



2356-2

Targeted Training Activity: ENSO-Monsoon in the Current and Future Climate

30 July - 10 August, 2012

Predictability and Prediction of Indian Summer Monsoon Rainfall

SHUKLA Jagadish

Center For Ocean Land Atmosphere Studies, COLA GMU Institute For Global Environment and Society, IGES 4041 Powder Mill Road Suite 302, Calverton 20705-3106 MD

ICTP, Trieste, Italy, 30 July, 2012

Predictability and Prediction of Indian Summer Monsoon Rainfall

Jagadish Shukla

Department of Atmospheric, Oceanic and Earth Sciences (AOES) George Mason University (GMU) Center for Ocean-Land-Atmosphere Studies (COLA) Institute of Global Environment and Society (IGES)

30 July 2012





Outline

1. Introduction

Monsoon Climatology ENSO-Monsoon Relationship Has ENSO-Monsoon Relationship Broken Down? Possible Influence of Indian Ocean SST Influence of Intraseasonal Variations

2. Prediction of Indian Summer Monsoon Rainfall (ISMR)

Statistical Prediction Dynamical Prediction (Model Fidelity and Predictability)

3. Summary























Annual cycle of Rainfall over India







The normal situation:

- The warmest water is not at the equator!
- There is a roughly 5°C (9°F) temperature contrast from west to east.
- Winds blow from the cooler to the warmer water, and converge on the West Pacific Warm Pool.



Mean wind vectors overlaid







JFM CMAP Precipitation



Normal

La Nina







JFM CMAP Precipitation



El Nino

Normal



























El Nino/Southern Oscillation

























PERSISTENCE OF RAINFALL IN ANOMALOUS MONSOON SEASONS

(a) ALL-INDIA RAINFALL

Heavy Rain Years							Deficient Rain Years				
Year	Total	Jun	Jul	Aug	Sep	Year	Total	Jun	Jul	Aug	Sep
1874	1.46	1.76	0.89	-0.26	0.93	1873	-1.13	-1.38	-0.27	-0.76	-0.14
1878	1.48	-0.90	0.54	2.49	1.09	1877	-2.96	-0.58	-3.21	-2.25	-0.60
1892	1.66	-0.15	1.08	1.63	1.12	1899	-2.66	0.87	-2.35	-2.58	-1.82
1893	1.23	2.14	-0.48	-0.33	1.44	1901	-1.55	-1.33	-1.41	0.40	-1.20
1894	1.42	1.46	1.01	-0.05	0.82	1904	-1.21	0.44	-0.73	-1.21	-1.18
1916	1.17	0.98	-0.44	1.20	0.86	1905	-1.62	-1.99	-0.78	-1.00	0.10
1917	1.81	1.39	-0.75	0.83	2.56	1911	-1.38	0.79	-3.29	-0.85	0.22
1933	1.47	1.06	-0.47	1.74	0.93	1918	-2.40	0.49	-3.56	-0.60	-1.72
1942	1.26	0.28	1.31	0.97	0.26	1920	-1.59	-0.57	0.47	-2.10	-1.29
1947	1.11	-1.11	0.45	1.37	1.71	1928	-1.01	-0.26	0.25	-1.18	-1.03
1956	1.56	1.22	1.83	0.27	0.24	1941	-1.48	0.02	-1.53	-0.88	-0.92
1959	1.09	-0.17	1.27	0.17	1.18	1951	-1.35	-0.23	-0.54	-0.93	-1.30
1961	2.00	0.64	1.14	0.79	1.91	1965	-1.70	-1.45	-0.20	-1.41	-0.74
1970	1.04	1.30	-1.40	1.44	0.96	1966	-1.34	0.06	-1.00	-1.18	-0.86
<u>1975</u>	1.32	0.46	0.43	0.55	1.50	1968	-1.17	-0.74	0.39	-1.27	-0.95
1983	1.23	-0.70	-0.02	1.30	2.11	1972	-2.38	-1.12	-2.46	-0.68	-1.10
1988	1.30	-0.19	1.33	0.85	0.91	1974	-1.24	-1.57	-0.02	-0.40	-0.81
1994	1.02	1.42	1.14	0.57	-0.79	1979	-1.72	-0.55	-1.32	-1.16	-0.83
-						1982	-1.40	-0.93	-1.58	0.65	-1.32
El Nino						1985	-1.10	-0.63	-0.57	-0.65	-0.62
1							-1.30	0.38	-0.98	-0.78	-1.52
La Nina 1987							-1.85	-1.30	-1.84	-0.17	-0.88
							-		1.7-11-11-11-11-11-11-11-11-11-11-11-11-11		the second second

