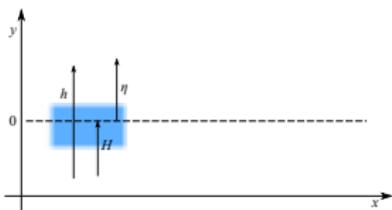


One-layer shallow water system

$$\begin{cases} \frac{\partial h}{\partial t} + \frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = 0 \\ \frac{\partial q_x}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x^2}{h} + \frac{g}{2} h^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_x q_y}{h} \right) = gh \frac{\partial H}{\partial x} + S_{f_1}(w) \\ \frac{\partial q_y}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x q_y}{h} \right) + \frac{\partial}{\partial y} \left(\frac{q_y^2}{h} + \frac{g}{2} h^2 \right) = gh \frac{\partial H}{\partial y} + S_{f_2}(w), \end{cases}$$



- $w = (h, q_x, q_y)^T$,
- h is the water depth,
- $\mathbf{q} = (q_x, q_y)$, is the mass flow,
- $\mathbf{u} = (u_x, u_y) = \frac{\mathbf{q}}{h}$ is the mean velocity,
- H is the bathymetry,
- g is the acceleration of gravity,
- $S_{f_i}(w)$, $i = 1, 2$, parametrize the friction terms.

One-layer shallow water system

$$\left\{ \begin{array}{l} \frac{\partial h}{\partial t} + \frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = 0 \\ \frac{\partial q_x}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x^2}{h} + \frac{g}{2} h^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_x q_y}{h} \right) = gh \frac{\partial H}{\partial x} + S_{f_1}(w) \\ \frac{\partial q_y}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x q_y}{h} \right) + \frac{\partial}{\partial y} \left(\frac{q_y^2}{h} + \frac{g}{2} h^2 \right) = gh \frac{\partial H}{\partial y} + S_{f_2}(w), \end{array} \right.$$

- Quadratic:

$$S_{f_1}(w) = -cu_x \|\mathbf{u}\| \quad S_{f_2}(w) = -cu_y \|\mathbf{u}\|;$$

- Chèzy:

$$S_{f_1}(w) = -\frac{g}{C^2} u_x \|\mathbf{u}\| \quad S_{f_2}(w) = -\frac{g}{C^2} u_y \|\mathbf{u}\|;$$

- Manning:

$$S_{f_1}(w) = -gh \frac{n^2}{h^{4/3}} u_x \|\mathbf{u}\| \quad S_{f_2}(w) = -gh \frac{n^2}{h^{4/3}} u_y \|\mathbf{u}\|.$$

One-layer shallow water system

$$\left\{ \begin{array}{l} \frac{\partial h}{\partial t} + \frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = 0 \\ \frac{\partial q_x}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x^2}{h} + \frac{g}{2} h^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_x q_y}{h} \right) = gh \frac{\partial H}{\partial x} + S_{f_1}(w) \\ \frac{\partial q_y}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x q_y}{h} \right) + \frac{\partial}{\partial y} \left(\frac{q_y^2}{h} + \frac{g}{2} h^2 \right) = gh \frac{\partial H}{\partial y} + S_{f_2}(w), \end{array} \right.$$

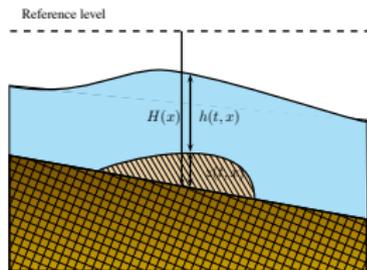


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Discrete and continuous dynamical systems- Series B 1 (1):89-102, 2001.

Shallow water-Exner equations

$$\left\{ \begin{array}{l} \frac{\partial h}{\partial t} + \frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = 0 \\ \frac{\partial q_x}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x^2}{h} + \frac{g}{2} h^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_x q_y}{h} \right) = -gh \frac{\partial z_b}{\partial x} + gh \frac{\partial \tilde{H}}{\partial x} + S_{f1}(w) \\ \frac{\partial q_y}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x q_y}{h} \right) + \frac{\partial}{\partial y} \left(\frac{q_y^2}{h} + \frac{g}{2} h^2 \right) = -gh \frac{\partial z_b}{\partial y} + gh \frac{\partial \tilde{H}}{\partial y} + S_{f2}(w), \\ \frac{\partial z_b}{\partial t} + \xi \left(\frac{\partial q_{b,x}}{\partial x}(w) + \frac{\partial q_{b,y}}{\partial y}(w) \right) = 0 \end{array} \right.$$



- $w = (h, q_x, q_y, z_b)^T$;
- h is the water depth, z_b the sediment layer depth;
- $q = (q_x, q_y)$, is the mass flow;
- \tilde{H} is the non-erodible bathymetry;
- $\xi = 1/(1 - \rho_0)$ and ρ_0 is the porosity of the sediment;
- $(q_{b,x}(w), q_{b,y}(w))$ the solid transport discharge.

Shallow water-Exner equations

$$\left\{ \begin{array}{l} \frac{\partial h}{\partial t} + \frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = 0 \\ \frac{\partial q_x}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x^2}{h} + \frac{g}{2} h^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_x q_y}{h} \right) = -gh \frac{\partial z_b}{\partial x} + gh \frac{\partial \tilde{H}}{\partial x} + S_{f_1}(w) \\ \frac{\partial q_y}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x q_y}{h} \right) + \frac{\partial}{\partial y} \left(\frac{q_y^2}{h} + \frac{g}{2} h^2 \right) = -gh \frac{\partial z_b}{\partial y} + gh \frac{\partial \tilde{H}}{\partial y} + S_{f_2}(w), \\ \frac{\partial z_b}{\partial t} + \xi \left(\frac{\partial q_{b,x}}{\partial x}(w) + \frac{\partial q_{b,y}}{\partial y}(w) \right) = 0 \end{array} \right.$$

- Fowler-Grass:

$$q_{b,x}(w) = A_g \frac{z_b}{\bar{z}_0} u_x \|\mathbf{u}\|^2 \quad q_{b,y}(w) = A_g \frac{z_b}{\bar{z}_0} u_y \|\mathbf{u}\|^2,$$

where \bar{z}_0 represents the mean value of the thickness of the sediment layer.

Shallow water-Exner equations

$$\left\{ \begin{array}{l} \frac{\partial h}{\partial t} + \frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = 0 \\ \frac{\partial q_x}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x^2}{h} + \frac{g}{2} h^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_x q_y}{h} \right) = -gh \frac{\partial z_b}{\partial x} + gh \frac{\partial \tilde{H}}{\partial x} + S_{f_1}(w) \\ \frac{\partial q_y}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_x q_y}{h} \right) + \frac{\partial}{\partial y} \left(\frac{q_y^2}{h} + \frac{g}{2} h^2 \right) = -gh \frac{\partial z_b}{\partial y} + gh \frac{\partial \tilde{H}}{\partial y} + S_{f_2}(w), \\ \frac{\partial z_b}{\partial t} + \xi \left(\frac{\partial q_{b,x}}{\partial x}(w) + \frac{\partial q_{b,y}}{\partial y}(w) \right) = 0 \end{array} \right.$$



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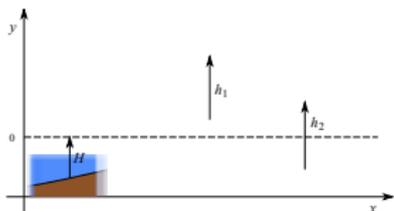


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Two-layer shallow-water model

$$\left\{ \begin{array}{l} \frac{\partial h_1}{\partial t} + \frac{\partial q_{1,x}}{\partial x} + \frac{\partial q_{1,y}}{\partial y} = 0 \\ \frac{\partial q_{1,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x}^2}{h_1} + \frac{g}{2} h_1^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) = -g h_1 \frac{\partial h_2}{\partial x} + g h_1 \frac{\partial H}{\partial x} + S_{i_1}(w) \\ \frac{\partial q_{1,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,y}^2}{h_1} + \frac{g}{2} h_1^2 \right) = -g h_1 \frac{\partial h_2}{\partial y} + g h_1 \frac{\partial H}{\partial y} + S_{i_2}(w) \\ \frac{\partial h_2}{\partial t} + \frac{\partial q_{2,x}}{\partial x} + \frac{\partial q_{2,y}}{\partial y} = 0 \\ \frac{\partial q_{2,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x}^2}{h_2} + \frac{g}{2} h_2^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) = -g r h_2 \frac{\partial h_1}{\partial x} + g h_2 \frac{\partial H}{\partial x} + S_{i_3}(w) + S_{f_1}(w) \\ \frac{\partial q_{2,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,y}^2}{h_2} + \frac{g}{2} h_2^2 \right) = -g r h_2 \frac{\partial h_1}{\partial y} + g h_2 \frac{\partial H}{\partial y} + S_{i_4}(w) + S_{f_2}(w) \end{array} \right.$$



- $w = (h_1, q_{1,x}, q_{1,y}, h_2, q_{2,x}, q_{2,y})^T$;
- h_i , is the layer depth, $i = 1, 2$;
- $q_i = (q_{i,x}, q_{i,y})$, is the mass flow at each layer, $i = 1, 2$;
- $r = \rho_1 / \rho_2$ is the ratio of the constant densities of the layers ($\rho_1 < \rho_2$);
- $S_{i_l}(w)$, $l = 1, \dots, 4$, parametrize the friction terms between the two layers;
- $S_{f_l}(w)$, $l = 1, 2$ parametrize the friction terms between the second layer and the bottom topography.

Two-layer shallow-water model

$$\left\{ \begin{aligned} \frac{\partial h_1}{\partial t} + \frac{\partial q_{1,x}}{\partial x} + \frac{\partial q_{1,y}}{\partial y} &= 0 \\ \frac{\partial q_{1,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x}^2}{h_1} + \frac{g}{2} h_1^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) &= -gh_1 \frac{\partial h_2}{\partial x} + gh_1 \frac{\partial H}{\partial x} + S_{i_1}(w) \\ \frac{\partial q_{1,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,y}^2}{h_1} + \frac{g}{2} h_1^2 \right) &= -gh_1 \frac{\partial h_2}{\partial y} + gh_1 \frac{\partial H}{\partial y} + S_{i_2}(w) \\ \frac{\partial h_2}{\partial t} + \frac{\partial q_{2,x}}{\partial x} + \frac{\partial q_{2,y}}{\partial y} &= 0 \\ \frac{\partial q_{2,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x}^2}{h_2} + \frac{g}{2} h_2^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) &= -grh_2 \frac{\partial h_1}{\partial x} + gh_2 \frac{\partial H}{\partial x} + S_{i_3}(w) + S_{f_1}(w) \\ \frac{\partial q_{2,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,y}^2}{h_2} + \frac{g}{2} h_2^2 \right) &= -grh_2 \frac{\partial h_1}{\partial y} + gh_2 \frac{\partial H}{\partial y} + S_{i_4}(w) + S_{f_2}(w), \end{aligned} \right.$$

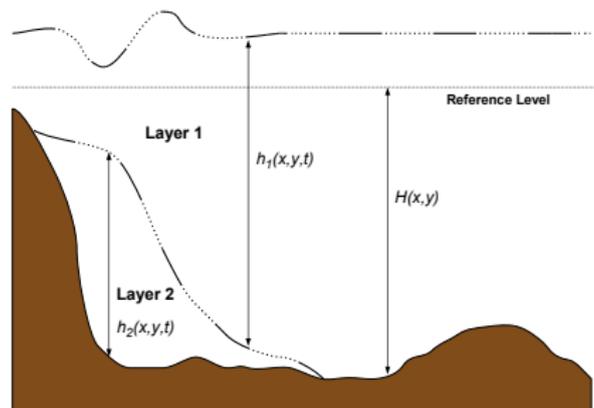
- $S_{i_1} = c(u_{2,x} - u_{1,x}) \| \mathbf{u}_1 - \mathbf{u}_2 \|$, $S_{i_2} = c(u_{2,y} - u_{1,y}) \| \mathbf{u}_1 - \mathbf{u}_2 \|$, $S_{i_3} = -rS_{i_1}$, $S_{i_4} = -rS_{i_2}$.

- $S_{i_1} = c \frac{h_1 h_2}{rh_1 + h_2} (u_{2,x} - u_{1,x}) \| \mathbf{u}_1 - \mathbf{u}_2 \|$, $S_{i_2} = c \frac{h_1 h_2}{rh_1 + h_2} (u_{2,y} - u_{1,y}) \| \mathbf{u}_1 - \mathbf{u}_2 \|$, $S_{i_3} = -rS_{i_1}$, $S_{i_4} = -rS_{i_2}$.

2D (Simplified) Two-layer Savage-Hutter shallow-water model

Hypothesis

- We consider an stratified media composed by a non viscous and homogeneous fluid with constant density ρ_1 (water) and a granular material with density ρ_s and porosity ψ_0 . We suppose the mean density of the granular material is given by:
$$\rho_2 = (1 - \psi_0)\rho_s + \psi_0\rho_1.$$
- Fluid and granular material are immiscible.



2D (Simplified) Two-layer Savage-Hutter shallow-water model

$$\left\{ \begin{aligned}
 & \frac{\partial h_1}{\partial t} + \frac{\partial q_{1,x}}{\partial x} + \frac{\partial q_{1,y}}{\partial y} = 0 \\
 & \frac{\partial q_{1,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x}^2}{h_1} + \frac{g}{2} h_1^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) = -g h_1 \frac{\partial h_2}{\partial x} + g h_1 \frac{\partial H}{\partial x} + S_{i_1}(w) \\
 & \frac{\partial q_{1,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,y}^2}{h_1} + \frac{g}{2} h_1^2 \right) = -g h_1 \frac{\partial h_2}{\partial y} + g h_1 \frac{\partial H}{\partial y} + S_{i_2}(w) \\
 & \frac{\partial h_2}{\partial t} + \frac{\partial q_{2,x}}{\partial x} + \frac{\partial q_{2,y}}{\partial y} = 0 \\
 & \frac{\partial q_{2,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x}^2}{h_2} + \frac{g}{2} h_2^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) = -g r h_2 \frac{\partial h_1}{\partial x} + g h_2 \frac{\partial H}{\partial x} + S_{i_3}(w) + \tau_x(w) \\
 & \frac{\partial q_{2,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,y}^2}{h_2} + \frac{g}{2} h_2^2 \right) = -g r h_2 \frac{\partial h_1}{\partial y} + g h_2 \frac{\partial H}{\partial y} + S_{i_4}(w) + \tau_y(w),
 \end{aligned} \right.$$

- h_i , is the layer depth, $i = 1, 2$,
- $q_i = (q_{i,x}, q_{i,y})$, is the mass flow at each layer, $i = 1, 2$,
- $w = (h_1, q_{1,x}, q_{1,y}, h_2, q_{2,x}, q_{2,y})^T$,
- $S_{i_l}(w)$, $l = 1, \dots, 4$, parametrize the friction between the layers,
- $\tau(w) = (\tau_x(w), \tau_y(w))$ parametrize the Coulomb friction term.

