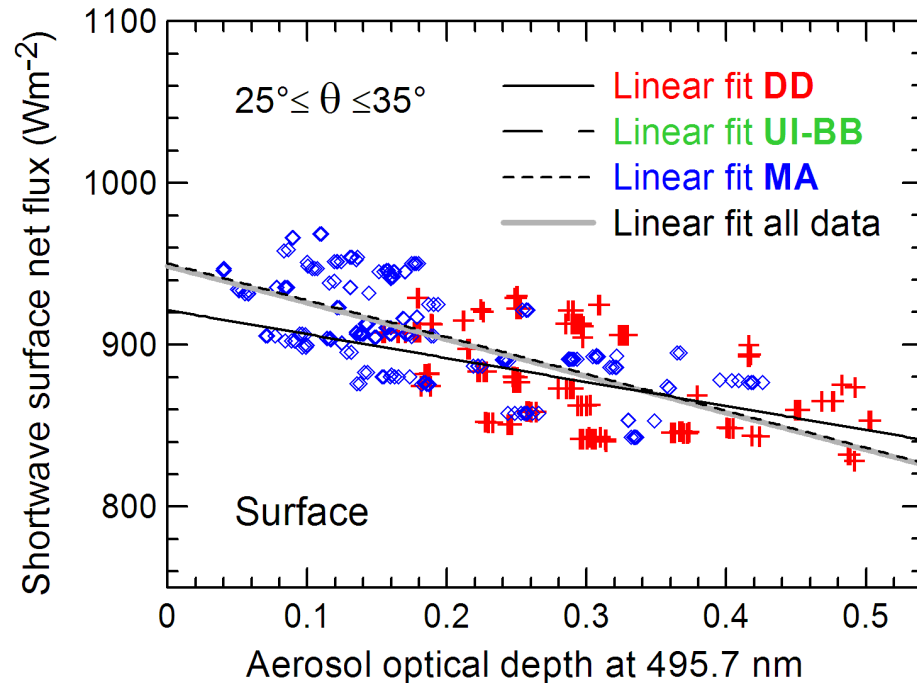
$FE_{TOA}$

$FE_{ATM}$

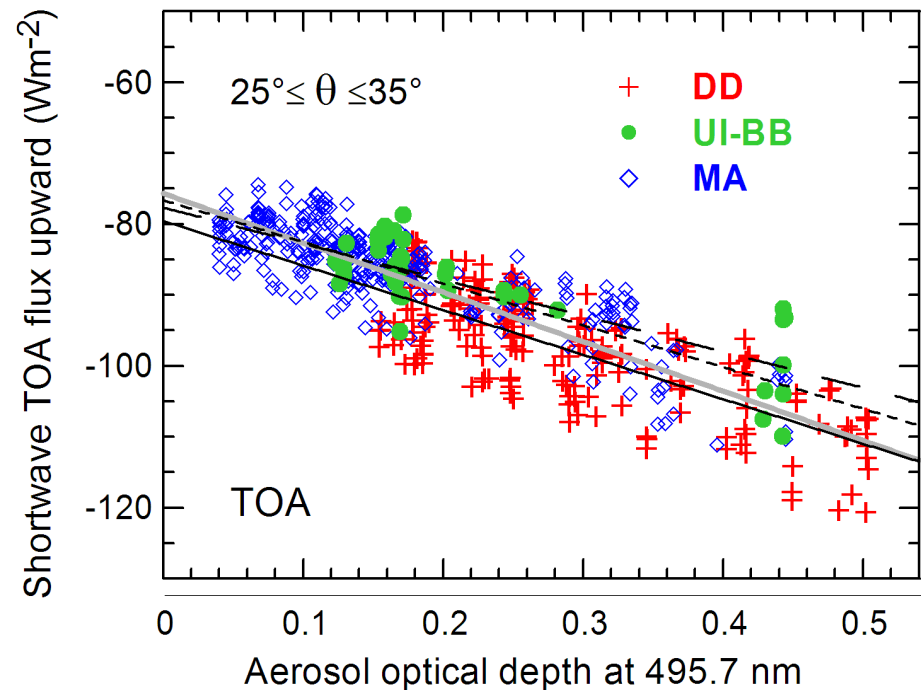


- surface observations (aerosol optical properties, surface sw fluxes)
- satellite observations (surface sw albedo, TOA sw outgoing fluxes)

SURFACE



TOA (top of atmosphere)



*Di Biagio et al., JGR 2010*

FE at different  $\theta$  → integrated to obtain the daily forcing efficiency  $(FE)_d$



The **daily mean forcing efficiency** ( $FE_d$ ) is largest for DD at TOA and for MA at surface; lowest values are obtained for UI-BB. The atmospheric forcing is  $\sim 30\text{-}50\%$  of the surface forcing for DD,  $\sim 70\%$  for UI-BB, and  $\sim 60\%$  for MA.

