

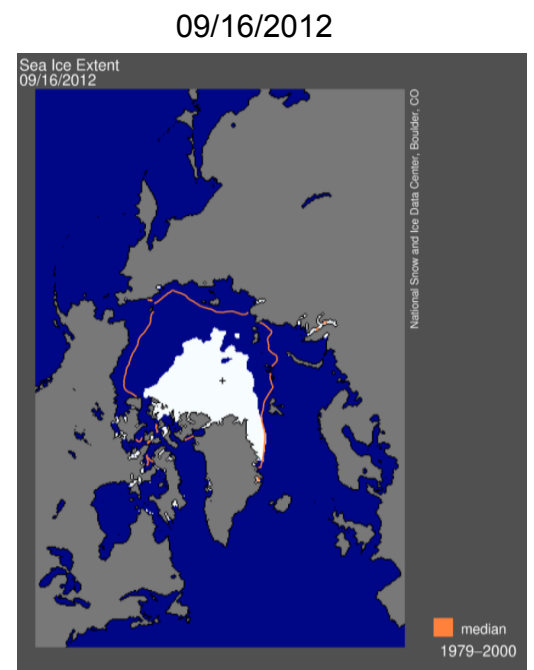
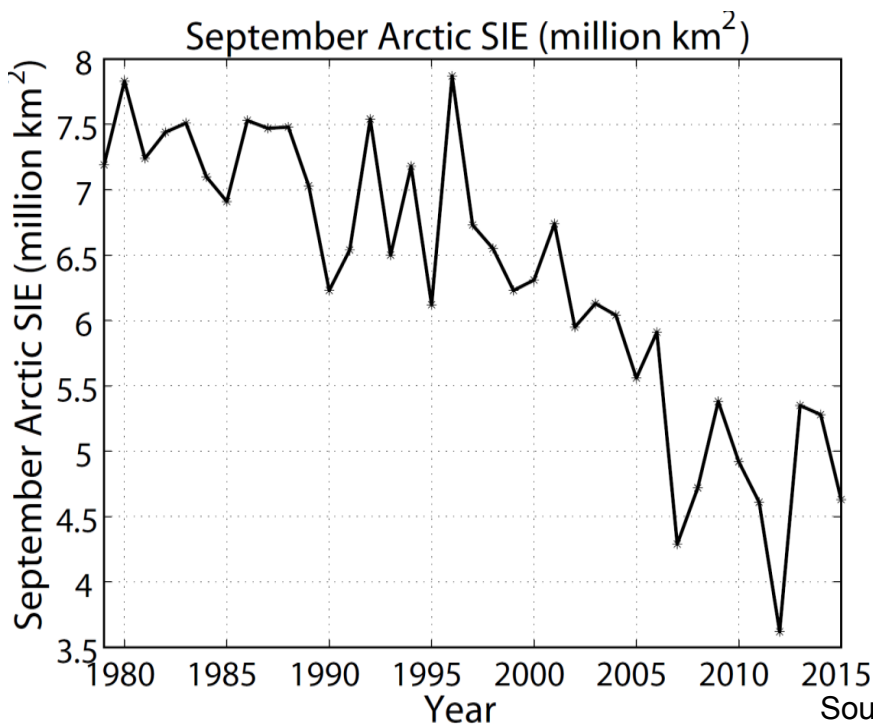
# Impact of Low Frequency Variability of the Atlantic Ocean on Arctic Sea Ice Extent

Rong Zhang

NOAA/GFDL, Princeton, NJ, USA

*CLIVAR-ICTP International Workshop on Decadal Climate Variability and Predictability  
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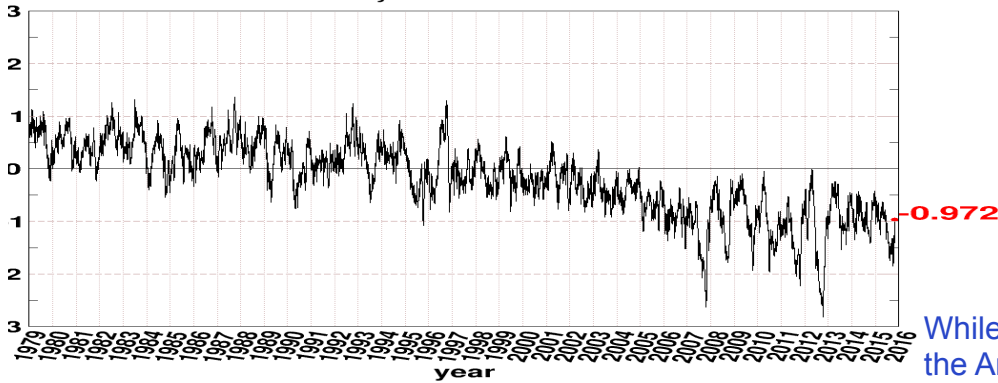
## Introduction