2D (Simplified) Two-layer Savage-Hutter shallow-water model

$$\left\{ \begin{aligned}
 & \frac{\partial h_1}{\partial t} + \frac{\partial q_{1,x}}{\partial x} + \frac{\partial q_{1,y}}{\partial y} = 0 \\
 & \frac{\partial q_{1,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x}^2}{h_1} + \frac{g}{2} h_1^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) = -g h_1 \frac{\partial h_2}{\partial x} + g h_1 \frac{\partial H}{\partial x} + S_{i_1}(w) \\
 & \frac{\partial q_{1,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,y}^2}{h_1} + \frac{g}{2} h_1^2 \right) = -g h_1 \frac{\partial h_2}{\partial y} + g h_1 \frac{\partial H}{\partial y} + S_{i_2}(w) \\
 & \frac{\partial h_2}{\partial t} + \frac{\partial q_{2,x}}{\partial x} + \frac{\partial q_{2,y}}{\partial y} = 0 \\
 & \frac{\partial q_{2,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x}^2}{h_2} + \frac{g}{2} h_2^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) = -g r h_2 \frac{\partial h_1}{\partial x} + g h_2 \frac{\partial H}{\partial x} + S_{i_3}(w) + \tau_x(w) \\
 & \frac{\partial q_{2,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,y}^2}{h_2} + \frac{g}{2} h_2^2 \right) = -g r h_2 \frac{\partial h_1}{\partial y} + g h_2 \frac{\partial H}{\partial y} + S_{i_4}(w) + \tau_y(w), \\
 \\
 & \tau(w) = \begin{cases} \left(g h_2 (1-r) \frac{u_{2,x}}{\|u\|} \tan(\delta_0), g h_2 (1-r) \frac{u_{2,y}}{\|u\|} \tan(\delta_0) \right) & \text{if } \tau > \sigma_c \\ \mathbf{u}_2 = \mathbf{0} & \text{otherwise} \end{cases}
 \end{aligned} \right.$$

2D (Simplified) Two-layer Savage-Hutter shallow-water model

$$\left\{ \begin{array}{l}
 \frac{\partial h_1}{\partial t} + \frac{\partial q_{1,x}}{\partial x} + \frac{\partial q_{1,y}}{\partial y} = 0 \\
 \frac{\partial q_{1,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x}^2}{h_1} + \frac{g}{2} h_1^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) = -g h_1 \frac{\partial h_2}{\partial x} + g h_1 \frac{\partial H}{\partial x} + S_{i_1}(w) \\
 \frac{\partial q_{1,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{1,x} q_{1,y}}{h_1} \right) + \frac{\partial}{\partial y} \left(\frac{q_{1,y}^2}{h_1} + \frac{g}{2} h_1^2 \right) = -g h_1 \frac{\partial h_2}{\partial y} + g h_1 \frac{\partial H}{\partial y} + S_{i_2}(w) \\
 \frac{\partial h_2}{\partial t} + \frac{\partial q_{2,x}}{\partial x} + \frac{\partial q_{2,y}}{\partial y} = 0 \\
 \frac{\partial q_{2,x}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x}^2}{h_2} + \frac{g}{2} h_2^2 \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) = -g r h_2 \frac{\partial h_1}{\partial x} + g h_2 \frac{\partial H}{\partial x} + S_{i_3}(w) + \tau_x(w) \\
 \frac{\partial q_{2,y}}{\partial t} + \frac{\partial}{\partial x} \left(\frac{q_{2,x} q_{2,y}}{h_2} \right) + \frac{\partial}{\partial y} \left(\frac{q_{2,y}^2}{h_2} + \frac{g}{2} h_2^2 \right) = -g r h_2 \frac{\partial h_1}{\partial y} + g h_2 \frac{\partial H}{\partial y} + S_{i_4}(w) + \tau_y(w),
 \end{array} \right.$$



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General framework

General formulation

$$\frac{\partial w}{\partial t} + \frac{\partial F_1}{\partial x}(w) + \frac{\partial F_2}{\partial y}(w) + B_1(w) \frac{\partial w}{\partial x} + B_2(w) \frac{\partial w}{\partial y} = S_1(w) \frac{\partial H}{\partial x} + S_2(w) \frac{\partial H}{\partial y} + S_F(w),$$

General framework

General formulation

$$\frac{\partial w}{\partial t} + \frac{\partial F_1}{\partial x}(w) + \frac{\partial F_2}{\partial y}(w) + B_1(w) \frac{\partial w}{\partial x} + B_2(w) \frac{\partial w}{\partial y} = S_1(w) \frac{\partial H}{\partial x} + S_2(w) \frac{\partial H}{\partial y} + S_F(w),$$

- $w(\mathbf{x}, t) : D \times (0, T) \mapsto \omega \subset \mathbb{R}^n$, the vector of unknowns,
- D bounded domain of \mathbb{R}^2 ; ω convex subset of \mathbb{R}^n ,
- $F_i : \omega \mapsto \mathbb{R}^n$, $i = 1, 2$, regular and locally bounded functions,
- $B_i : \Omega \mapsto \mathcal{M}_{N \times N}(\mathbb{R})$, $i = 1, 2$ regular and locally bounded matrix-valued functions,
- $S_i, S_F : \omega \mapsto \mathbb{R}^n$, $i = 1, 2$, regular and locally bounded functions,
- $H : D \subset \mathbb{R}^2 : \mapsto \mathbb{R}$ known function.

General framework

General formulation

$$\frac{\partial w}{\partial t} + \frac{\partial F_1}{\partial x}(w) + \frac{\partial F_2}{\partial y}(w) + B_1(w) \frac{\partial w}{\partial x} + B_2(w) \frac{\partial w}{\partial y} = S_1(w) \frac{\partial H}{\partial x} + S_2(w) \frac{\partial H}{\partial y} + S_F(w),$$

Nonconservative products

The nonconservative products $B_i(w) \frac{\partial w}{\partial \alpha}$, $S_i(w) \frac{\partial w}{\partial \alpha}$, $i = 1, 2$, $\alpha = x, y$ do not make sense in general within the framework of distributions. Here, we follow the theory developed by [Dal Maso, LeFloch and Murat] to give a sense to these products as Borel measures. This theory is based on the choice of a family of paths.

General Formulation: Nonconservative Form

General formulation

$$\frac{\partial w}{\partial t} + \frac{\partial F_1}{\partial x}(w) + \frac{\partial F_2}{\partial y}(w) + B_1(w) \frac{\partial w}{\partial x} + B_2(w) \frac{\partial w}{\partial y} = S_1(w) \frac{\partial H}{\partial x} + S_2(w) \frac{\partial H}{\partial y} + S_F(w)$$

[LeFloch] Introducing the trivial equation $\partial_t H = 0$ and taking H as a new unknown

$W := [w \ H]^\top \in \mathbb{R}^N$, $N = n + 1$, $\mathcal{J}_k(w) = \frac{\partial F_k}{\partial w}$ and

$$\mathcal{A}_k = \left[\begin{array}{c|c} \mathcal{J}_k(w) + B_k(w) & -S_k(w) \\ \hline 0 & 0 \end{array} \right], \quad k = 1, 2 \in \mathcal{M}_{N \times N},$$

General Formulation: Nonconservative Form

General formulation

$$\frac{\partial w}{\partial t} + \frac{\partial F_1}{\partial x}(w) + \frac{\partial F_2}{\partial y}(w) + B_1(w) \frac{\partial w}{\partial x} + B_2(w) \frac{\partial w}{\partial y} = S_1(w) \frac{\partial H}{\partial x} + S_2(w) \frac{\partial H}{\partial y} + S_F(w)$$

[LeFloch] Introducing the trivial equation $\partial_t H = 0$ and taking H as a new unknown

$W := [w \ H]^\top \in \mathbb{R}^N$, $N = n + 1$, $\mathcal{J}_k(w) = \frac{\partial F_k}{\partial w}$ and

$$\mathcal{A}_k = \left[\begin{array}{c|c} \frac{\mathcal{J}_k(w) + B_k(w)}{0} & -S_k(w) \\ \hline & 0 \end{array} \right], \quad k = 1, 2 \in \mathcal{M}_{N \times N},$$

the system can be written as follows:

$$\frac{\partial W}{\partial t} + \mathcal{A}_1(W) \frac{\partial W}{\partial x} + \mathcal{A}_2(W) \frac{\partial W}{\partial y} = S_F(W).$$

General Formulation: Nonconservative form

Given a unitary vector $\boldsymbol{\eta} = (\eta_1, \eta_2) \in \mathbb{R}^2$:

$$\mathcal{A}(W, \boldsymbol{\eta}) = \mathcal{A}_1(W)\eta_1 + \mathcal{A}_2(W)\eta_2.$$

We assume that the system is strictly **hyperbolic**, i.e.,

- $\forall W \in \Omega \subset \mathbb{R}^N$ and $\forall \boldsymbol{\eta} \in S^1$, the matrix $\mathcal{A}(W, \boldsymbol{\eta})$ has N **real eigenvalues**: $\lambda_1(W, \boldsymbol{\eta}) < \dots < \lambda_N(W, \boldsymbol{\eta})$, being $R_j(W, \boldsymbol{\eta})$, $j = 1, \dots, N$ the associated **eigenvectors**.
- $\mathcal{A}(W, \boldsymbol{\eta})$ is diagonalizable.

$$\mathcal{A}(W, \boldsymbol{\eta}) = \mathcal{K}(W, \boldsymbol{\eta})\mathcal{L}(W, \boldsymbol{\eta})\mathcal{K}^{-1}(W, \boldsymbol{\eta}),$$

$\mathcal{L}(W, \boldsymbol{\eta})$ **diagonal matrix** whose coefficients are the eigenvalues of $\mathcal{A}(W, \boldsymbol{\eta})$,

$\mathcal{K}(W, \boldsymbol{\eta})$ matrix whose columns are the eigenvectors.

General Formulation: Nonconservative form

A **family of paths** (Dal Maso, LeFloch, Murat) in $\Omega \subset \mathbb{R}^N$ is a locally Lipschitz map

$$\Phi: [0, 1] \times \Omega \times \Omega \times \mathcal{S}^1 \rightarrow \Omega,$$

where $\mathcal{S}^1 \subset \mathbb{R}^2$ denotes the unit sphere, that satisfies some regularity conditions and

- 1 $\Phi(0; W_L, W_R, \boldsymbol{\eta}) = W_L$ and $\Phi(1; W_L, W_R, \boldsymbol{\eta}) = W_R$, for any $W_L, W_R \in \Omega$, $\boldsymbol{\eta} \in \mathcal{S}^1$.
- 2 $\Phi(s; W_L, W_R, \boldsymbol{\eta}) = \Phi(1 - s; W_R, W_L, -\boldsymbol{\eta})$, for any $W_L, W_R \in \Omega$, $s \in [0, 1]$, $\boldsymbol{\eta} \in \mathcal{S}^1$.
- 3 $\Phi(s; W, W, \boldsymbol{\eta}) = W$, for any $s \in [0, 1]$, $\boldsymbol{\eta} \in \mathcal{S}^1$.

General Formulation: Nonconservative form

A **family of paths** (Dal Maso, LeFloch, Murat) in $\Omega \subset \mathbb{R}^N$ is a locally Lipschitz map

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- ❸ $\Phi(s; W, W, \boldsymbol{\eta}) = W$, for any $s \in [0, 1]$, $\boldsymbol{\eta} \in \mathcal{S}^1$.

A piecewise regular function W is a **weak solution** if and only if the two following conditions are satisfied:

- (i) W is a classical solution where it is smooth.
- (ii) At every point of a discontinuity W satisfies the jump condition

$$\int_0^1 (\sigma \mathcal{I} - \mathcal{A}(\Phi(s; W^-, W^+, \boldsymbol{\eta}), \boldsymbol{\eta})) \frac{\partial \Phi}{\partial s}(s; W^-, W^+, \boldsymbol{\eta}) ds = 0, \quad (1)$$

where \mathcal{I} is the identity matrix; σ , the speed of propagation of the discontinuity; $\boldsymbol{\eta}$ a unit vector normal to the discontinuity at the considered point; and W^-, W^+ , the lateral limits of the solution at the discontinuity.

General Formulation: Nonconservative form

A **family of paths** (Dal Maso, LeFloch, Murat) in $\Omega \subset \mathbb{R}^N$ is a locally Lipschitz map

$$\Phi: [0, 1] \times \Omega \times \Omega \times \mathcal{S}^1 \rightarrow \Omega,$$

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- ❶ $\Phi(0; W_L, W_R, \boldsymbol{\eta}) = W_L$ and $\Phi(1; W_L, W_R, \boldsymbol{\eta}) = W_R$, for any $W_L, W_R \in \Omega$, $\boldsymbol{\eta} \in \mathcal{S}^1$.
 - ❷ $\Phi(s; W_L, W_R, \boldsymbol{\eta}) = \Phi(1 - s; W_R, W_L, -\boldsymbol{\eta})$, for any $W_L, W_R \in \Omega$, $s \in [0, 1]$, $\boldsymbol{\eta} \in \mathcal{S}^1$.
 - ❸ $\Phi(s; W, W, \boldsymbol{\eta}) = W$, for any $s \in [0, 1]$, $\boldsymbol{\eta} \in \mathcal{S}^1$.
- The choice of the family of paths is important because it determines the speed of propagation of discontinuities.
 - It has to be based on the physical background of the problem: limit of viscous profiles . . . , but in practice can be very difficult.
 - From the mathematical point of view, some hypotheses concerning the relation of the paths with the integral curves of the characteristic fields can be imposed.

High order numerical schemes

Goal

Develop **high order finite volume schemes** for problems that can be written under the form:

$$\frac{\partial W}{\partial t} + \mathcal{A}_1(W) \frac{\partial W}{\partial x} + \mathcal{A}_2(W) \frac{\partial W}{\partial y} = 0,$$

Properties:

- Shock capturing property: the solutions of this kind of systems may generate discontinuities in finite time, even for very regular initial conditions. The appearance and subsequent propagation of these shocks has to be correctly simulated.
- High accuracy in areas where the solution is regular.
- Well-balanced property: in cases where regular non-trivial stationary solutions exist, the schemes must be able to solve these solutions (or some of them) exactly or at least with an enhanced accuracy.

Third order compact polynomial reconstructions: procedure

- For $R \in \{N, S, E, W\}$, build the smoothness indicators

$$\beta_R = \iint_V \Delta x \Delta y ([\partial_y P_R(\mathbf{x})]^2 + [\partial_x P_R(\mathbf{x})]^2) d\mathbf{x}.$$

- If some $\beta_R > \text{tol} = \Delta^2$ then redefine $P_R(\mathbf{x})$ to be a linear function providing a second order conservative approximation to $F(\mathbf{x})$ on V .
- Apply the WENO idea: define the nonlinear weights ω_R as

$$\omega_R = \frac{\alpha_R}{\alpha_N + \alpha_S + \alpha_E + \alpha_W}, \quad \alpha_R = \frac{\lambda_R}{(\varepsilon + \beta_R)^r}.$$

The linear weights λ_R are given the value $1/4$, as they are not chosen to improve the accuracy. Typically, $\varepsilon = 10^{-5}$ and $r = 4$.