 \hat{e}

In either heavy or deficient rainfall years (> 1sd), the monthly values tend to be either enhanced or depressed throughout the monsoon season

1



















ENSO and Monsoon rainfall over India





Time lagged correlation between all-India JJAS Monsoon Rainfall and the NINO3 index during the period 1960-2005. The red hatching indicates the JJAS period, the horizontal red dashed line indicates zero, and the grey shading indicates the 95%

confidence interval for the time lagged correlation.



Center of Ocean-Land-Atmosphere studies





Has ENSO-Monsoon relationship broken down?

No; it is sampling variability.





JJAS NINO3 (°C) wrt 1871-1970 Mean











21-year running cc: obs. all India JJAS rainfall and July NINO3 (top left);

Simulated pairs of bi-variate random variables with population correlation -0.54 (top right, bottom left)

$$y = \rho x + w \sqrt{1 - \rho^2}$$

where x and w are independent, normally distributed random variables with unit variance, and ρ is a constant (less than or = 1).

Population correlation between random pairs of variables (*x* and *y*) produced by this model, will be equal to ρ .









Running 21-year correlations between observed ISMR and observed July NINO3 (thick black), between ensemble mean ISMR and ensemble mean JJAS NINO3 (colored solid lines), and between ISMR and JJAS NINO3 of a selected ensemble member from each model (dots connected by solid lines). All-India rainfall in dynamical models is defined as the total land precipitation within 70E – 90E and 10N – 25N. The center year of the 21-year period is plotted on the horizontal axis. The 5% significance threshold for a 21-year

correlation is shown as a thin horizontal black line.



Center of Ocean-Land-Atmosphere studies







Possible Influence of Indian Ocean SST Anomalies on Predictability of Summer Monsoon Rainfall










JJAS DMI (°C) wrt 1871-1970 Mean











The Atmospheric Influence of Tropical Diabatic Heating Associated with Developing ENSO on Indian Monsoon

> Youkyoung Jang (Ph D Thesis) (George Mason University)

AGCM is run with climatological SST with, and without a prescribed diabatic heating/cooling anomaly over **West Pacific** and **Indian Ocean** derived from observed ENSO events.









Influence of Intraseasonal Variations



Fig.9.7(h) Synoptic charts 0300 GMT 10 July 1968.

.

FIVE HIGHEST MONSOON RAINFALL

Active and Break Composites of Rainfall and Depressions

(a) Active composite and (b) break composite of daily rainfall (mm day⁻¹) over India, and (c) active composite and (d) break composite of low pressure systems (depressions only) during JJAS 1901-1970. The active (break) phase is defined when the daily all-India average rainfall is above (below) a threshold of one half of the standard deviation of all-India average rainfall for at least five consecutive days. The composites are averages over all active/break days during 1901-70.

Statistical Prediction of ISMR

What is the skill of historical (operational) forecasts issued in April/May for ISMR?

"Much has been done since 1880... but much more needs to be done."

> Sir Gilbert Thomas Walker (1868-1958)

History of Forecasting Indian Summer Monsoon Rainfall (ISMR)

1877: Major ENSO; ISMR; highly deficient; great famine

1886: First official forecast (Blanford: high Spring snow cover - low ISMR)

1899: Another major ENSO; another great famine

1904: Gilbert Walker joined IMD (DGO: 1904-1924) Regression Equations to predict ISMR Discovered Southern Oscillation & Northern Oscillations

1924-2011: IMD: Regression Equations (16,8,6 predictors)

2009: Artificial skill; too many predictors; Data fishing (DelSole)

2012: IMD: Regression Equations + Dynamical Models

Possible causes for failure of skillful empirical forecasts:

1. Questionable procedures for selecting predictors (fishing!)

2. Too many predictors (over-fitting)

DelSole and Shukla (2012)

•Canonical Correlation Analysis (CCA) for 1880-1959 for May SST & Indian Subdivisional rainfall. (No Skill)

•CCA: May SST and all India rainfall (1880-1959): 3 SST patterns; significant skill. In-sample CC = 0.54. (This is the highest possible CC using antecedent SST.)