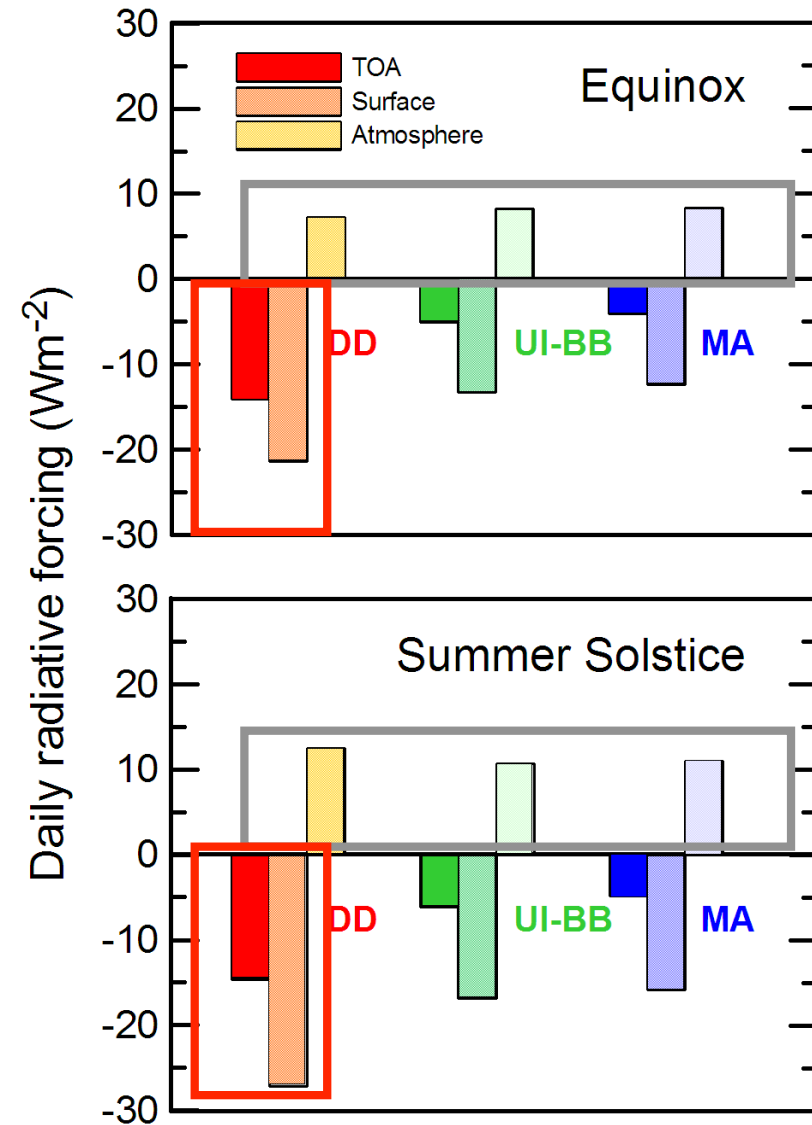
	$FE_d$ at the equinox ( $Wm^{-2}$ )		
	TOA	Surface	Atm
<b>DD</b>	-45.5 $\pm$ 5.4	-68.9 $\pm$ 4.0	23.4 $\pm$ 6.7
<b>UI-BB</b>	-19.2 $\pm$ 3.3	-59.0 $\pm$ 4.3	39.8 $\pm$ 5.4
<b>MA</b>	-36.2 $\pm$ 1.7	-94.9 $\pm$ 5.1	58.7 $\pm$ 5.4

	$FE_d$ at the summer solstice ( $Wm^{-2}$ )		
	TOA	Surface	Atm
<b>DD</b>	-47.3 $\pm$ 5.6	-87.5 $\pm$ 5.0	40.2 $\pm$ 7.5
<b>UI-BB</b>	-23.3 $\pm$ 4.1	-75.6 $\pm$ 7.9	52.3 $\pm$ 8.9
<b>MA</b>	-44.2 $\pm$ 2.1	-120.5 $\pm$ 6.5	76.3 $\pm$ 6.8

The **daily mean radiative forcing** ( $\mathbf{RF}_d$ ) at TOA and at the surface are largest for DD due to the high value of both  $FE_d$  and average  $\tau$ .