- Finally, define the reconstruction operator $P(\mathbf{x})$ as

$$P(\mathbf{x}) = \omega_N P_N(\mathbf{x}) + \omega_S P_S(\mathbf{x}) + \omega_E P_E(\mathbf{x}) + \omega_W P_W(\mathbf{x})$$

High Order 2D scheme: conservative systems

Using the divergence theorem, the previous semi-discrete numerical scheme can be rewritten as follows:

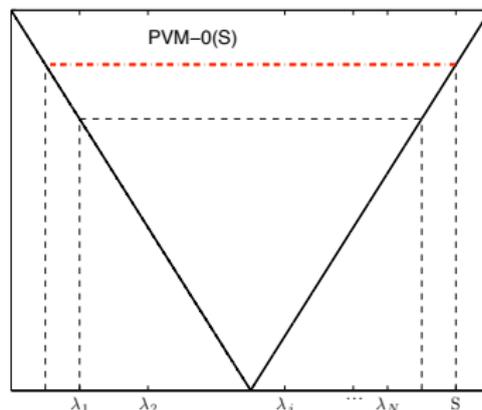
$$\begin{aligned}
 W_i'(t) &= -\frac{1}{|V_i|} \sum_{j \in \mathcal{N}_i} \int_{E_{ij}} \left(G(W_{ij}^-(\gamma, t), W_{ij}^+(\gamma, t), \boldsymbol{\eta}_{ij}) - F(W_{ij}^-(\gamma, t)) \cdot \boldsymbol{\eta}_{ij} \right) d\gamma \\
 &\quad - \frac{1}{|V_i|} \sum_{j \in \mathcal{N}_i} \int_{E_{ij}} \left(F(W_{ij}^-(\gamma, t)) \cdot \boldsymbol{\eta}_{ij} \right) d\gamma \\
 &= -\frac{1}{|V_i|} \sum_{j \in \mathcal{N}_i} \int_{E_{ij}} \left(G(W_{ij}^-(\gamma, t), W_{ij}^+(\gamma, t), \boldsymbol{\eta}_{ij}) - F(W_{ij}^-(\gamma, t)) \cdot \boldsymbol{\eta}_{ij} \right) d\gamma \\
 &\quad - \frac{1}{|V_i|} \int_{V_i} \nabla \cdot (F \circ P_i^t)(\mathbf{x}) d\mathbf{x} \\
 &= -\frac{1}{|V_i|} \sum_{j \in \mathcal{N}_i} \int_{E_{ij}} \left(G(W_{ij}^-(\gamma, t), W_{ij}^+(\gamma, t), \boldsymbol{\eta}_{ij}) - F(W_{ij}^-(\gamma, t)) \cdot \boldsymbol{\eta}_{ij} \right) d\gamma \\
 &\quad - \frac{1}{|V_i|} \int_{V_i} \left(\mathcal{J}_1(P_i^t(\mathbf{x})) \frac{\partial P_i^t}{\partial x}(\mathbf{x}) + \mathcal{J}_2(P_i^t(\mathbf{x})) \frac{\partial P_i^t}{\partial y}(\mathbf{x}) \right) dx,
 \end{aligned}$$

PVM-0(S_0) methods: Rusanov, Lax-Friedrichs and modified Lax-Friedrichs schemes

$$P_0(x) = S_0,$$

$$S_0 \in \{S_{Rus}, S_{LF}, S_{LF}^{mod}\}, \quad S_{Rus} = \max_j |\lambda_{j,i+1/2}|, \quad S_{LF} = \frac{\Delta x}{\Delta t} \quad \text{and} \quad S_{LF}^{mod} = \alpha \frac{\Delta x}{\Delta t}.$$

- Rusanov scheme corresponds to the choice $S_0 = S_{Rus}$,
- Lax-Friedrichs with $S_0 = S_{LF}$
- modified Lax-Friedrichs with $S_0 = S_{LF}^{mod}$.



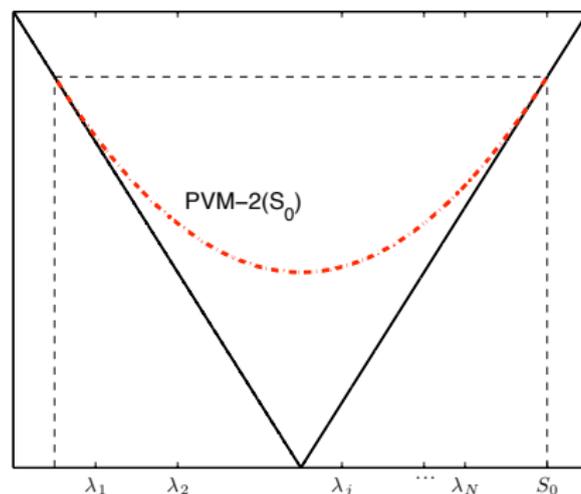
PVM-2(S_0) methods or FORCE type methods

$$P_2(x) = \alpha_0 + \alpha_2 x^2, \text{ such as } P_2(S_0) = S_0, P_2'(S_0) = 1, S_0 \in \{S_{Rus}, S_{LF}, S_{LF}^{mod}\}.$$

Remarks

- If $S_0 = S_{LF}$ then we obtain FORCE method.
- GFORCE scheme can be obtained by imposing

$$P_2(S_{LF}^{mod}) = S_{LF}^{mod}, \quad P_2'(S_{LF}^{mod}) = \frac{2\alpha}{1 + \alpha},$$

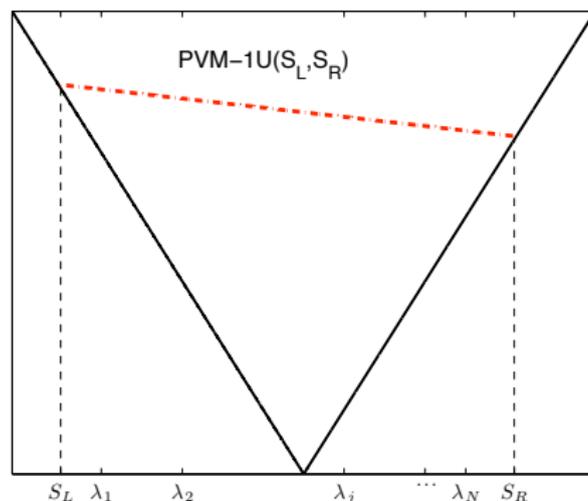


PVM-1U(S_L, S_R) or HLL method

Let S_L (respectively S_R) be an approximation of the minimum (respectively maximum) wave speed, we consider the polynomial

$$P_1(x) = \alpha_0 + \alpha_1 x \quad \text{such as } P_1(S_L) = |S_L|, \quad P_1(S_R) = |S_R|.$$

$$Q_{i+1/2} = \alpha_0 Id + \alpha_1 A_{i+1/2}$$



PVM-2U(S_L, S_R) method

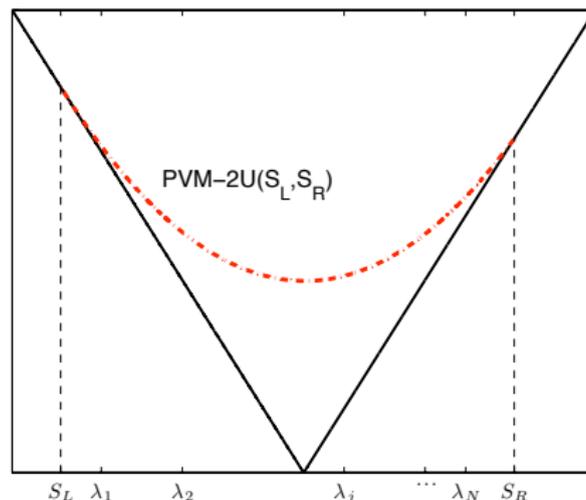
$$P_2(x) = \alpha_0 + \alpha_1 x + \alpha_2 x^2,$$

such as

$$P_2(S_m) = |S_m|, \quad P_2(S_M) = |S_M|, \quad P_2'(S_M) = \text{sgn}(S_M),$$

where

$$S_M = \begin{cases} S_L & \text{if } |S_L| \geq |S_R|, \\ S_R & \text{if } |S_L| < |S_R|. \end{cases} \quad S_m = \begin{cases} S_R & \text{if } |S_L| \geq |S_R|, \\ S_L & \text{if } |S_L| < |S_R|. \end{cases}$$



PVM-2U(S_L, S_R, S_{int}) method or IFCP method

$$P_2(x) = \alpha_0 + \alpha_1 x + \alpha_2 x^2,$$

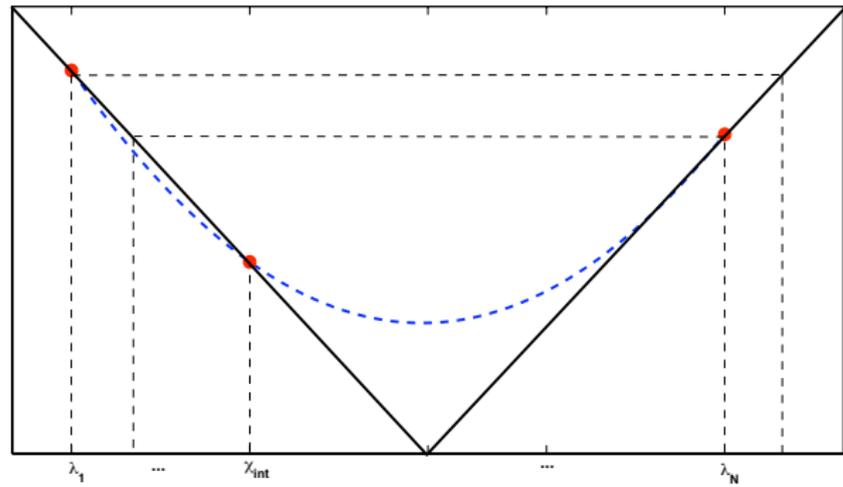
$$\begin{pmatrix} 1 & S_L & S_L^2 \\ 1 & S_R & S_R^2 \\ 1 & S_{int} & S_{int}^2 \end{pmatrix} \begin{pmatrix} \alpha_0 \\ \alpha_1 \\ \alpha_2 \end{pmatrix} = \begin{pmatrix} |S_L| \\ |S_R| \\ |S_{int}| \end{pmatrix},$$

S_L (respectively S_R) is an approximation of the minimum (respectively maximum) wave speed and

$$S_{int} = S_{ext} \max(|S_2|, \dots, |S_{N-1}|),$$

$$S_{ext} = \begin{cases} \text{sgn}(S_L + S_R), & \text{if } (S_L + S_R) \neq 0, \\ 1, & \text{otherwise.} \end{cases}$$

PVM-2U(S_L, S_R, S_{int}) method or IFCP method



PVM-Chebyshev method

- Chebyshev polynomials provide optimal approximation to the absolute value function in $[-1, 1]$:

$$|x| = \frac{2}{\pi} + \sum_{k=1}^{\infty} \frac{4}{\pi} \frac{(-1)^{k+1}}{(2k-1)(2k+1)} T_{2k}(x), \quad x \in [-1, 1],$$

where the Chebyshev polynomials of even degree $T_{2k}(x)$ are recursively defined as

$$T_0(x) = 1, \quad T_2(x) = 2x^2 - 1, \quad T_{2k}(x) = 2T_2(x)T_{2k-2}(x) - T_{2k-4}(x).$$

- The polynomial

$$\tau_{2p}(x) = \frac{2}{\pi} + \sum_{k=1}^p \frac{4}{\pi} \frac{(-1)^{k+1}}{(2k-1)(2k+1)} T_{2k}(x), \quad x \in [-1, 1],$$

verifies the equality

$$\| |x| - \tau_{2p}(x) \|_{\infty} = \frac{2}{\pi} \frac{1}{2p+1},$$

which is optimal in $L^{\infty}(-1, 1)$ (see [Bernstein 1913](#)).

-

$$Q_{i+1/2} = P_{2p}(A_{i+1/2}) = |\lambda_{\max}| \tau_{2p} \left(\frac{1}{|\lambda_{\max}|} A_{i+1/2} \right) \approx |A_{i+1/2}|,$$

where λ_{\max} is the eigenvalue of $A_{i+1/2}$ with maximum absolute value.

PVM-Chebyshev method

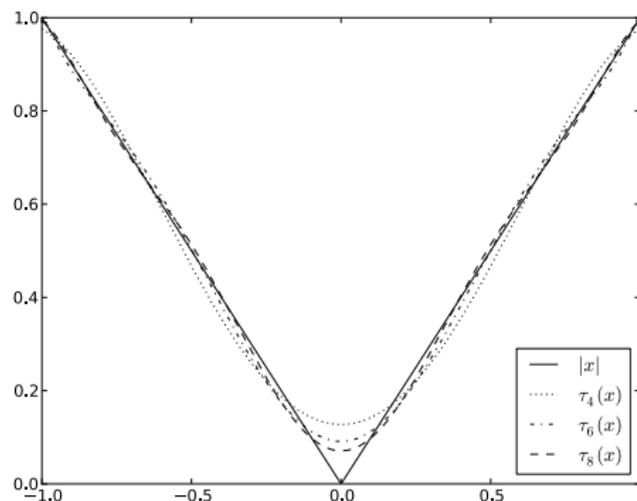


Figure : The Chebyshev approximations $\tau_{2p}(x)$ for $p = 2, 3, 4$.

- Notice that $\tau_{2p}(x)$ do not strictly satisfy the stability condition $\tau_{2p}(x) \geq |x|$. This drawback can be avoided by substituting $\tau_{2p}(x)$ by $\tau_{2p}^\epsilon(x)$, such that $\tau_{2p}^\epsilon(x) \geq |x|$.

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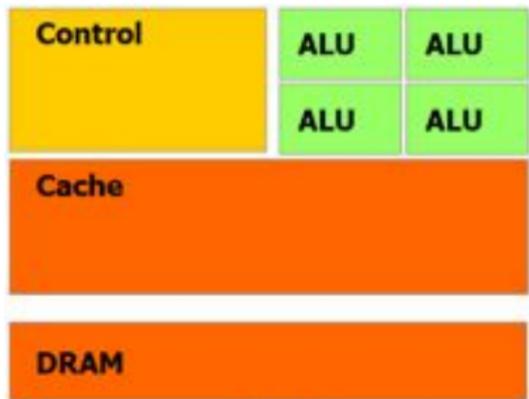
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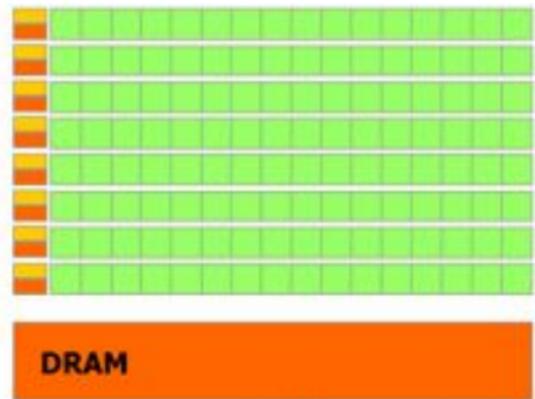
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CPU vs GPU architectures



CPU



GPU

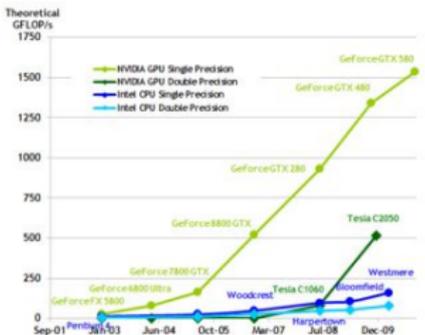
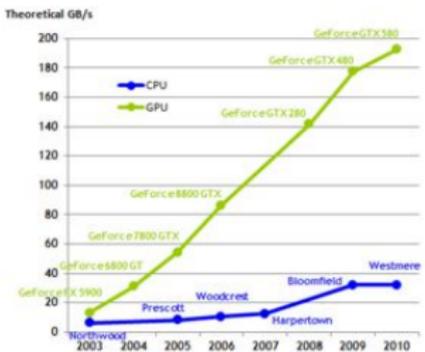
- GPUs are specialized for compute-intensive highly parallel computations.
- GPU are especially well-suited to address problems that present a high data-parallel paradigm.
- The same program is executed for each data element in GPUs, so the flow control of the code is simple.
- The code it is executed on many data elements in GPUS, so the memory access latency can be hidden with calculations instead of big data caches.