•No skill in independent sample for 1960-2005. No skill even in the In-Sample CCA for 1960-2005 (Statistical prediction has no future!)

CCA: NO Skill

EOFs of subdivisional JJAS rainfall (1880-2010)

Correlation Between ISMR and RANDOMIZED MAY SST

Correlation Between ISMR and RANDOMIZED MAY SST

Correlation between JJAS all-Indian Monsoon Rainfall and May tropical sea surface temperature during the period 1880-1959 (top), and analogous correlation maps derived from randomized May SST fields (middle and bottom

panels). The same color scale is used in each panel. Insignificant correlations have been masked out.

DelSole and Shukla (2012)

•Canonical Correlation Analysis (CCA) for 1880-1959 for May SST & Indian Subdivisional rainfall. (No Skill)

•CCA: May SST and all India rainfall (1880-1959): 3 SST patterns; significant skill. In-sample CC = 0.54. (This is the highest possible CC using antecedent SST.)

•No skill in independent sample for 1960-2005. No skill even in the In-Sample CCA for 1960-2005 (Statistical prediction has no future!)

Prospects for Dynamical Prediction of JJAS Mean Monsoon Rainfall

- Model Fidelity and Predictability
- Hindcasts using CFS and ENSEMBLES Models

Hypothesis

Models that simulate climatology "better" make better predictions.

Definition: Fidelity refers to the degree to which the climatology of the forecasts (including the mean and variance) matches the observed climatology

Climate Model Fidelity and Predictability

Relative Entropy: The relative entropy between two distributions, $p_1(x)$ and $p_2(x)$, is defined as

$$R(p_1, p_2) = \int_{\mathbb{R}^M} p_1 \log\left(\frac{p_1}{p_2}\right) dx$$
(1)

where the integral is a multiple integral over the range of the M-dimensional vector x.

$$R(p_1, p_2) = \frac{1}{2} \log \left(\frac{|\Sigma_2|}{|\Sigma_1|} \right) + \frac{1}{2} Tr \left\{ \Sigma_1 \left(\Sigma_2^{-1} - \Sigma_1^{-1} \right) \right\} + \sum_{k=1}^4 \frac{1}{2} \left(\mu_1^k - \mu_2^k \right)^T \Sigma_1^{-1} \left(\mu_1^k - \mu_2^k \right)$$
(2)

where μ_j^k is the mean of $p_j(x)$ in the *k*th season, representing the annual cycle, Σj is the covariance matrix of $p_j(x)$, assumed independent of season and based on seasonal anomalies. The distribution of observed temperature is appropriately identified with p_1 , and the distribution of model simulated temperature with p_2 .

Fidelity vs. Skill

Fidelity vs. Skill DEMETER 1980-2001 Seasonal Forecasts

7 models, 4 initial conditions

Lead Time = 0 months

Fidelity and Skill are related.

Models with poor climatology tend to have poor skill.

Models with better climatology tend to have better skill.

Courtesy of Tim DelSole

Fundamental barriers to advancing weather and climate diagnosis and prediction on timescales from days to years are (partly) (almost entirely?) attributable to gaps in knowledge and the limited capability of contemporary operational and research numerical prediction systems to represent precipitating convection and its multi-scale organization, particularly in the tropics.