The atmospheric  $RF_d$ , conversely, is approximately independent of the aerosol type.

Aerosol type	$\tau$ (495.7 nm)
<b>DD</b>	$0.31 \pm 0.13$
<b>UI-BB</b>	$0.21 \pm 0.09$
<b>MA</b>	$0.14 \pm 0.08$



# DEPENDENCE OF THE FE ON THE SINGLE SCATTERING ALBEDO ( $\omega$ )

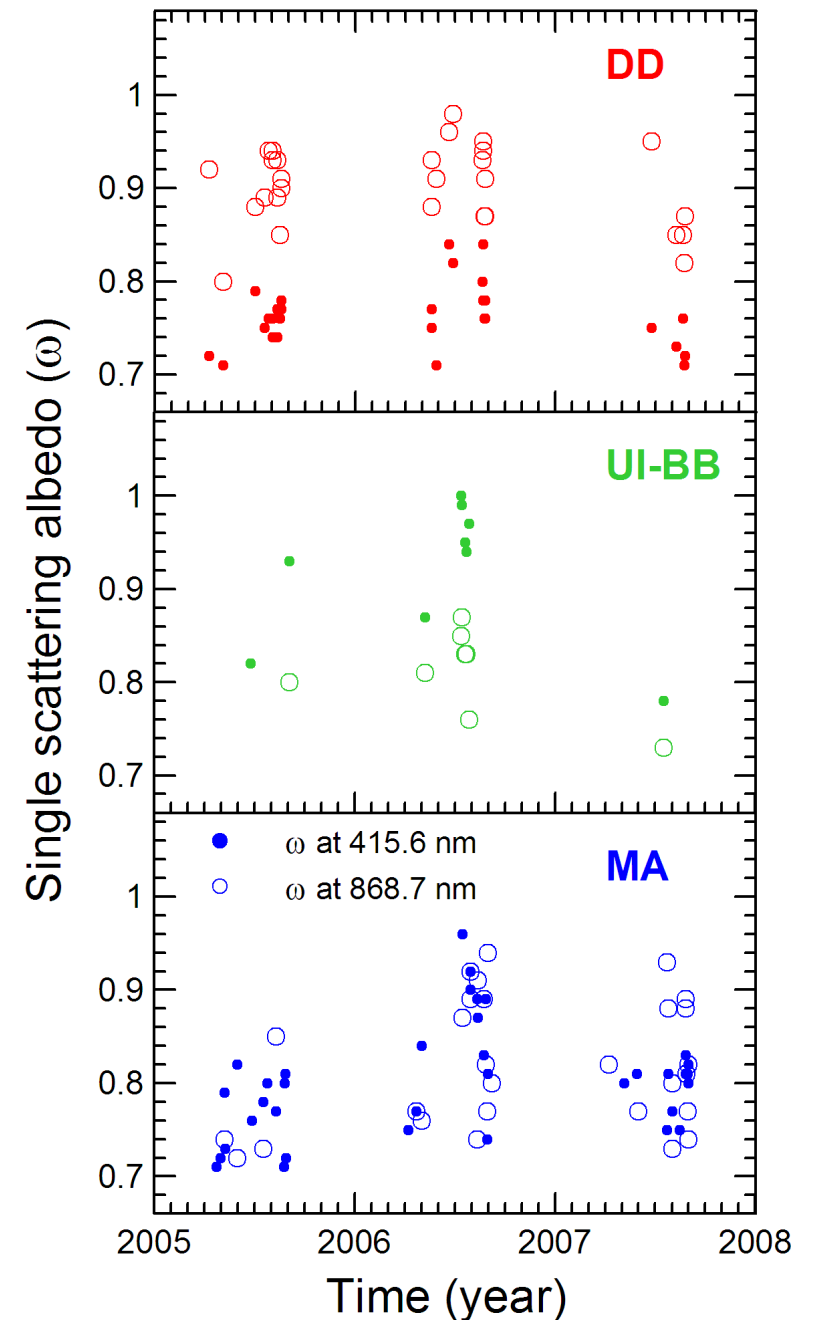
measurements (DDR +  $\tau$ )  
+  
radiative transfer model (UVSPEC)