Satellite observations reveal a substantial decline trend in September Arctic sea ice extent since 1979, which has often been attributed in large part to the increase in greenhouse gases

### Northern Hemisphere Sea Ice Anomaly

Anomaly from 1979-2008 mean

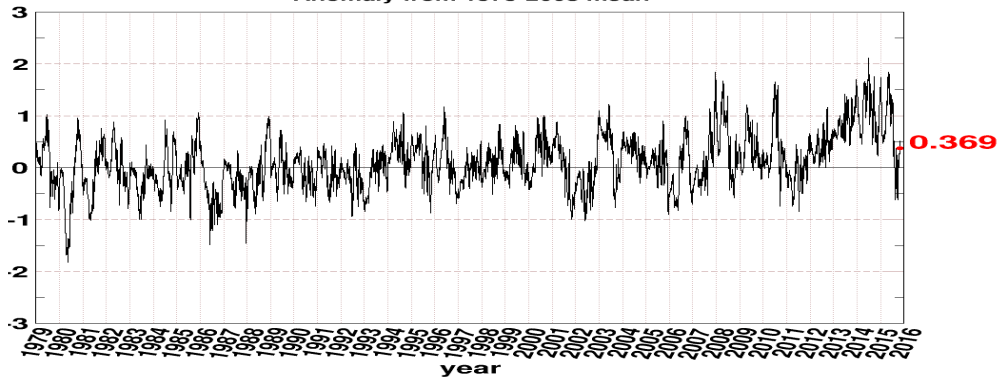


UIUC, Cryosphere Today

While the Arctic sea ice extent (SIE) dec the Antarctica sea ice extent (SIE) incre since 1979.

### Southern Hemisphere Sea Ice Anomaly

Anomaly from 1979-2008 mean



What causes the opposite trends?