Why GPUs?

L3 Cache

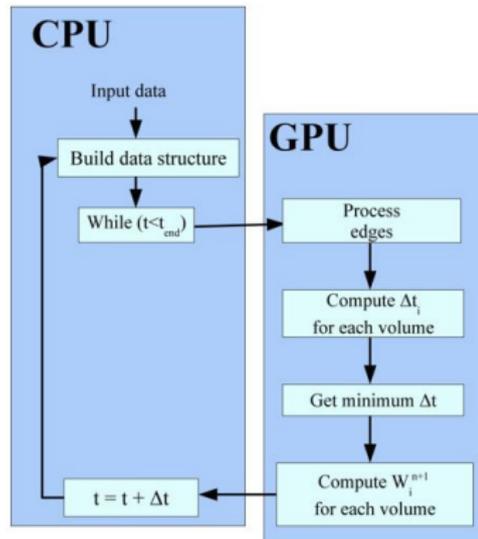
• GPUs

- GeForce FX 5900: 450 MHz, 256 MB
- GeForce 8800 GTX: 128 Cores, 575 MHz, 728 MB
- GeForce GTX 580: 512 Cores, 1544 MHz, 1546 MB



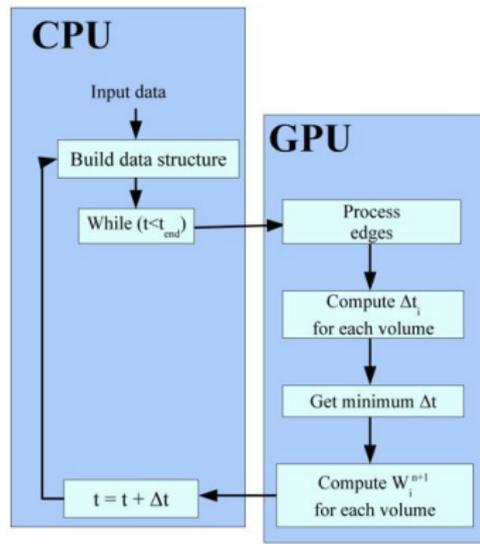
GPU Code Example

- 1 Input data in CPU.
- 2 Build data structure as arrays in CPU.
- 3 Loop in time.



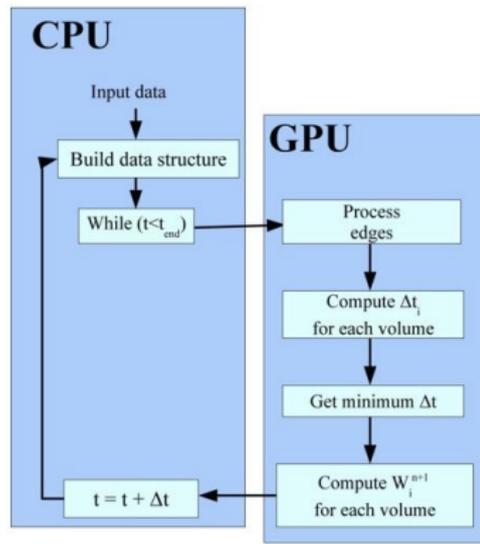
GPU Code Example

- 1 Input data in CPU.
- 2 Build data structure as arrays in CPU.
- 3 Loop in time.
 - Copy CPU arrays in GPU arrays.
 - Data is stored in GPU in float textures.



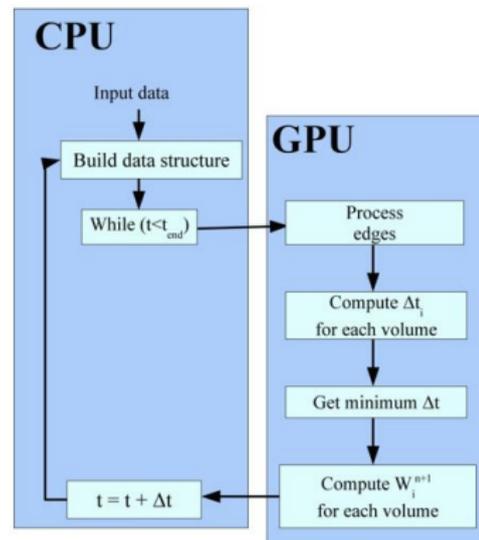
GPU Code Example

- 1 Input data in CPU.
- 2 Build data structure as arrays in CPU.
- 3 Loop in time.
 - 1 Copy CPU arrays in GPU arrays.
Data is stored in GPU in *float4* textures.
 - 2 Calculate and copy the contributions D_i^{n+1} and D_i^n to *float4* arrays.



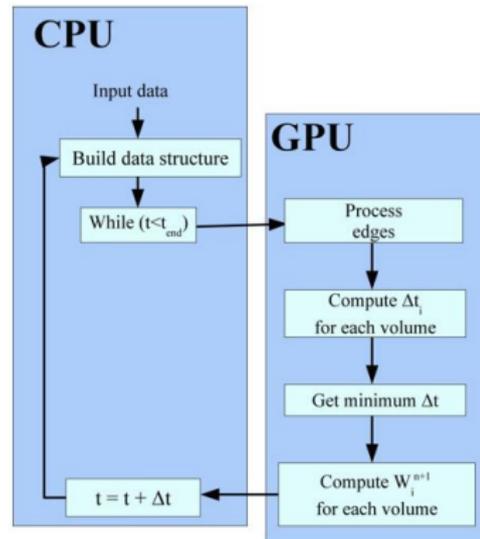
GPU Code Example

- 1 Input data in CPU.
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- 3 Loop in time.
 - 1 Copy CPU arrays in GPU arrays. Data is stored in GPU in *float4 textures*.
 - 2 Calculate and copy the contributions $D_{i-\frac{1}{2}}^+$ and $D_{i+\frac{1}{2}}^-$ to *float4* arrays.
 - 3 Compute Δt_i for each volume.



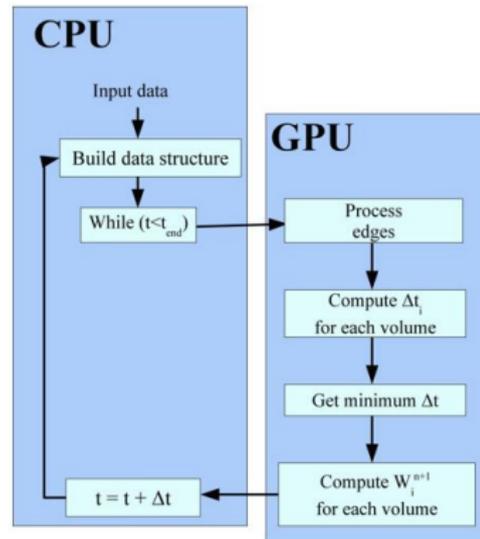
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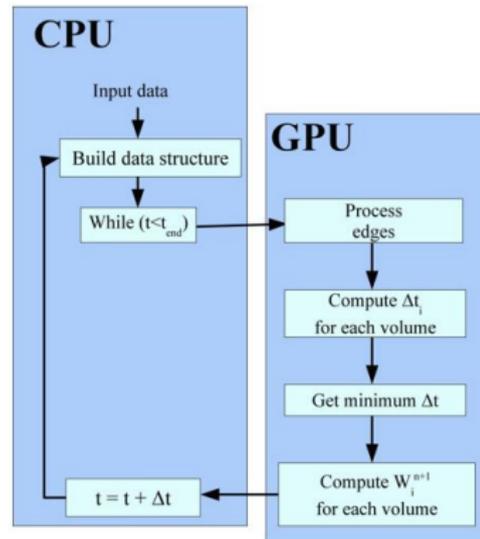
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 - 5 Calculate $W_i^{n+1} = W_i^n + \Delta t (D_{i-\frac{1}{2}}^+ + D_{i+\frac{1}{2}}^-)$.



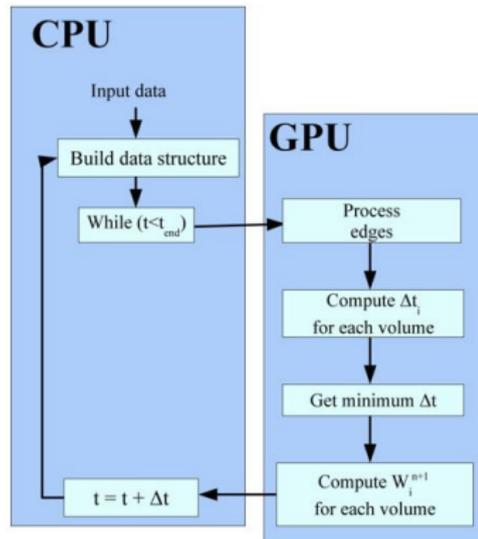
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 - 3 Compute Δt_i for each volume.
 - 4 Get the minimum of Δt_i .
 - 5 Calculate $W_i^{n+1} = W_i^n - \frac{\Delta t}{\Delta x} (D_{i-\frac{1}{2}}^+ + D_{i+\frac{1}{2}}^-)$.
 - 6 Next step in time.



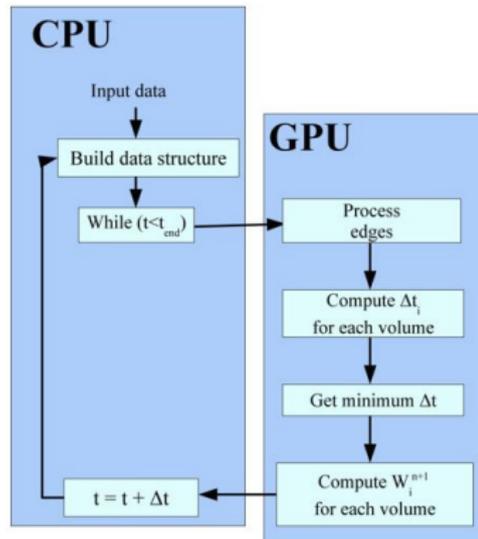
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 - 6 Next step in time.



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An example: First order PVM scheme for the two-layer shallow-water system

For each PVM method:

- **Two implementations in CPU** using C++ and the Eigen library: one for one thread and another for 4 threads using OpenMP. Both programs use double precision.
- **One implementation in GPU** using CUDA and single precision.
- Computer: Intel Core i7 920, 4 GB RAM.
- Graphics card: GeForce GTX 570.

An example: First order PVM scheme for the two-layer shallow-water system

- Initial state:

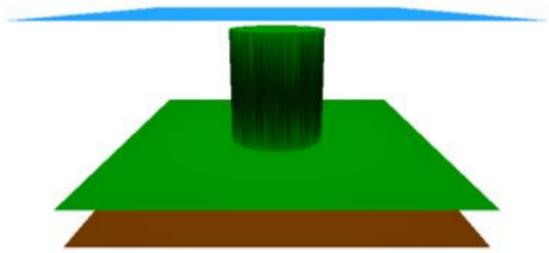
$$H(x, y) = 5$$

$$W_i^0(x, y) = (h_1(x, y), 0, 0, h_2(x, y), 0, 0)^T$$

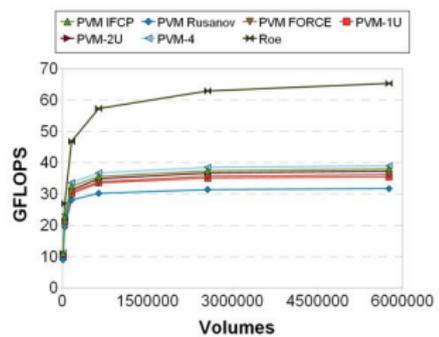
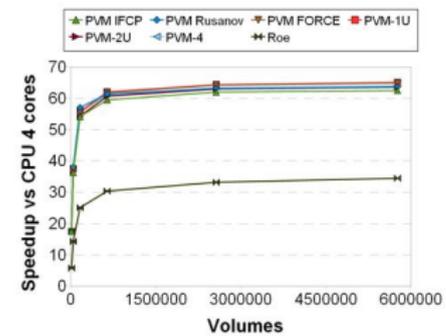
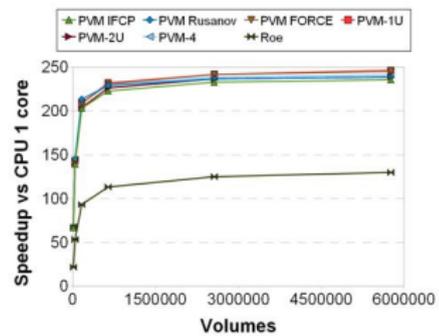
$$h_1(x, y) = \begin{cases} 4 & \text{if } \sqrt{x^2 + y^2} > 1.5 \\ 0.5 & \text{otherwise} \end{cases}$$

$$h_2(x, y) = 5 - h_1(x, y)$$

- Wall boundary conditions
- Time interval [0, 0.1]
- $\gamma = 0.9, r = 0.998$
- Different mesh sizes, up to 5760000 volumes

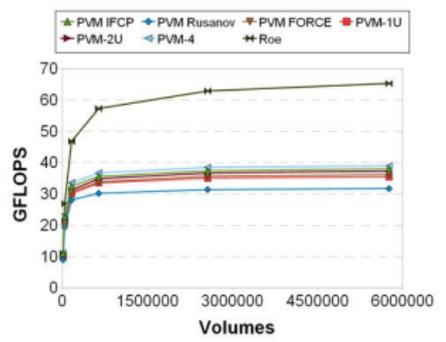
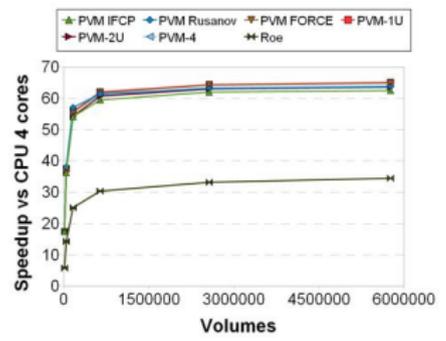
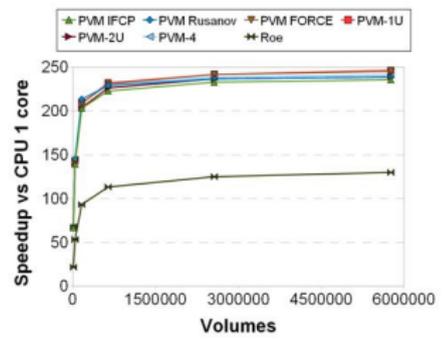


An example (Two-layer shallow-water system)



- PVM methods have been 1.9 times faster than Roe method on a GTX 570.
- Roe method has reached more GFLOPS than PVM methods.

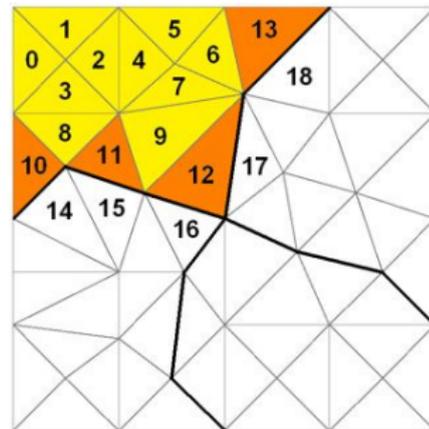
An example (Two-layer shallow-water system)



- PVM methods have been 1.9 times faster than Roe method on a GTX 570.
- Roe method has reached more GFLOPS than PVM methods.