*Meloni et al., ACP 2006*

Single scattering albedo:  $0 \leq \omega \leq 1$

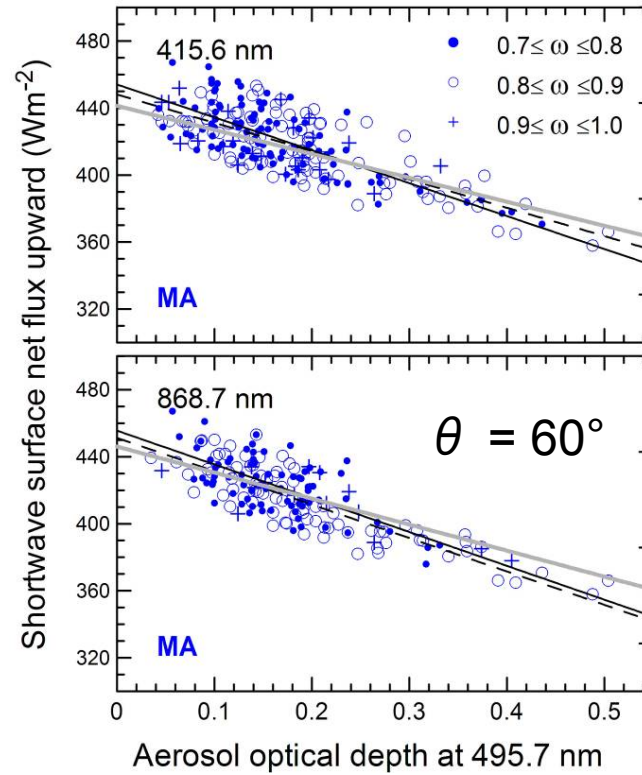
2004 - 2007

aerosol type	$\omega$ (415.6 nm)	$\omega$ (868.7 nm)
<b>DD</b>	$0.76 \pm 0.03$	$0.89 \pm 0.05$
<b>UI-BB</b>	$0.91 \pm 0.06$	$0.81 \pm 0.04$
<b>MA</b>	$0.81 \pm 0.06$	$0.82 \pm 0.07$