(Moncrieff, Shapiro, Slingo, Molteni, 2007)

Low skills in current dynamical seasonal predictions are not due to intrinsic limits of predictability, but due to large errors in ICs and models. (a la NWP: large dedicated effort; slow & steady progress)

Dynamical Seasonal Prediction (DSP)

Source of predictability: Dynamical memory of atmos. IC + Boundary forcing (SST, SW, snow, sea ice)

DSP = NWP + IC of Ocean, Land, Atmosphere

- dynamically coupled and consistent IC
- Global ocean (especially upper ocean); sea ice (volume)
- Global Atmos. including stratosphere (IC)
- Global GHG (especially CO₂, O₃)
- Global land (soil moisture, vegetation, snow depth) IC

Tier 1: Fully coupled models (CGCM) to predict Boundary Forcing Tier 2: Predict Boundary Forcing separately; use AGCM

•(NWP=Atmos. IC + SST IC)

Total JJAS Precipitation (mm)

Box- and whisker- plot of JJAS all India rainfall (mm) for 1960-2005: Observed (IITM) and hindcasts from ENSEMBLES project. Centerline is the median, first and last quartile as the ends of the rectangle, and data not included in the whiskers are open circles.

Analysis of Variance: F as a measure of predictability 5 CGCMs, 46 years, 9 ensembles

Measure of predictability is

$$F = E \frac{\hat{\sigma}_S^2}{\hat{\sigma}_N^2}$$

where

$$\hat{\sigma}_{S}^{2} = \frac{1}{Y - 1} \sum_{y=1}^{Y} \left(P_{y,e} - \overline{\overline{P}} \right)^{2}$$
$$\hat{\sigma}_{N}^{2} = \frac{1}{Y(E - 1)} \sum_{y=1}^{Y} \sum_{e=1}^{E} \left(P_{y,e} - \overline{P}_{y} \right)^{2}$$
$$\overline{P}_{y} = \frac{1}{E} \sum_{e=1}^{E} P_{y,e}$$
$$\overline{\overline{P}} = \frac{1}{Y} \sum_{y=1}^{Y} \overline{P}_{y}$$

For samples drawn independently from the same normal distribution, and for Y = 46 and E = 9, the 5% significance threshold of F is 1.40

F for JJAS Precip in IFM-GEOMAR

F for JJAS Precip in UK Met Office

F for JJAS Precip in Meteo-France

F-values for JJAS precip. For 46-years and 9 ensemble members the 5% significance is F=1.4. Gray color indicates not statistically significant at 95% confidence interval.

F for JJAS Precip in CMCC–Bologna

F values for JJAS All–India Rainfall from ENSEMBLES

(46 years (1960-2005); Ens.=9)

Correlation between observed and predicted JJAS all-India rainfall for hindcasts in the ENSEMBLES data set for the period 1960-2005. All-India rainfall in dynamical models is defined as the total land precipitation within 70E – 90E and 10N – 25N.









Correlation between observed NINO3, and ensemble mean NINO3 predicted by the ENSEMBLES models, for hindcasts in the period 1960-2005, as a function of calendar month. Also shown is the correlation between observed NINO3 and the least squares prediction of NINO3 based on the observed May NINO3 value (thick grey). The 'x'-symbols on the far right give the correlations

between the observed and predicted JJAS NINO3 index.



Center of Ocean-Land-Atmosphere studies







month of NINO3

Correlation between observed JJAS all-India rainfall, and ensemble mean model predicted NINO3 index, for hindcasts from the ENSEMBLES data set during 1960-2005. The 'x'-symbols on the far right give the correlations for the JJAS NINO3 index in each model. The correlation with the mean of all models is shown as

"multimodel."



Center of Ocean-Land-Atmosphere studies







For 1960-2005 Obs, CC (April Nino3, ISMR): -0.18 CC (May Nino3, ISMR): -0.21

(46 years (1960-2005); Ens.=9)

Correlation between observed and predicted JJAS all-India rainfall for hindcasts in the ENSEMBLES data set for the period 1960-2005. All-India rainfall in dynamical models is defined as the total land

precipitation within 70E – 90E and 10N – 25N. Last row shows empirical prediction using observed May NINO3.



Center of Ocean-Land-Atmosphere studies





Summary

•Canonical Correlation Analysis (CCA) for 1880-1959 for May SST & Indian Subdivisional rainfall. (No Skill)

•CCA: May SST and all India rainfall (1880-1959): 3 SST patterns; significant skill. In-sample CC = 0.54. (This is the highest possible CC using antecedent SST.)