Multi-GPU Implementation

- The triangular mesh is divided into submeshes. Each submesh is assigned to a CPU process, which uses a GPU to perform its computations.
- Volume indexation:
 - 1 Non-communication volumes.
 - 2 Comm. volumes of the submesh.
 - 3 Comm. volumes of adjacent submeshes.
- The communication volumes of the submesh must be reordered, overlapping the sendings.



References

References



M. Lastra, J. M. Mantas, C. Ureña, M. J. Castro, J. A. García- Rodríguez,

Simulation of shallow-water systems using graphics processing units, Mathematics and Computers in Simulation 80 (3),598-618, 2009.



M. de la Asunción, J. M. Mantas, M. J. Castro,

Simulation of one-layer shallow water systems on multicore and CUDA architectures, The Journal of Supercomputing, 2010. Online.



M. J. Castro, S. Ortega, M. de la Asunción, J. M. Mantas, J. M. Gallardo,

GPU computing for shallow water flow simulation based on finite volume schemes, Comptes Rendus Mécanique 339 (2-3),165-184, 2011. Special issue on High Performance Computing.



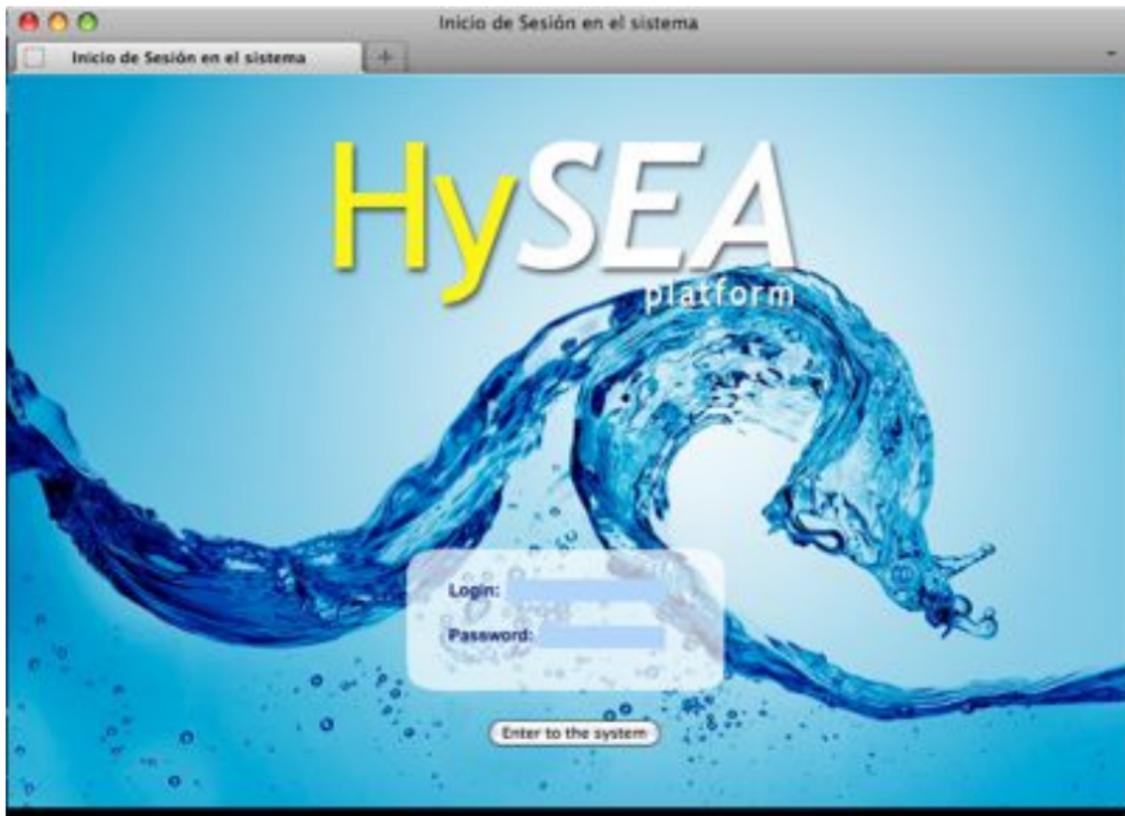
M. de la Asunción, J.M. Mantas, M.J. Castro, E.D. Fernández-Nieto.

An MPI-CUDA implementation of an improved Roe method for two-layer shallow water systems on non-structured meshes. submitted to Journal of Parallel and Distributed computing, 72 (9): 1065-1072, 2012.



M de la Asunción, MJ Castro, ED Fernández-Nieto, JM Mantas, S. Ortega, J. Manuel González-Vida.

Efficient GPU implementation of a two waves TVD-WAF method for the two-dimensional one layer shallow water system on structured meshes. Accepted on Computers & Fluids, 2012.



The screenshot shows a web browser window titled "HySea Project". The browser's address bar contains "HySea Project" and a plus sign. The page header features a blue banner with the text "HySEA v3.0" and "Hyperbolic Systems and Efficient Algorithms". Below the banner is a navigation menu with links: Home, LAM/MPJ Admin, System Status, Computations, Utilities, User Admin, Help, and Logout. The main content area is a dark blue rectangle with the text "Hi, mosaiccol." and "Please, select one option from the above menu." At the bottom of this area is a quote: "Hay algo que Dios ha hecho mal. A todo le puse limites menos a la tontería." (Konrad Adenauer). The browser window has standard OS window controls (red, yellow, green buttons) in the top-left corner.

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tsunami-HySEA-Is for structured grids - General Data

Experiment Comments:
(max 200 digits)

Output file name:
(without extension) (max 40 digits)

Grid from a file?:
Select your Grid file:
(max 150 digits)

xmin:
(max 15 digits)

xmax:
(max 15 digits)

ymin:
(max 15 digits)

ymax:
(max 15 digits)

nx:
(max 15 digits)

ny:
(max 15 digits)

HySea Project

HySea Project

HySEA
Hydrodynamic Systems and Efficient Algorithms

Version: 1.0.0

Home | LAN/MPI Admin | System Status | Computations | Utilities | User Admin | Help | Logout

System Status
Restart up.

(max 700 digits)

Second Layer y-discharge (q2y):

Scalar qy:
qy=0.;
return qy;

(max 700 digits)

Boundary Conditions

Generate Contour Conditions

Tipo de CC para la referencia 1

Free

Wall

Choose option

Tipo de CC para la referencia 2

Free

Wall

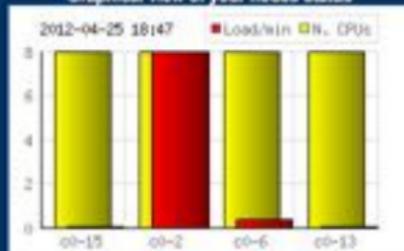


Please, select your execution scripts:

turbo-HySEA_1D-522 Ex.Date: 2011-06-10 12:29:04. Comments: Acta Nuevo completo - Experimento 3 - 3 especies - viscosidad cinematica 1.36e-6

Don't forget to run the parallel machines before start your execution.

Graphical view of your nodes status



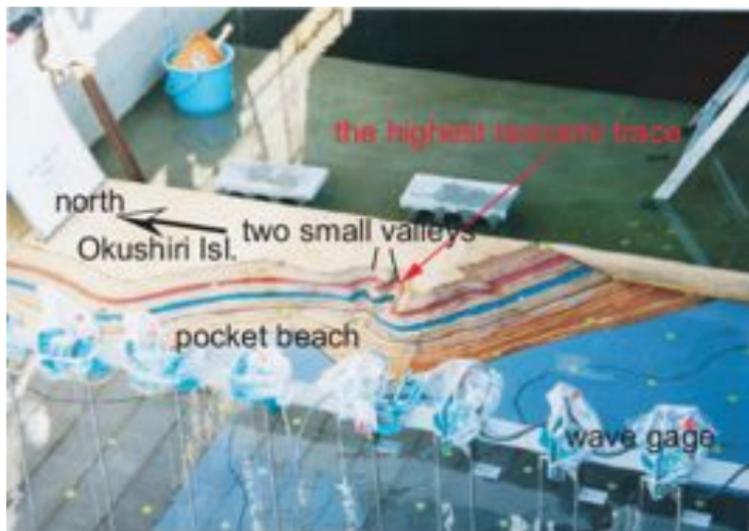
Select the startup node for this job:

Run code

Monai valley: Laboratory Experiment

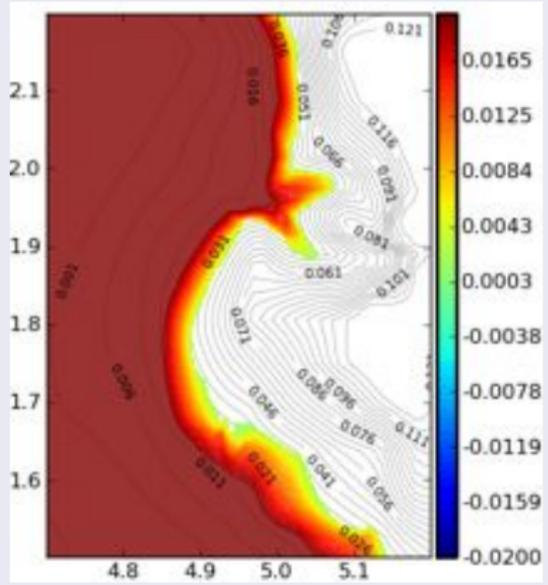
Matsuyama and Tanaka (2001)

- Maximum run-up height at Monai zone in Okushiri Island (1993)
- Scale 1:400
- 3.4 m wide, 205 m long flume
- Maximum depth 6 m.

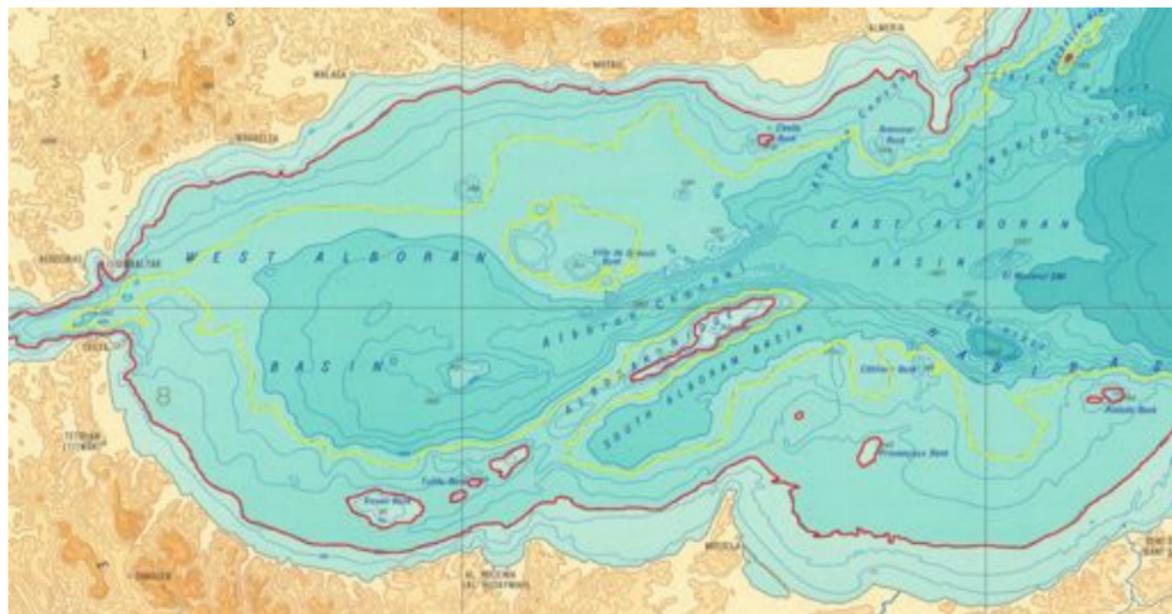


Monai valley: Frontal wave location

Physical model (Frame 70) - Numerical simulation (Time 17 s)



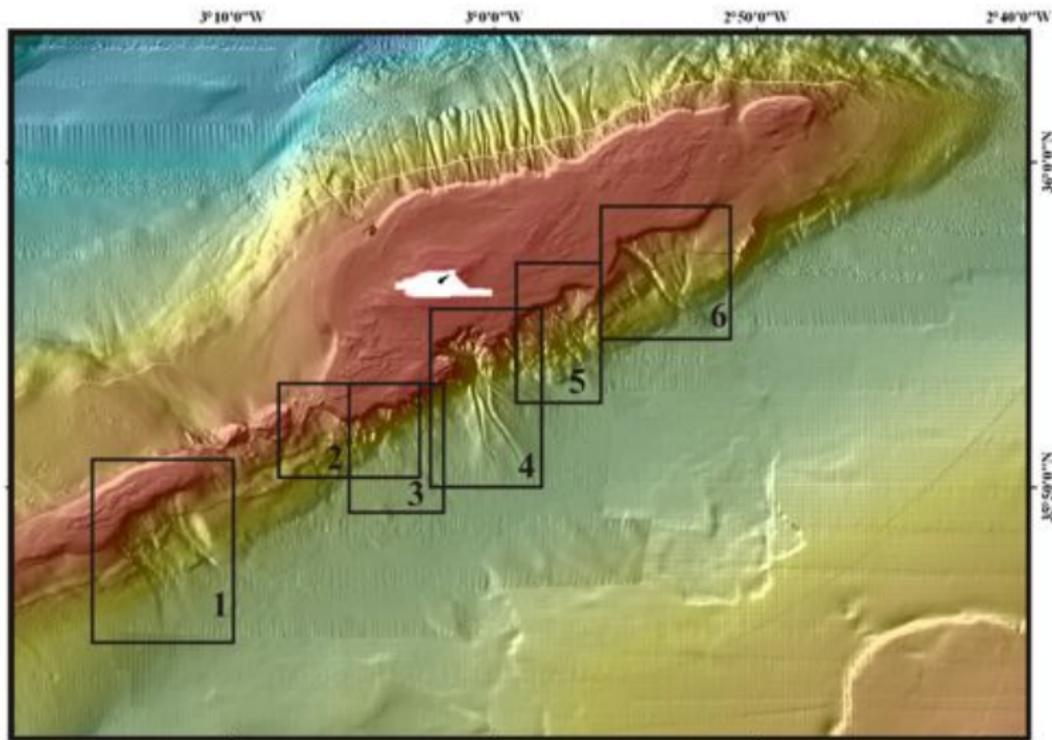
Tsunami at the Alboran Sea I



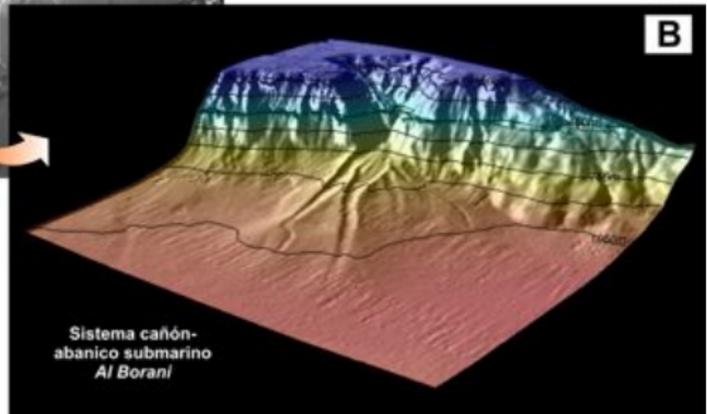
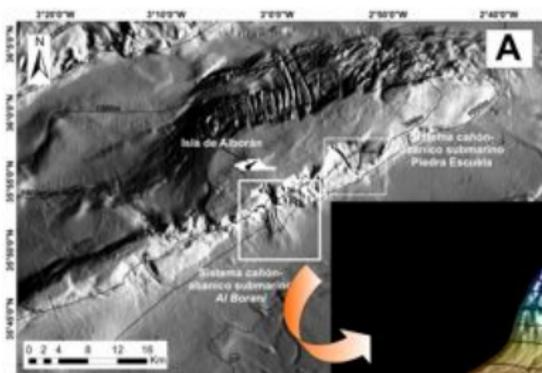
Tsunami at the Alboran Sea I