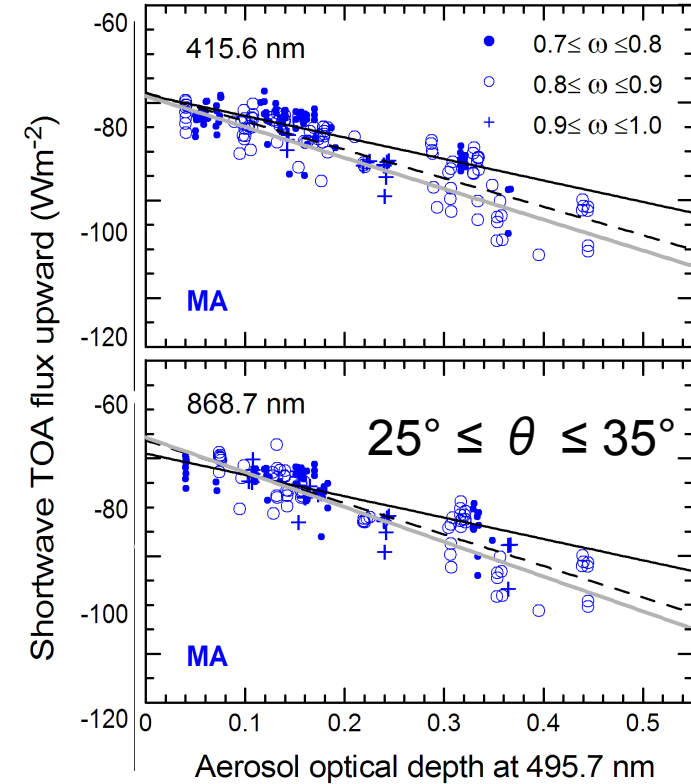




## Surface



## TOA (top of atmosphere)



FE for DD and MA

- $0.7 \leq \omega < 0.8$
- $0.8 \leq \omega < 0.9$
- $0.9 \leq \omega \leq 1.0$

$FE_{TOA}$

$\Delta\omega = +0.1 \rightarrow \Delta FE_{TOA} \approx \downarrow \sim 25\%$

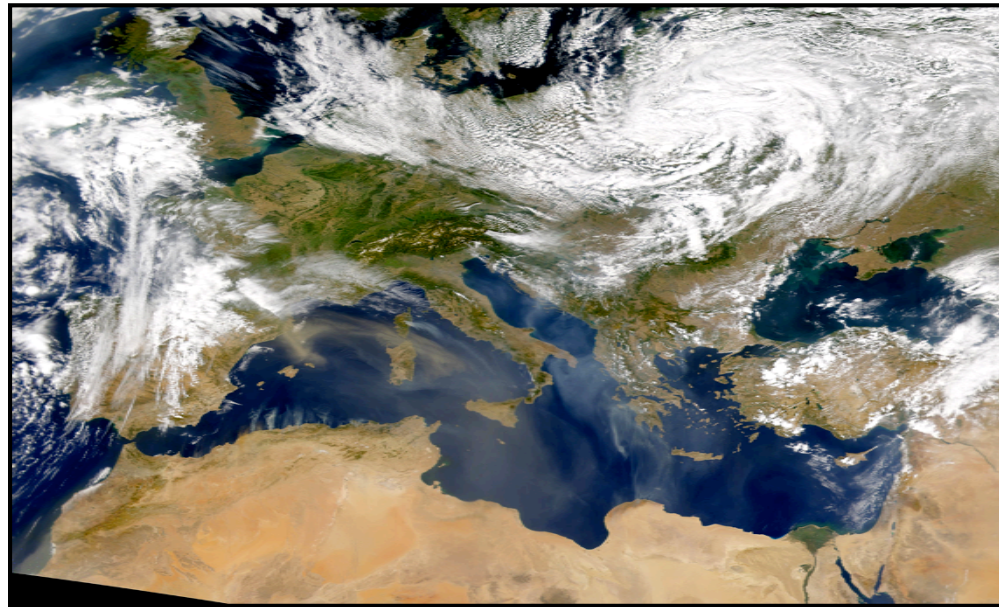
$FE_S$

$\Delta\omega = +0.1 \rightarrow \Delta FE_S \approx \uparrow \sim 15\%$

$FE_{ATM}$

$\Delta\omega = +0.1 \rightarrow \downarrow \sim 40\%$

## **THE INTENSE SAHARAN DUST EVENT OF MARCH 25-26, 2010**

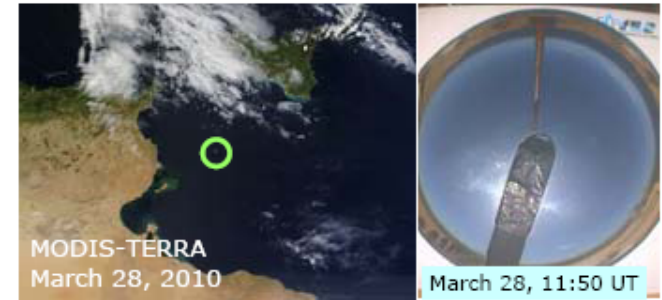
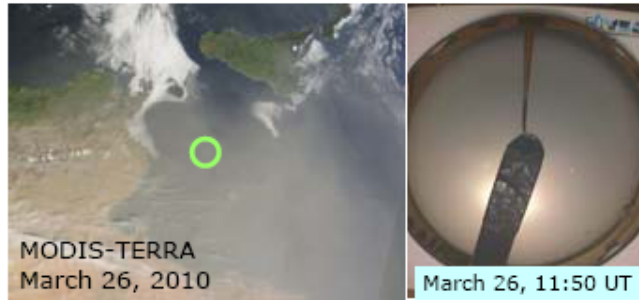
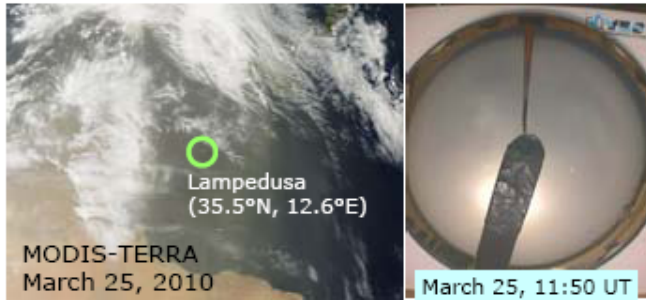