### Observed Multidecadal Variations in Arctic SAT

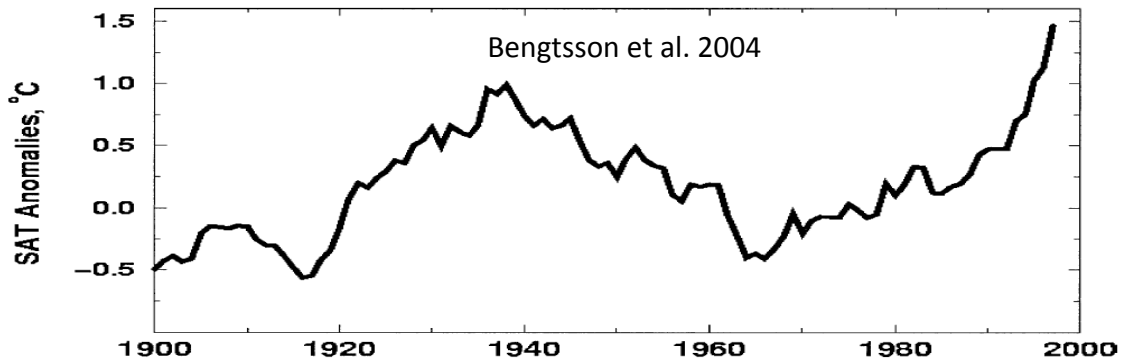
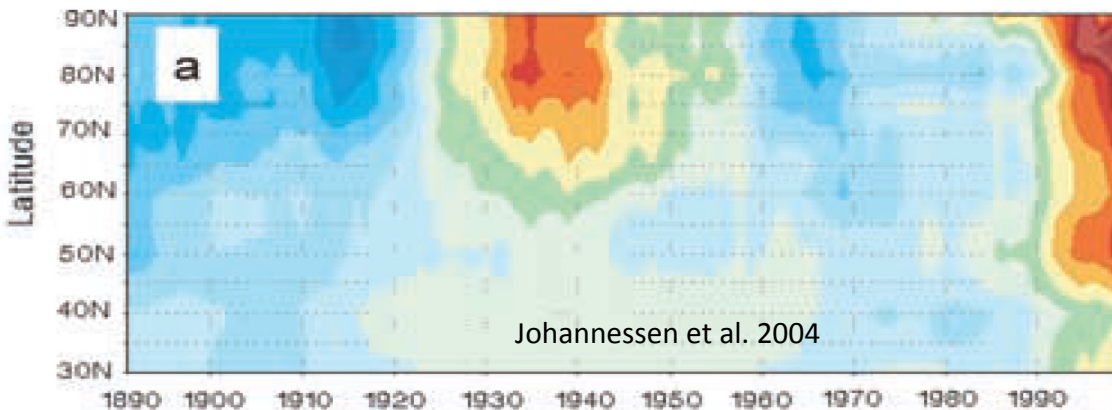
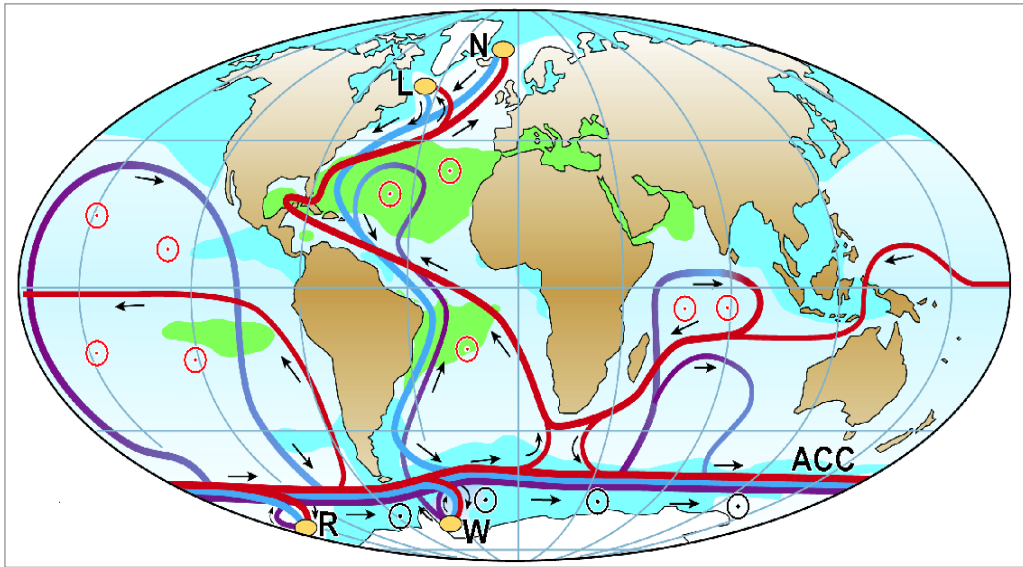


FIG. 1. Annual mean Arctic SAT anomalies (°C, area averaged from 60°–90°N) from Johannessen et al. (2004), 5-yr running mean.



# Atlantic Meridional Overturning Circulation (AMOC)

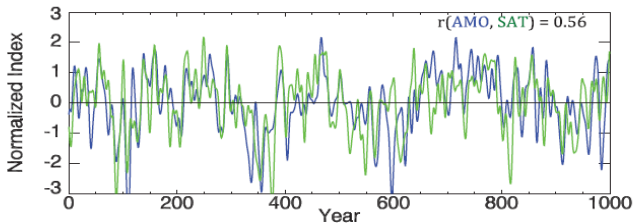


Kuklbrodt et al. 2007

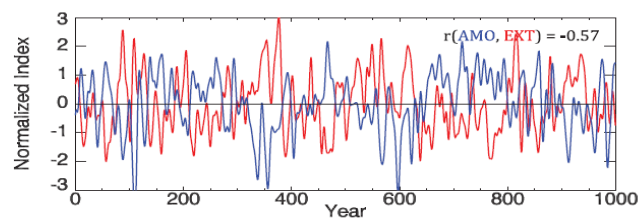
What is the role of low frequency AMOC variability in the observed Arctic sea ice decline since 1979?