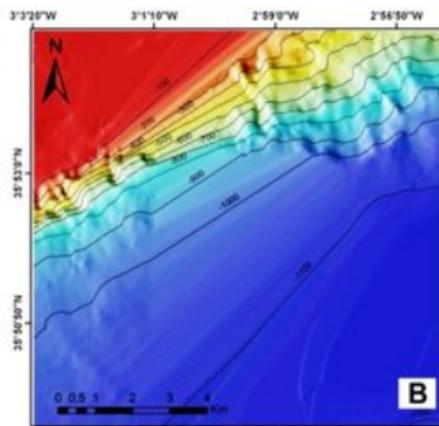
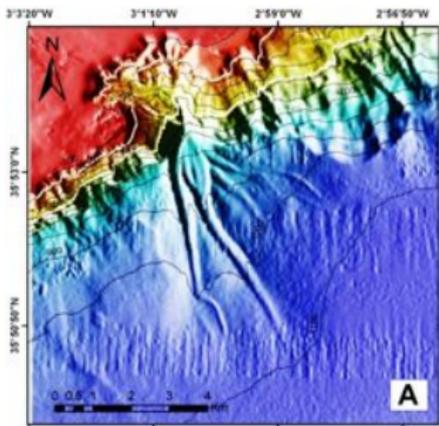
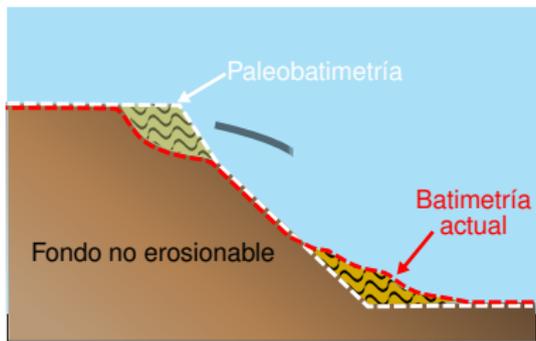
Tsunami at the Alboran Sea I



Tsunami at the Alboran Sea I

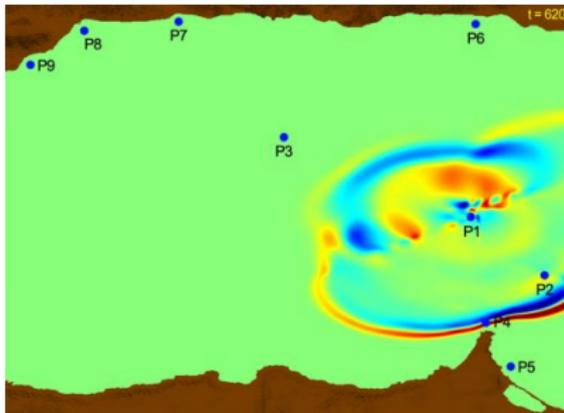
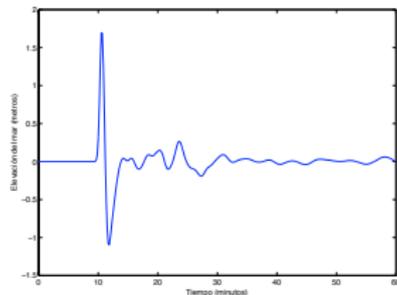
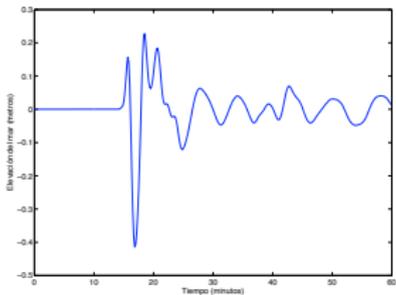


Tsunami at the Alboran Sea II: Bathymetry reconstruction



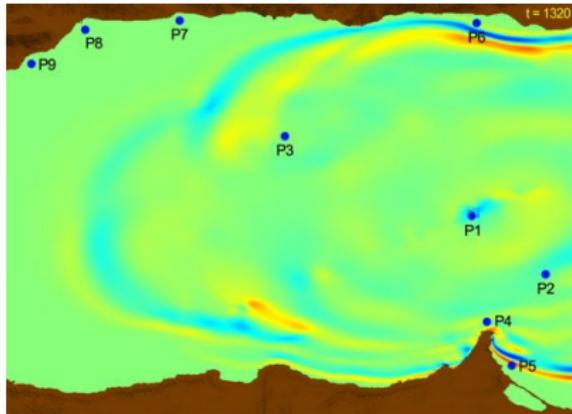
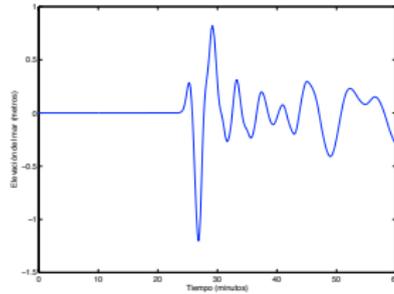
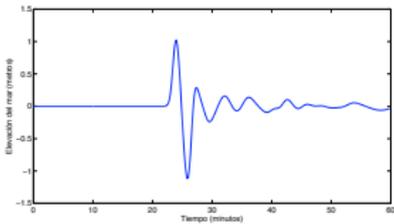
Tsunami at the Alboran Sea V: simulation

Puntos P3 (NO - open sea) - P4 (Tres Forcas)



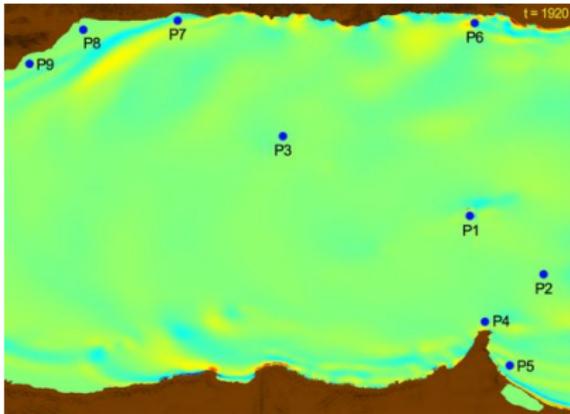
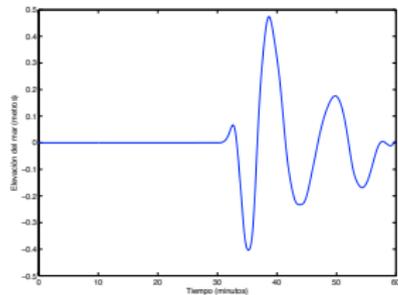
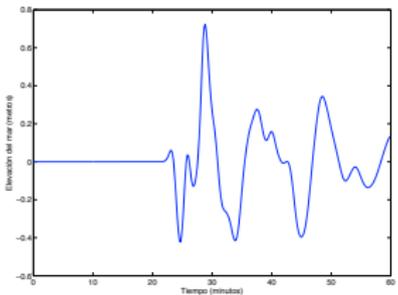
Tsunami at the Alboran Sea V: simulation

P5 (Melilla) - P2 (Adra)



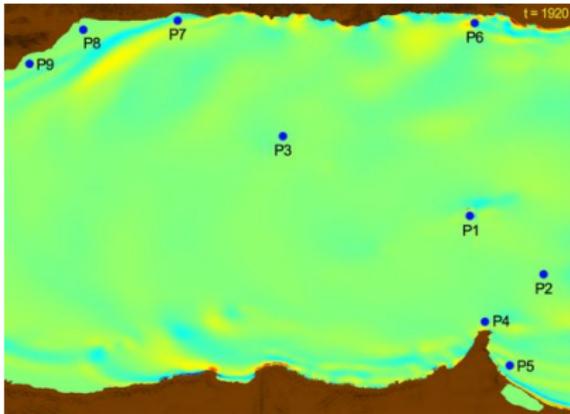
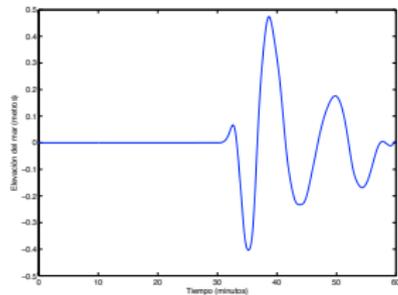
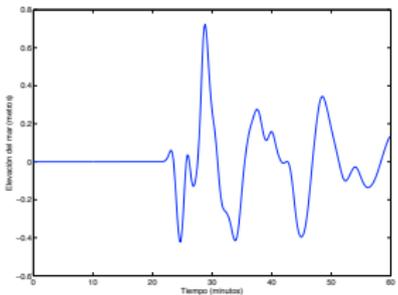
Tsunami at the Alboran Sea V: simulation

P7 (Torre del Mar) - P8 (Málaga)



Tsunami at the Alboran Sea V: simulation

P7 (Torre del Mar) - P8 (Málaga)



Lituya Bay: Earthquake & Landslide

- At 10:16pm (local time) on July 10, 1958. M_w 8.3 earthquake.
- Southwest sides and bottoms of Gilbert and Crillon inlets moved northwestward and relative to the northeast shore at the head of the bay, on the opposite side of the Fairweather fault.
- Shaking lasted about 4 minutes. Estimated total movements of 6.4m horizontally and 1 m vertically.
- About 2 minutes after the beginning of the earthquake the landslide was triggered.
- Slide volume estimated by Miller, 1960 in $30.6 \times 10^6 m^3$.

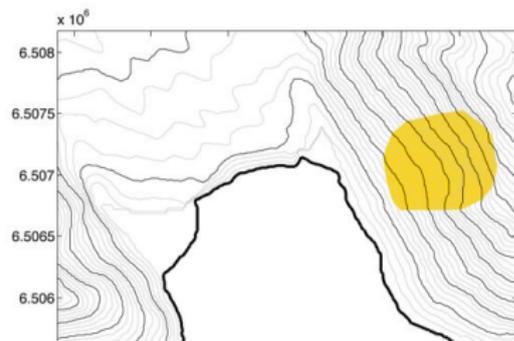
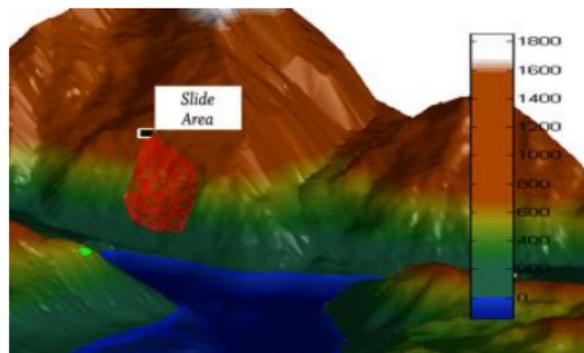
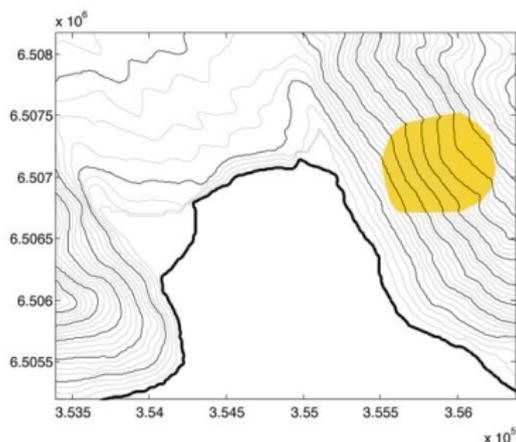


Picture from Weiss et. al., 2009

Lituya Bay: Tsunami source: reconstruction of the slide

Landslide setup

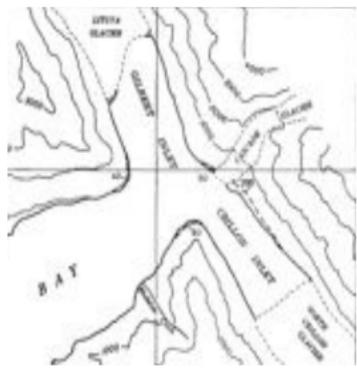
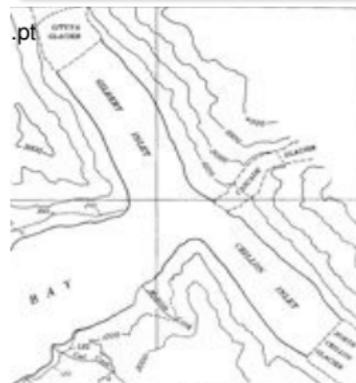
- Determine the slide perimeter.
- Surface centroid located at 610 m elevation.
- Volume of reconstructed slide: $30.625 \times 10^6 \text{ m}^3$.



Lituya Bay: Topo-bathymetry reconstruction

Gilbert Inlet Glacier front reconstruction

- Prior to the 1958 low deltas of gravel had built out into Gilbert Inlet at the southwest and northeast margins of the Lituya Glacier front.
- After the 1958 tsunami the delta on the northeast side of Gilbert Inlet completely disappeared, and the delta on the southwest side was much smaller.
- According to [Miller, 1960], the glacier surface for several hundred feet from the front was severely crevassed.



Head of Lituya Bay, showing changes in shoreline and glacier fronts over time from [Miller, 1960].

