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Earthquake Tectonics and Hazards on the Continents

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GPS and applications to the earthquake cycle, and use in particular earthquakes

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Elastic Dislocation Modelling



Y. Okada, 1985. Surface deformation due to **shear** and tensile faults in a half-space. *Bull. Seism. Soc. Am.*, 75, 1135-1154

To define a rectangular fault dislocation, need 10 parameters:

- Location of fault x,y,z (x=y=0, z = -d) in fault-centred reference frame with the x-axis pointing along strike [1]
- Length, Width and dip of the fault (L, W, δ) [3]
- Slip components (u_1 = strike-slip; u_2 = dip-slip; u_3 = tensile) [3]
- Shift and rotate 3D Displacements calculated in the fault-centred reference frame to fit location and strike of model fault. [3]









<figure><figure>

Interferogram for the Bam Earthquake

The first major earthquake for which a radar interferogram was constructed using Envisat ASAR data.

Elastic dislocation modelling of the interferogram showed that earthquake occurred on an unknown fault with no obvious surface features and not on any of the mapped faults.





Interferometric Coherence for the Bam Earthquake

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The line of incoherence – disturbed ground – south of Bam was used to lead field workers to the right area.









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Elastic dislocation modelling of the interferogram showed that earthquake occurred on an unknown fault with no obvious surface features and not on the mapped fault.

This fault dips in towards the fault on which the main rupture occurred and may interact with it. A small amount of slip may have occurred on the mapped fault near it intersection with the main fault.









The earthquake occurred at 23:49 on 13 April, 2010 (UTC; 07:49, 14 April 2010, local)

Approximately 70% of houses collapsed. ~2,700 people were killed and over 12,000 injured

Gyegu Monastery













Earthquake occurred on the Yushu-Garzê-Xianshuihe Fault System Surface rupture traced for about 35 km with a peak slip of 1.8m

Estimating the Length of the Fault from the Seismic Moment

$$M_0 = \mu L W \overline{u}$$

Assume:

 μ = 3.2×10¹⁰ Pa, W = 15 km, u = 1.5 m, M₀ = 2.5 ×10¹⁹ N m (GCMT), then L = 35 km







What are the Advantages of InSAR over Seismology in Studying Earthquakes?

- Satellite radar interferometry can be used to study all stages of the earthquake process – interseismic strain accumulation, the earthquake itself, and transient postseismic adjustments.
- Modelling allows the factors of the amount of slip and the length and width of fault, that determine the magnitude of the earthquake, to be determined separately.
- Perhaps the most important, often unmentioned, advantage is that InSAR provides an accurate reference frame within which the surface deformation due to earthquakes is located, and hence accurate earthquake fault locations.









