## Impact of AMOC on Winter Arctic Sea Ice Variability

Time-series: AMO index and Arctic Surface Air Temperature (SAT)

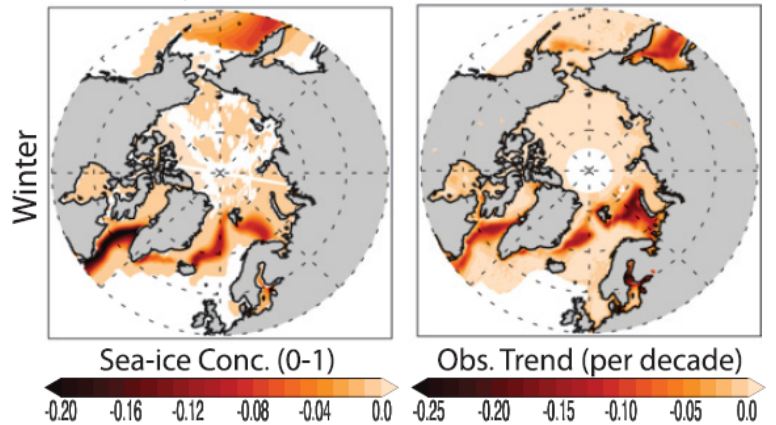


Time-series: AMO index and Arctic sea-ice extent (EXT)



GFDL CM2.1 1000-year control simulation

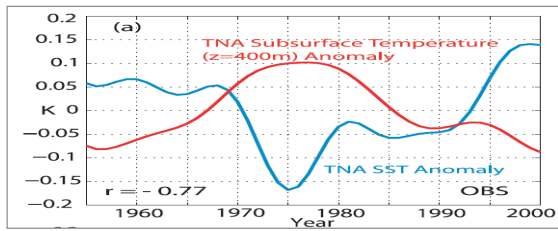
Modeled Regression on AMO Observed Trend (1979-2008)



(Mahajan, Zhang, and Delworth, 2011)

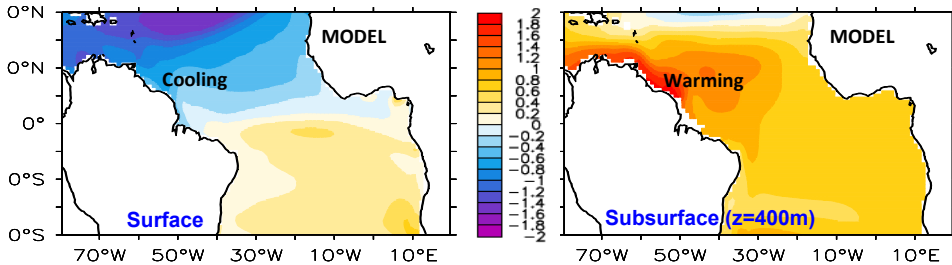
- Winter Arctic sea ice in the Atlantic side declines with an intensified AMOC
- Similar spatial patterns suggest a possible role of the AMOC in the observed sea ice decline
- The anti-correlation between AMO and winter Arctic sea ice is also found in other CMIP3 models (Day et al. 2012) and paleo records (Miles et al. 2014)

# Tropical Fingerprint of AMOC Variations



Warming  
↑  
↓  
Cooling

Zhang, 2007



The weakening of the AMOC leads to:

- A southward shift of the Intertropical Convergence Zone (ITCZ) and surface cooling
- Thermocline deepening/weakening of the western boundary current and subsurface warming in the TNA