Lituya Bay: Topo-bathymetry reconstruction

Gilbert Inlet Glacier front reconstruction

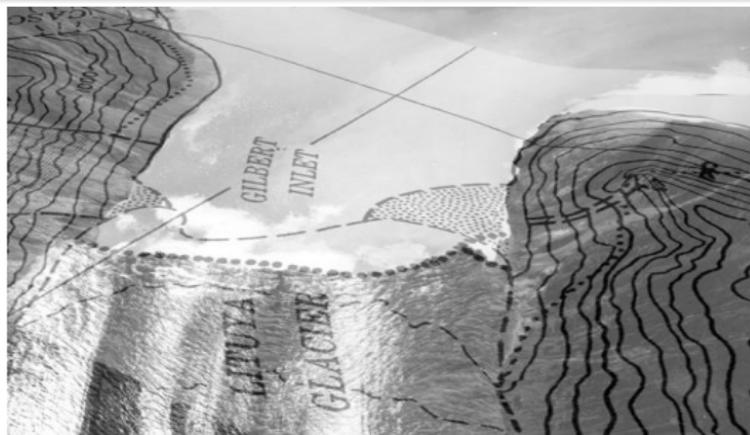
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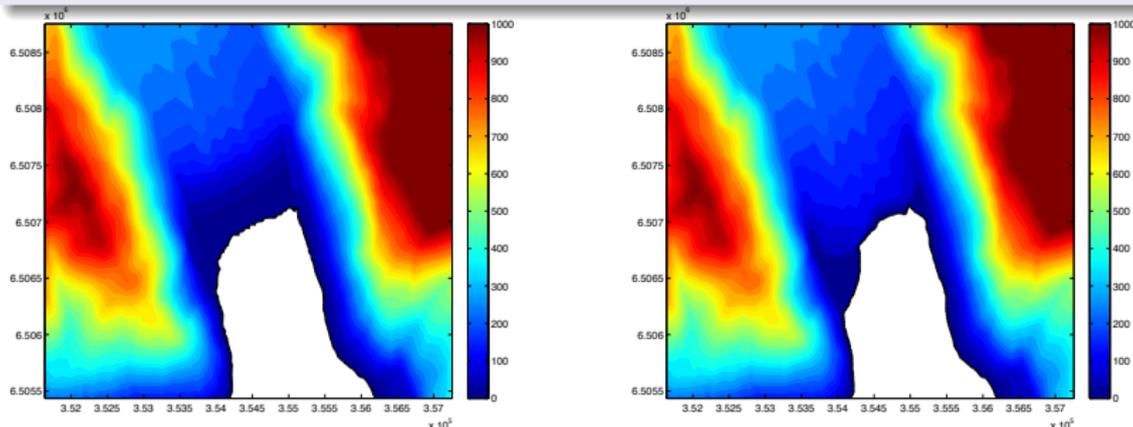
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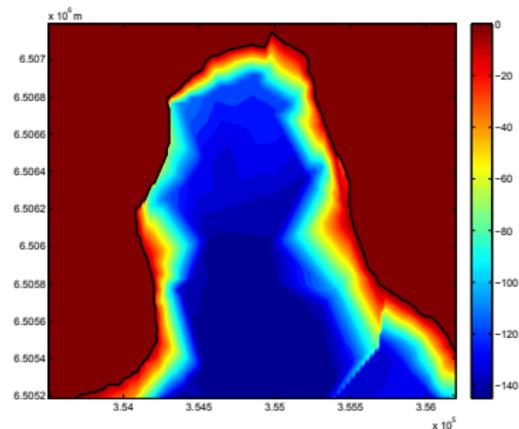
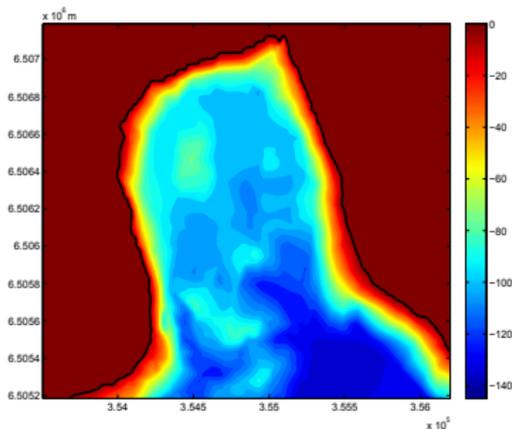


Gilbert Inlet shoreline and glacier front before and after the 1958 tsunami.

Lituya Bay: reconstruction of the bathymetry

Bathymetric data (National Ocean Service (USA))

- Hydrographic Survey ID: H08492 (Digital sounding), 1959.
- Hydrographic Survey ID: H04608, 1926.

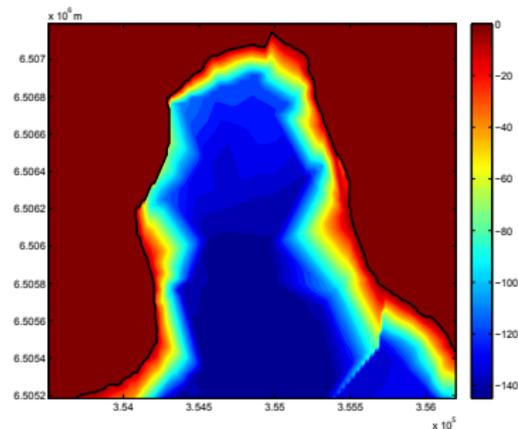
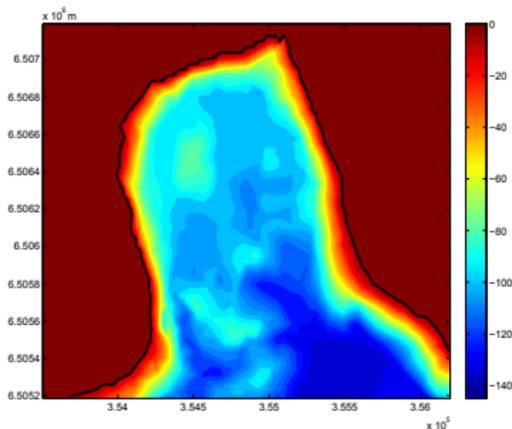


Gilbert Inlet bathymetries and shorelines before and after 1958.

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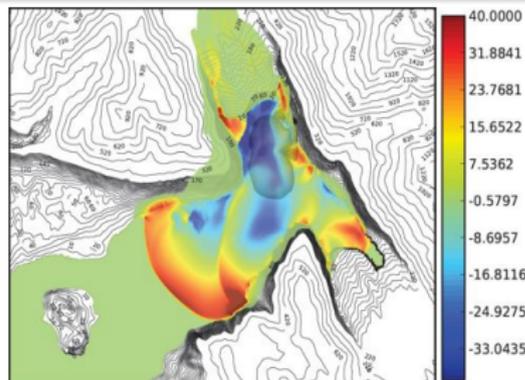
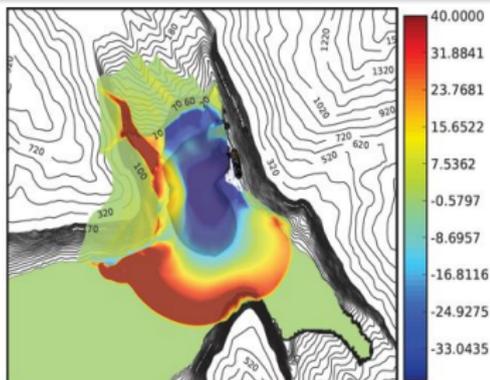


Gilbert Inlet bathymetries and shorelines before and after 1958.

Lituya Bay: Parameter sensitivity analysis

Criteria selected in order to get the optimal parameters:

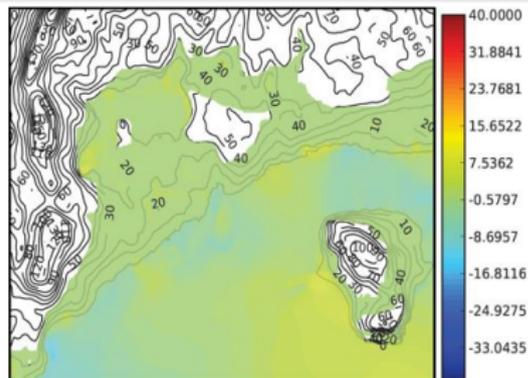
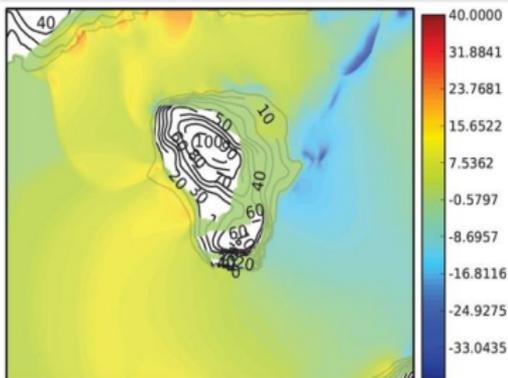
- the runup on the spur southwest of Gilbert Inlet had to be the closest to the optimal 524 m,
- the wave moving southwards towards the main stem of Lituya Bay had a peak close to 208 m in the vicinity of Mudslide Creek,
- the simulated wave had run over through the Cenotaph Island, opening a narrow channel through the trees [Miller, 1960],
- the trimline maximum distance of 1,100 m from high-tide shoreline at Fish Lake had to be reached.



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- the trimline maximum distance of 1,100 m from high-tide shoreline at Fish Lake had to be reached.



Parameter sensitivity analysis

Optimal parameters found:

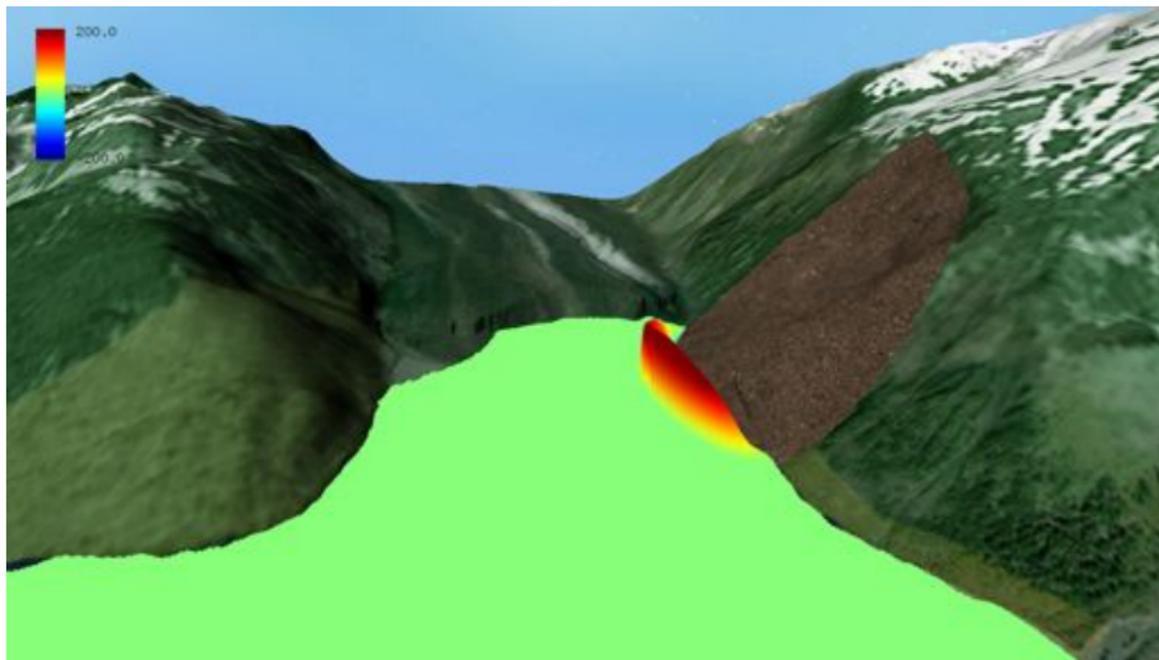
α = 13 degrees,
 r = 0.44,
 m_f = 0.08.

Lituya Bay: Settings

- Simulated time: 10 min.
- Mesh resolution: $3650 \times 1271 = 4639150$ volumes ($4\text{m} \times 7.5\text{m}$).
- CFL: 0.9
- Ratio of densities: $r = 0.44$
- Coulomb angle of repose: 13° .
- Coef. Friction between layers: 0.08.
- Coef. Friction water-bottom: $2 \cdot 10^{-3}$.
- Velocity of impact sediment/water 110 m/s.