•No skill in independent sample for 1960-2005. No skill even in the In-Sample CCA for 1960-2005 (Statistical prediction has no future!)





Summary

• Model's ability to simulate SST and Q in West Pacific and Indian Ocean are critical for accurate monsoon prediction.

• Analysis of Variance (F test) calculation for 5 coupled model ("ENSEMBLES" Project) seasonal predictions for 46 years, 9 member ensembles show high predictability ISMR, but skill of hindcast for 1960-2005 is rather modest.

• Coupled O-A models for monsoon prediction is the future.





THANK YOU!

ANY QUESTIONS?





Statements & Conjectures

- The apparent breakdown of ENSO Monsoon correlation could be due, in large part, to sampling variability.
- Realistic Land ICs enhance weekly-monthly predictions (high resolution land rainfall data required)
- Model predictability depends on model's fidelity to simulate climate.
- Intraseasonal variations appear to be due, in part, to coupled convection-dynamics interactions. Realistic simulation of diabatic heat sources in West-Pac. & IO will be required for ISO prediction.





Statements & Conjectures

- ENSO prediction skill is sensitive to ocean IC (NCEP, ECMWF).
- AGCM forced with SST predicted by coupled O-A models have comparable skill, but larger variance.

(coupling damps the heat flux variability)

• Low skills in current dynamical seasonal predictions are not due to intrinsic limits of predictability, but due to large errors in ICs and models.

(a la NWP: large dedicated effort; slow & steady progress)







Fig. 16

: 53 :



Correlation between JJAS All-India Rainfall and SST indices

	1880-1959		1960-2005			
SST Predictor	April	May	JJAS	April	May	JJAS
canonical predictor trained in 1960-2005	-	_	-	NA	NA	0.39^{*}
canonical predictor trained in 1880-1959	0.47^{*}	0.54	0.66	0.10	0.15	0.34
NINO3	-0.21	-0.41	-0.60	-0.18	-0.21	-0.45
canonical tendency predictor trained in 1960-2005 (w/ Jan)	-	_	-	0.46^{*}	0.54^*	0.6
NINO3 tendency (w/ Jan)	0.24	0.43	0.60	0.44	0.38	0.51





Trends of annual-mean Surface Air Temperatures and Precipitation over 1951-1999



Observed Trends

Multi-model ensemble-mean trends in 76 **COUPLED** climate model simulations with prescribed observed radiative forcings

Multi-model ensemble-mean trends in 87 **UNCOUPLED** atmospheric GCM simulations with prescribed observed global or tropical SSTs, but no explicitly specified radiative forcings.

From Shin and Sardeshmukh Climate Dynamics 2011

Page 1









Seamless Prediction of Weather and Climate

From Cyclone Resolving Global Models to Cloud System Resolving Global Models

- 1. Planetary Scale Resolving Models (1970~): Δx~500Km
- 2. Cyclone Resolving Models (1980~): Δx ~100-300Km
- 3. Mesoscale Resolving Models (1990~): Δx~10-30Km
- 4. Cloud System Resolving Models (2000 ~): Δx~3-5Km



Towards a Hypothetical "Perfect" Model

- Replicate the statistical properties of the past observed climate
 - Means, variances, covariances, and patterns of covariability
- Utilize this model to estimate the limits of predicting the sequential evolution of climate variability
- Better model → Better prediction (??)









ENSO, EQWIN, and Monsoon rainfall over India

















All–India JJAS Rainfall vs. MAY NINO3 1880–1982 (blue), 1983–2010 (red), CC= –0.34











$$p(y_{t+\tau}|x_t) = p(y_{t+\tau})$$

forecast distribution climatological distribution







Record performance of the deterministic forecasting system. The useful range of the deterministic forecasts for Europe reached its highest ever monthly value in February 2010. Overall the performance has been consistently good during 2010. The useful forecast range is determined by the time at which the anomaly correlation for 500 hPa height operational forecasts at 12 UTC reached 60%.













The Tale of Two Monsoon Seasons

All India JJAS Monsoon Rainfall as % of Normal

	Actual	Forecast (IMD)
1994	+13 %	-8 %
1997	+2 %	-8 %





