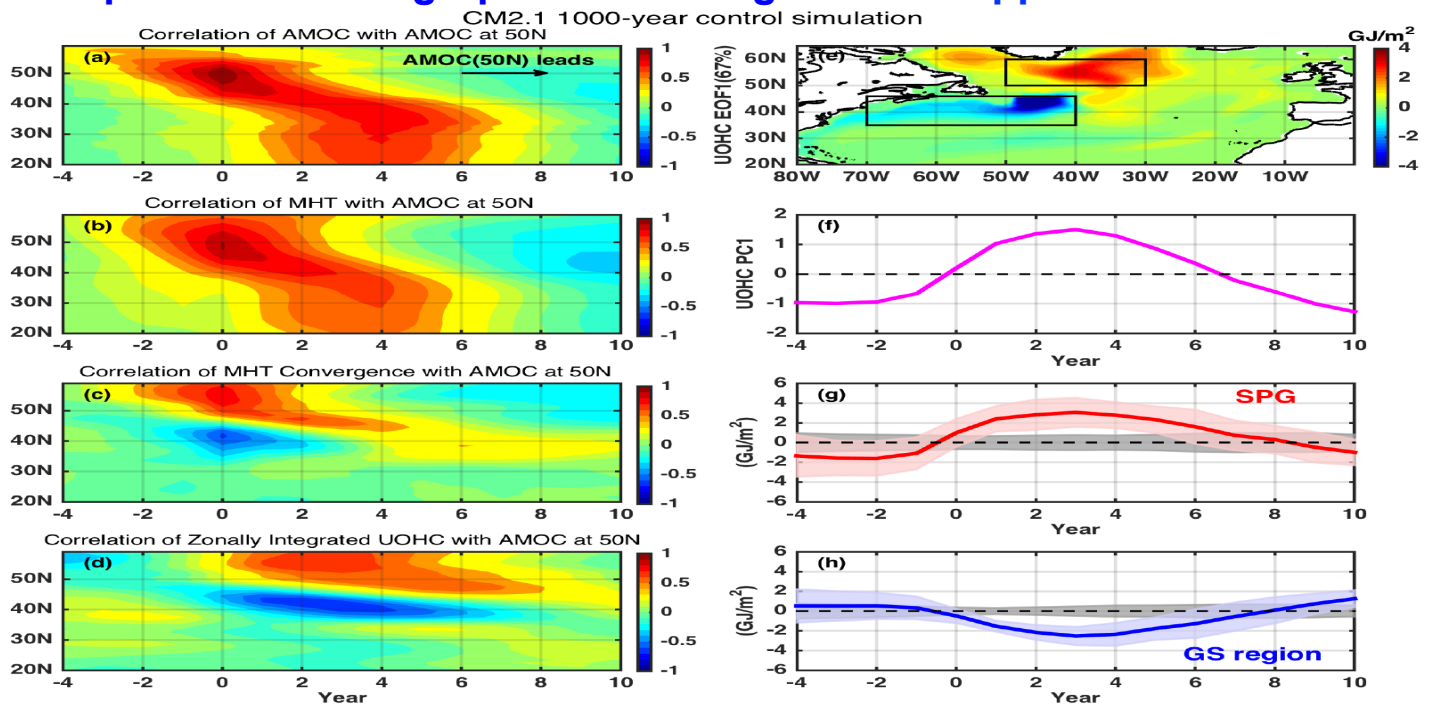
Ocean temperature anomaly due to the weakening of AMOC from GFDL CM2.1 water hosing experiment

Observed Tropical North Atlantic (TNA) SST is anti-correlated with TNA subsurface ocean temperature

The anti-correlated variations are shown to be a fingerprint of AMOC variations, suggesting the AMOC was weakened during the 70's and strengthened since then

The AMOC induced anti-correlated TNA surface and subsurface temperature variations are also found in CMIP5 models (Wang and Zhang, 2013) and paleo records (Schmidt et al. 2012, PNAS)

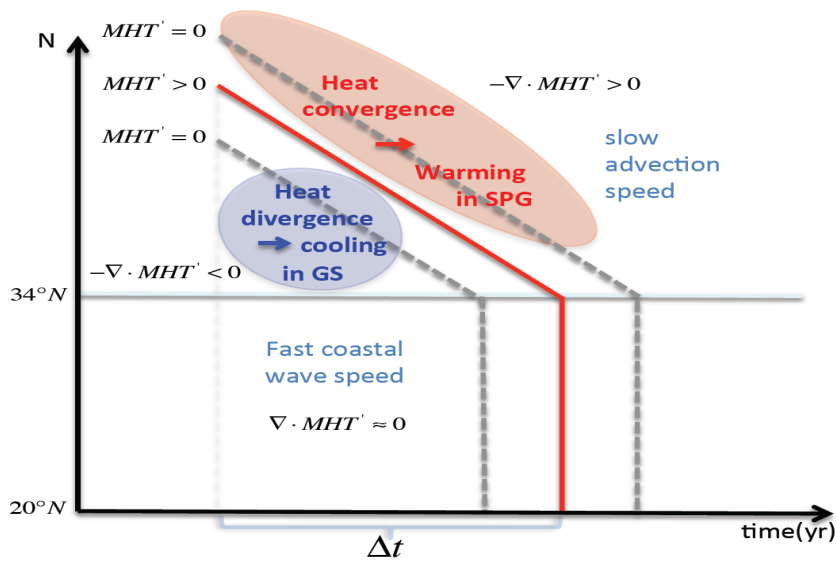
## Extra-tropical AMOC Fingerprint – Leading Mode of Upper Ocean Heat Content