Lituya Bay: Model results

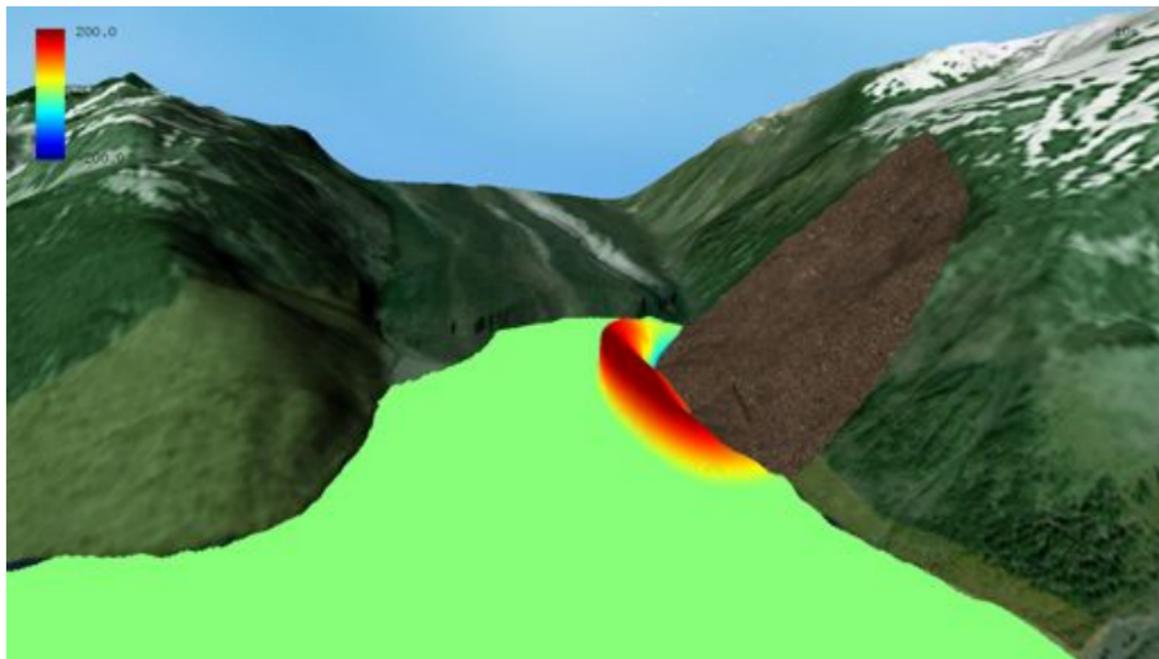
Giant wave generation



$t = 8$ s. The giant wave reaches 272.4m height

Lituya Bay: Model results

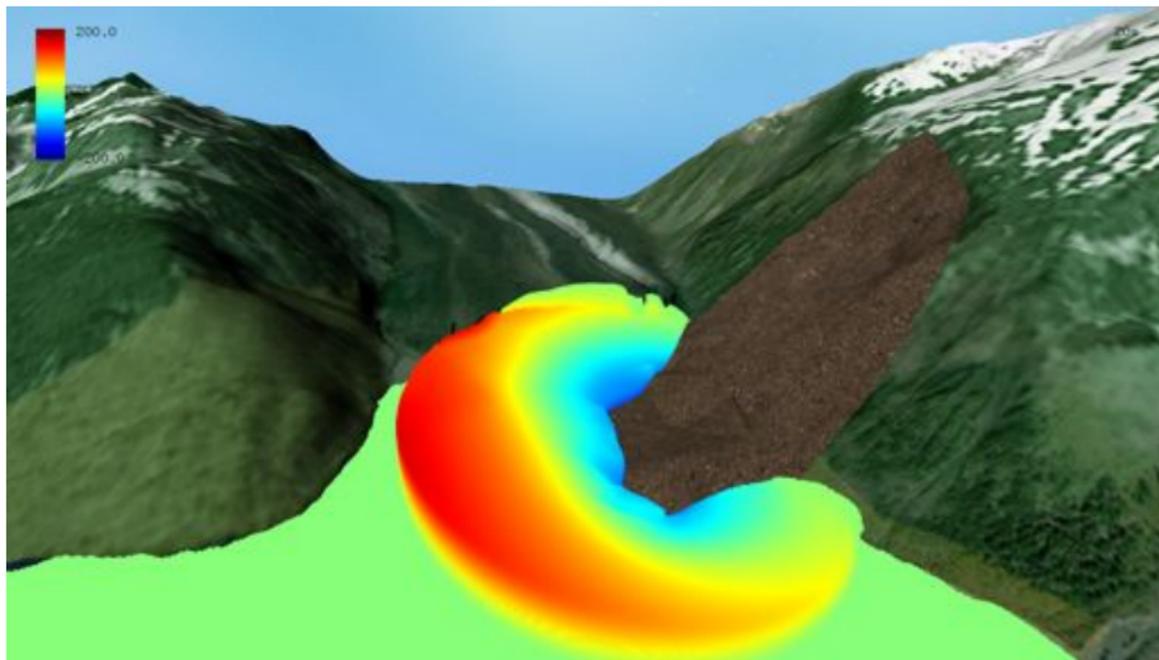
Giant wave generation



$t = 10$ s. The giant wave reaches 251.1m height

Lituya Bay: Model results

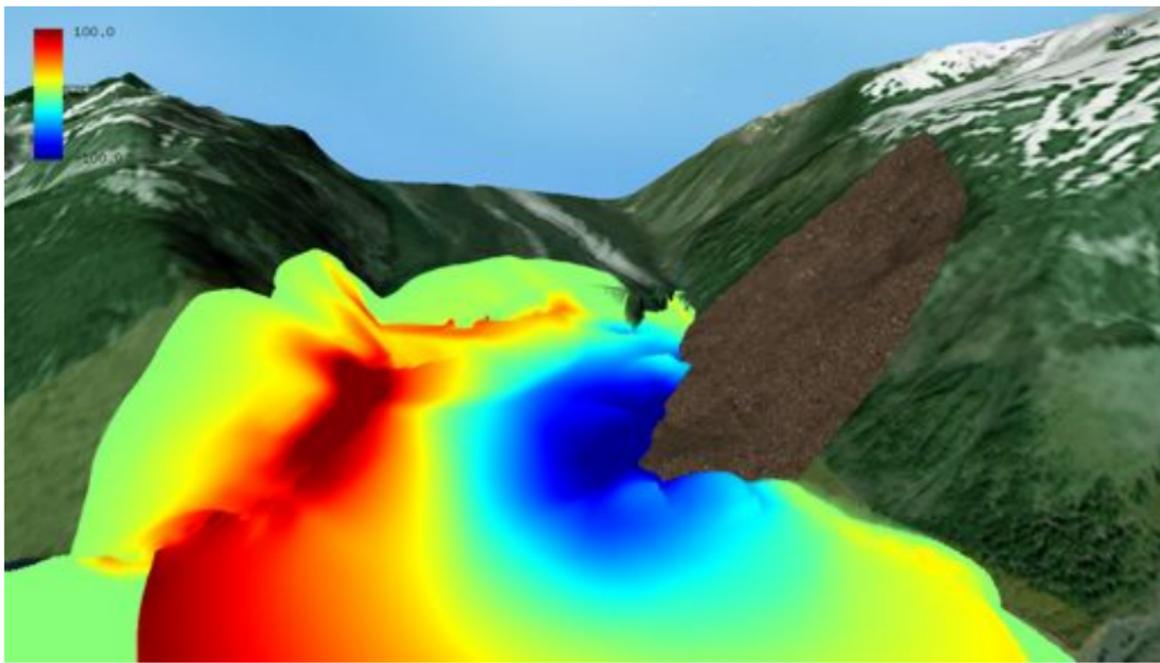
Giant wave generation



$t = 20$ s. Maximum wave height reaches 161.5m height

Lituya Bay: Model results

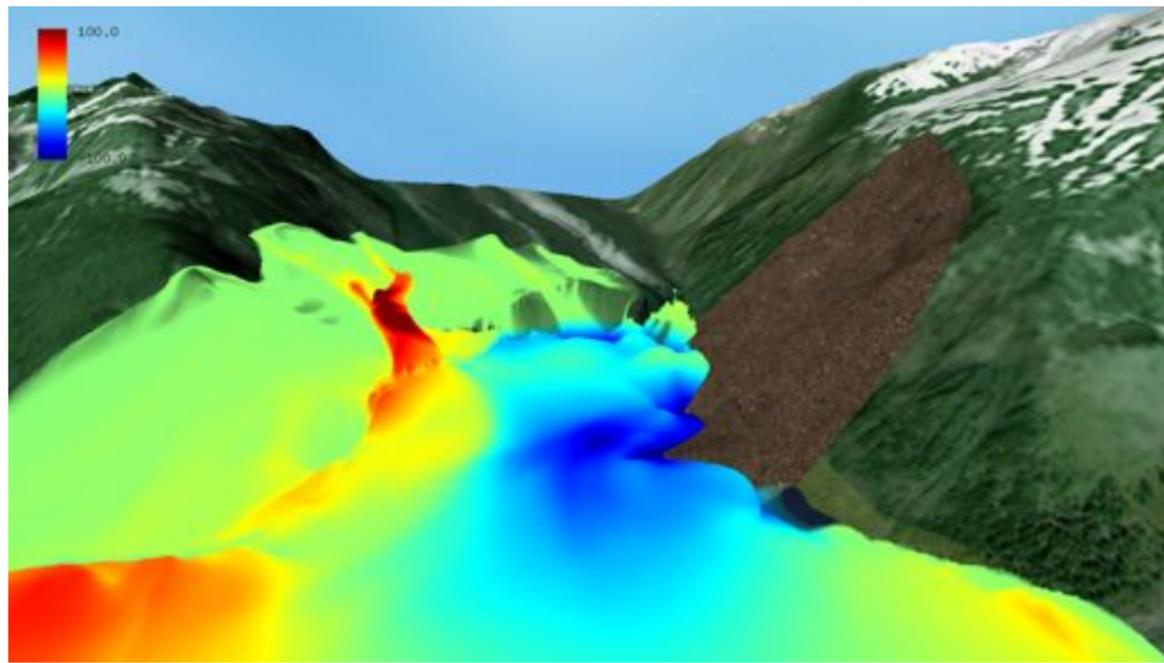
Giant wave generation



$t = 30$ s. Giant wave hits the spur southwest of Gilbert Inlet

Lituya Bay: Model results

Giant wave generation



$t = 39$ s. Maximum runup: 523.9 m

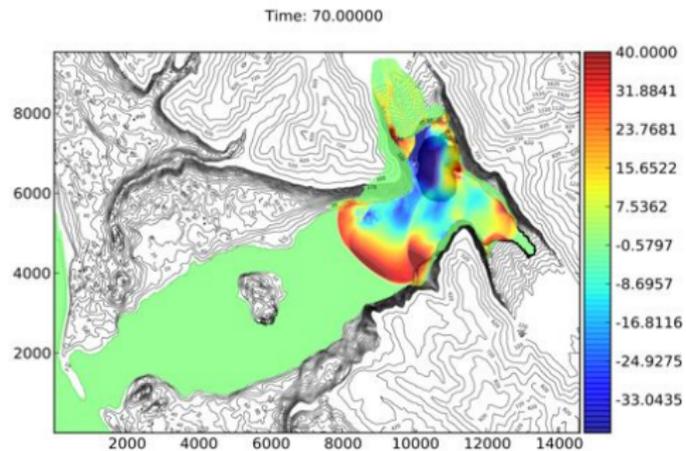
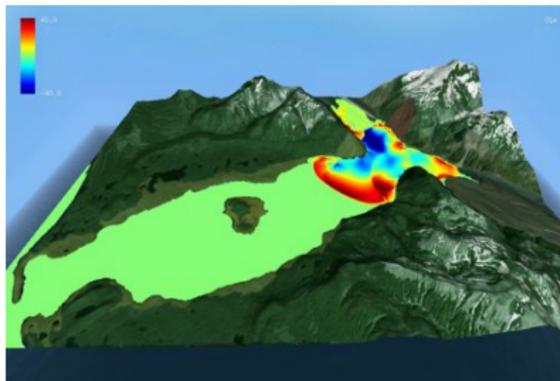
Lituya Bay: Model results

Wave evolution



Lituya Bay: Model results

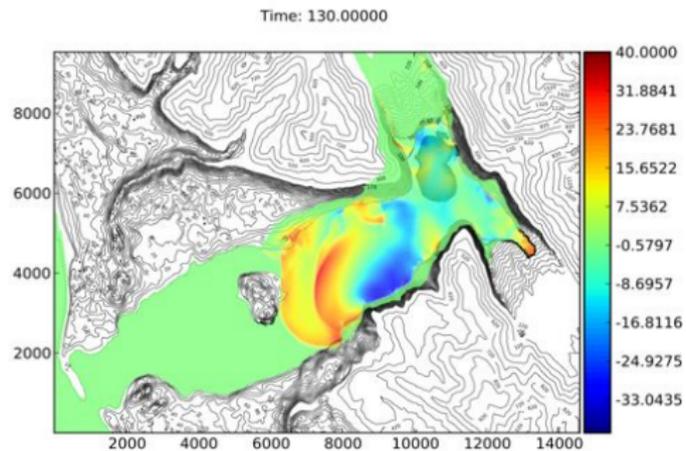
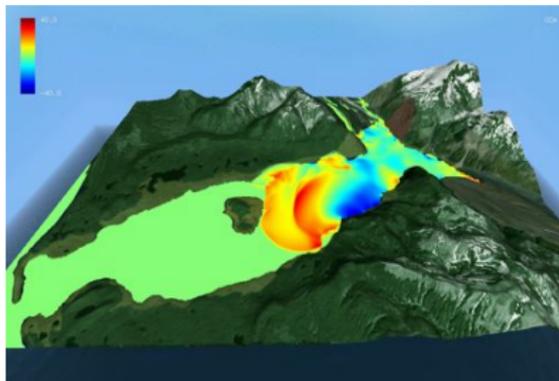
Wave evolution. Some snapshots



$t = 1 \text{ m } 10 \text{ s}$. Waves amplitude.

Lituya Bay: Model results

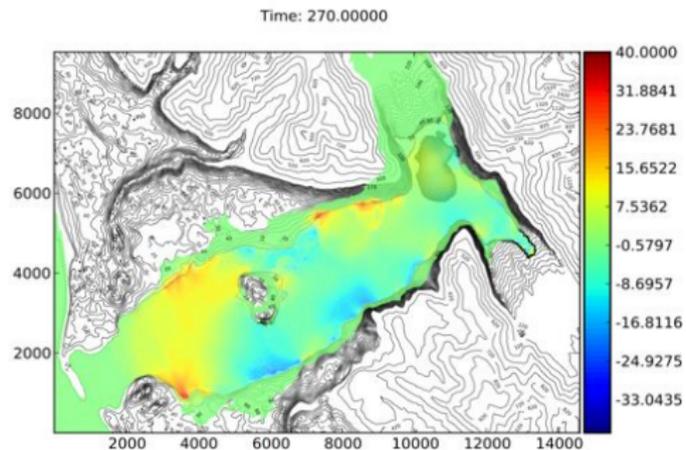
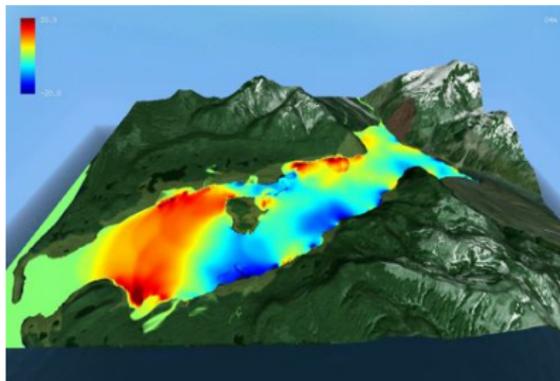
Wave evolution. Some snapshots



$t = 2 \text{ m } 10 \text{ s}$. Waves amplitude.

Lituya Bay: Model results

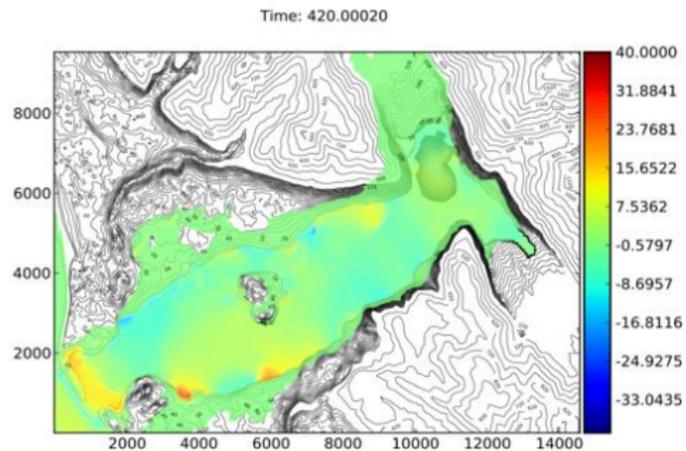
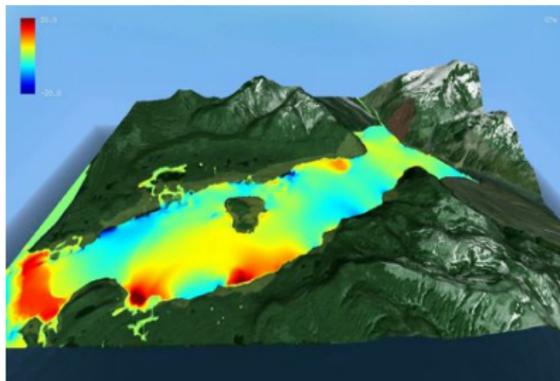
Wave evolution. Some snapshots



$t = 4 \text{ m } 30 \text{ s}$. Waves amplitude.

Lituya Bay: Model results

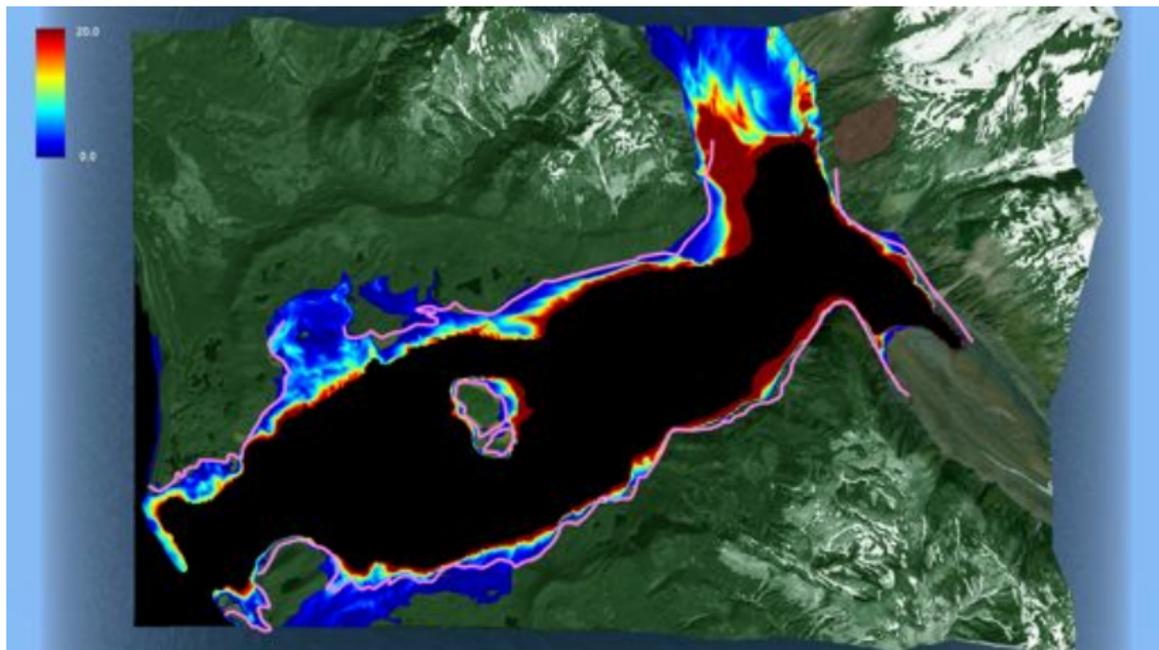
Wave evolution. Some snapshots



$t = 7$ m. Waves amplitude.

Lituya Bay: Model results

Inundation area and trimline comparison



Total inundation areas and trimline