# THE INTENSE SAHARAN DUST EVENT OF MARCH 25-26, 2010

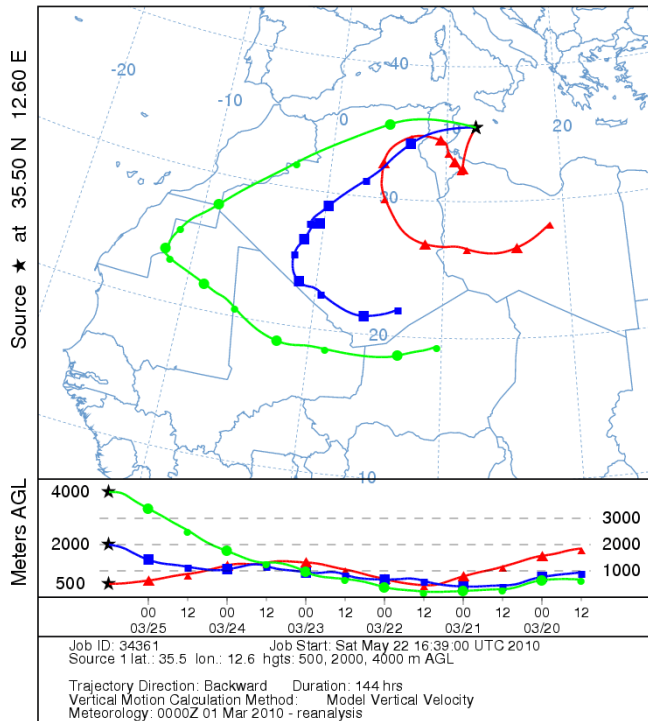
March 25

March 26

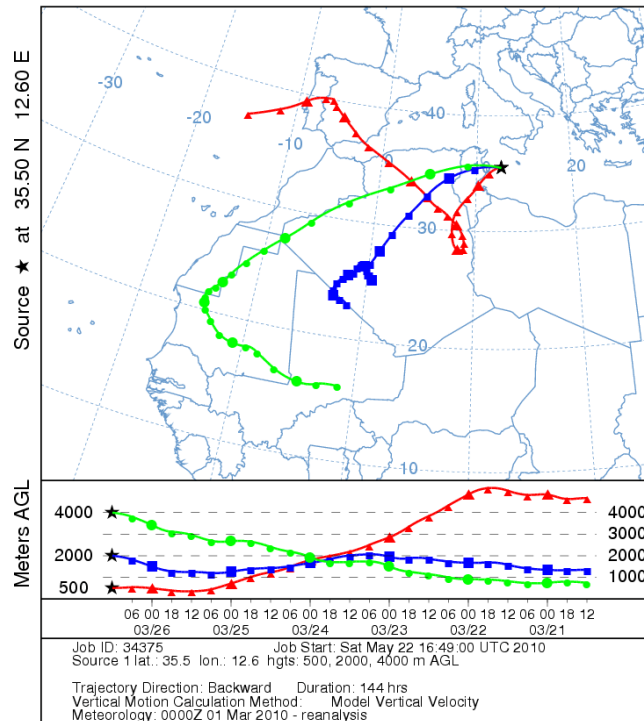
March 28



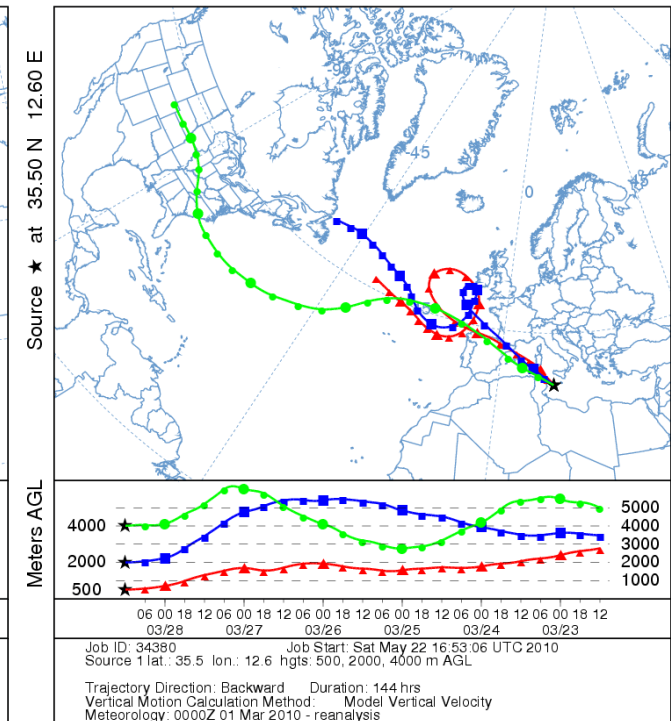
NOAA HYSPLIT MODEL  
Backward trajectories ending at 1200 UTC 25 Mar 10  
CDC1 Meteorological Data