Zhang and Zhang, 2015

Similar southward AMOC propagation also exists in isopycnal coordinate model GFDL CM2G (Wang et al., 2015), and high-resolution models GFDL CM2.5 (Zhang et al., 2011) and UK HiGEM (M. Thomas, personal communication, 2015)

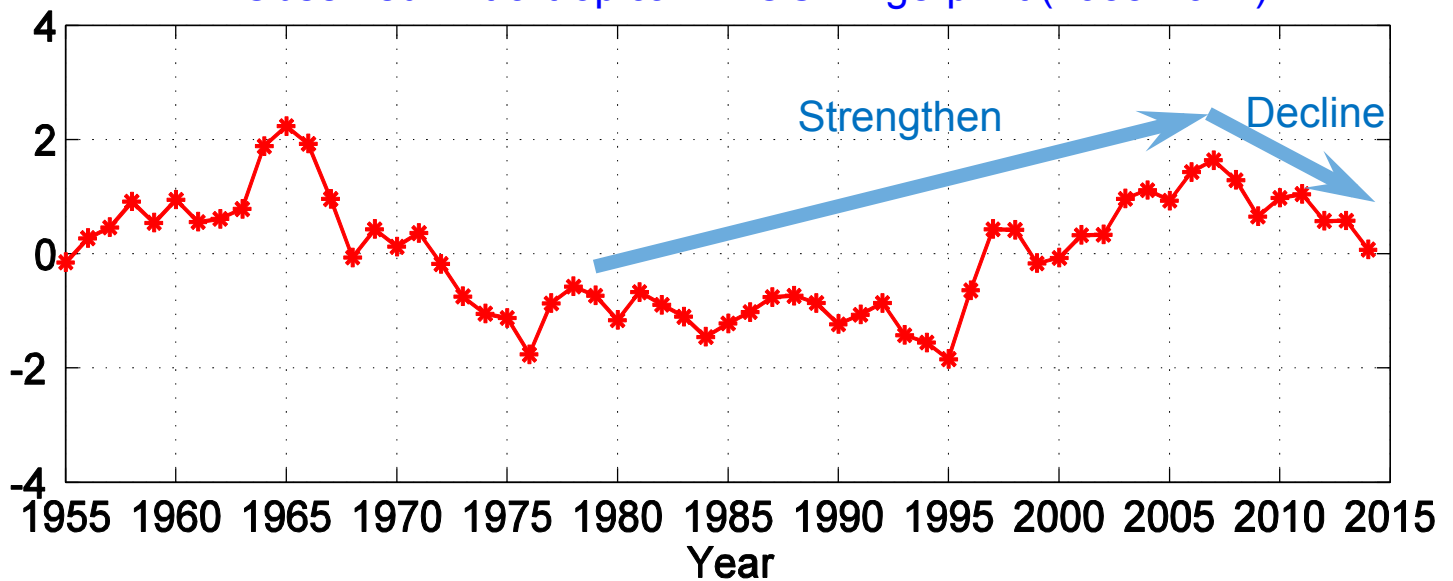
## Schematic Diagram for the Mechanism of the Evolution of the AMOC Fingerprint



Zhang and Zhang, 2015

slow propagation of the AMOC anomaly is crucial for the evolution and the enhanced decadal predictability of the AMO fingerprint, consistent with recent decadal prediction studies that successfully predicted the warm shift in the North Atlantic during the mid 1990s by initializing a stronger AMOC at northern high latitudes (Robson et al., 2012; Yeager et al., 2012; Yang 2013; Msadek et al., 2014)