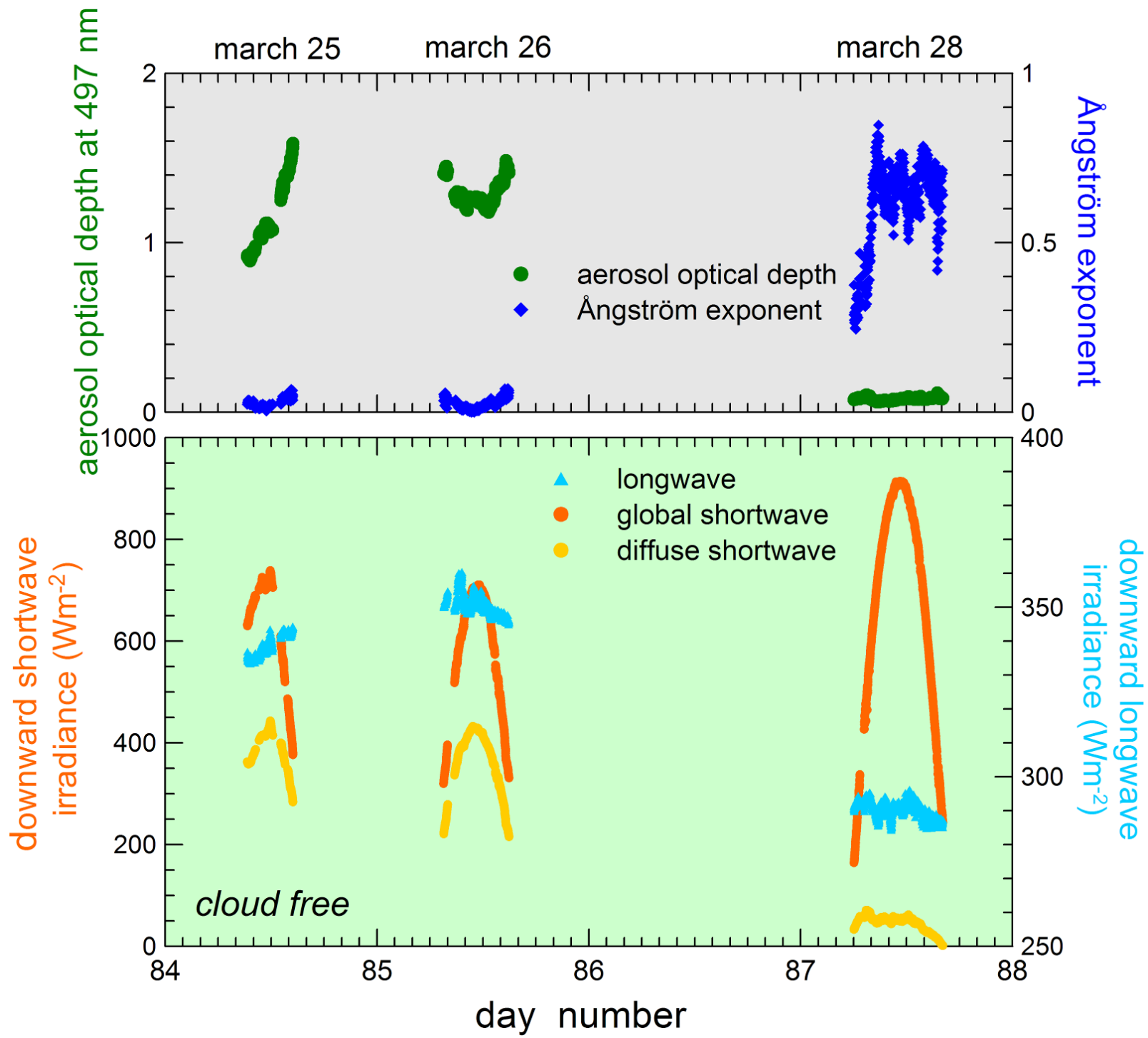
NOAA HYSPLIT MODEL  
Backward trajectories ending at 1200 UTC 26 Mar 10  
CDC1 Meteorological Data



NOAA HYSPLIT MODEL  
Backward trajectories ending at 1200 UTC 28 Mar 10  
CDC1 Meteorological Data







$$RF = F_{net} - (F_{net})_{\text{aerosol-free}}$$

March 25, 26

March 28

- surface observations
- satellite observations

no satellite data available for this day

	SHORTWAVE DAILY RF (Wm <sup>-2</sup> )		
	TOA	Surface	Atm
March 25		-158.2	
March 26	-98.7	-173.5	74.8

**7-10 times larger than daily RF at the equinox calculated for DD**

- surface observations
- satellite observations
- MODTRAN simulation

	LONGWAVE DAILY RF (Wm <sup>-2</sup> )		
	TOA	Surface	Atm
March 25	12.4	32.7	-20.3
March 26	15.5	42.7	-27.2

→ it is the first time we have such a complete characterization of both sw and lw aerosol forcing for a very intense event, but.....

*....it is still work in progress!*



Thank you



# References

Di Biagio, C., A. di Sarra, D. Meloni, F. Monteleone, S. Piacentino, and D. Sferlazzo (2009), Measurements of Mediterranean aerosol radiative forcing and influence of the single scattering albedo, *J. Geophys. Res.*, 114, D06211, doi:10.1029/2008JD011037.

Di Biagio, C., A. di Sarra, and D. Meloni (2010), Large atmospheric shortwave radiative forcing by Mediterranean aerosols derived from simultaneous ground-based and spaceborne observations and dependence on the aerosol type and single scattering albedo, *J. Geophys. Res.*, 115, D10209, doi: 10.1029/2009JD012697.

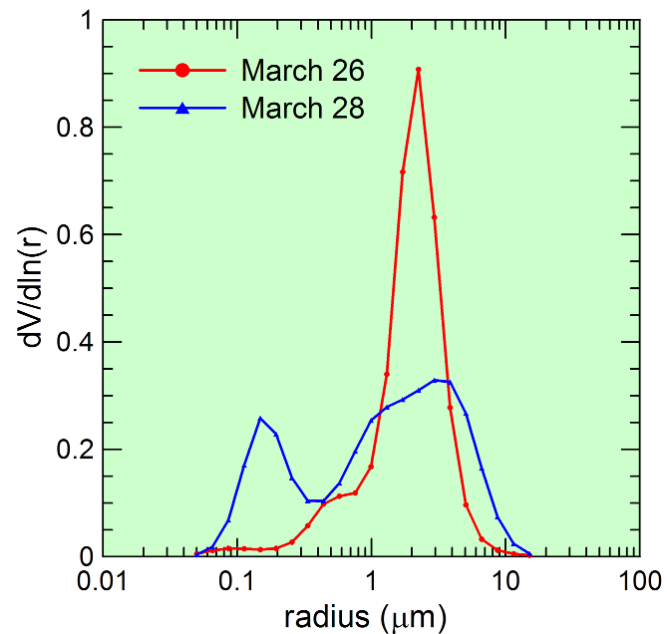
Forster, P., et al. (2007), Changes in atmospheric constituents and in radiative forcing, in *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by S. Solomon et al., pp. 129– 234, Cambridge Univ. Press, Cambridge, U. K.

Meloni, D., A. di Sarra, G. Pace, and F. Monteleone (2006), Aerosol optical properties at Lampedusa (central Mediterranean). 2. Determination of single scattering albedo at two wavelengths for different aerosol types, *Atmos. Chem. Phys.*, 6, 715– 727.

Pace, G., A. di Sarra, D. Meloni, S. Piacentino, and P. Chamard (2006), Aerosol optical properties at Lampedusa (central Mediterranean). 1. Influence of transport and identification of different aerosol types, *Atmos. Chem. Phys.*, 6, 697–713.

Satheesh, S. K., and V. Ramanathan (2000), Large differences in tropical aerosol forcing at the top of the atmosphere and Earth's surface, *Nature*, 405, 60– 63, doi:10.1038/35011039.

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# SINGLE SCATTERING ALBEDO (SSA) AND ASYMMETRY FACTOR ( $g$ )