## Observed Extra-tropical AMOC Fingerprint (1955-2014)





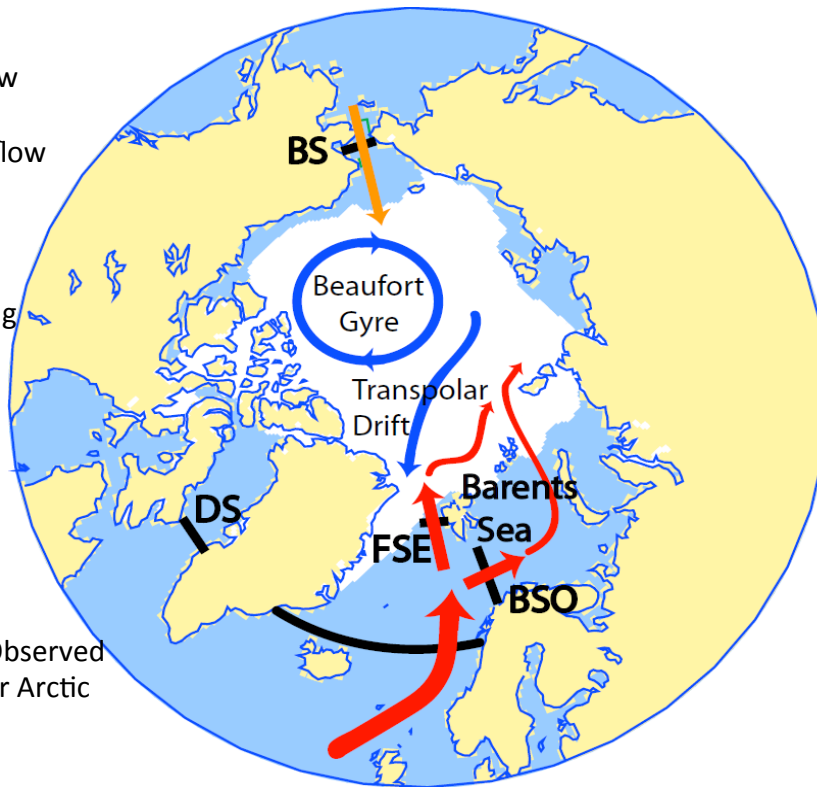
# Mechanisms for Low Frequency Variability of Summer Arctic Sea Ice Ext

Red Arrow: Atlantic Inflow

Orange Arrow: Pacific Inflow

Barents Sea Opening  
East Fram Strait  
Davis Strait  
Nares Strait

The color represents Observed  
Historical September Arctic  
Ice Extent (SIE) over 1979-2013

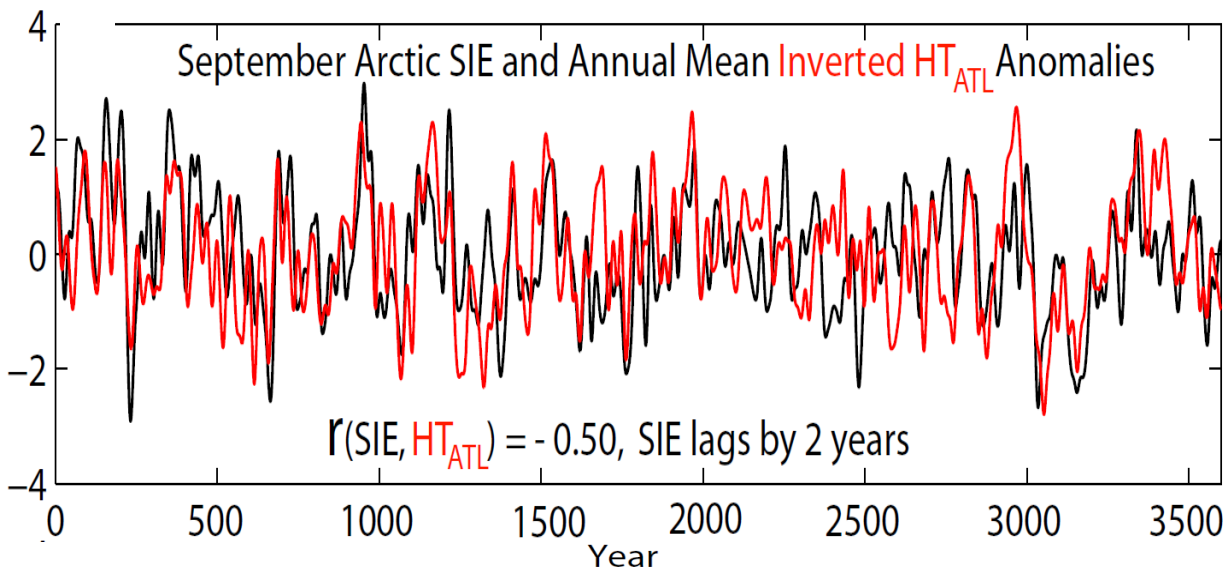


The Atlantic inflow enters the mainly through the Iceland-Scotland Ridge and further splits into two main branches: one enters the Barents Sea across the Barents Sea Opening (BSO) the other flows northward as the West Spitsbergen Current across the East Fram Strait (EFS); both eventually reach central Arctic.

Zhang, 2015, PNAS

GFDL CM2.1  
3600-year Control Simulation

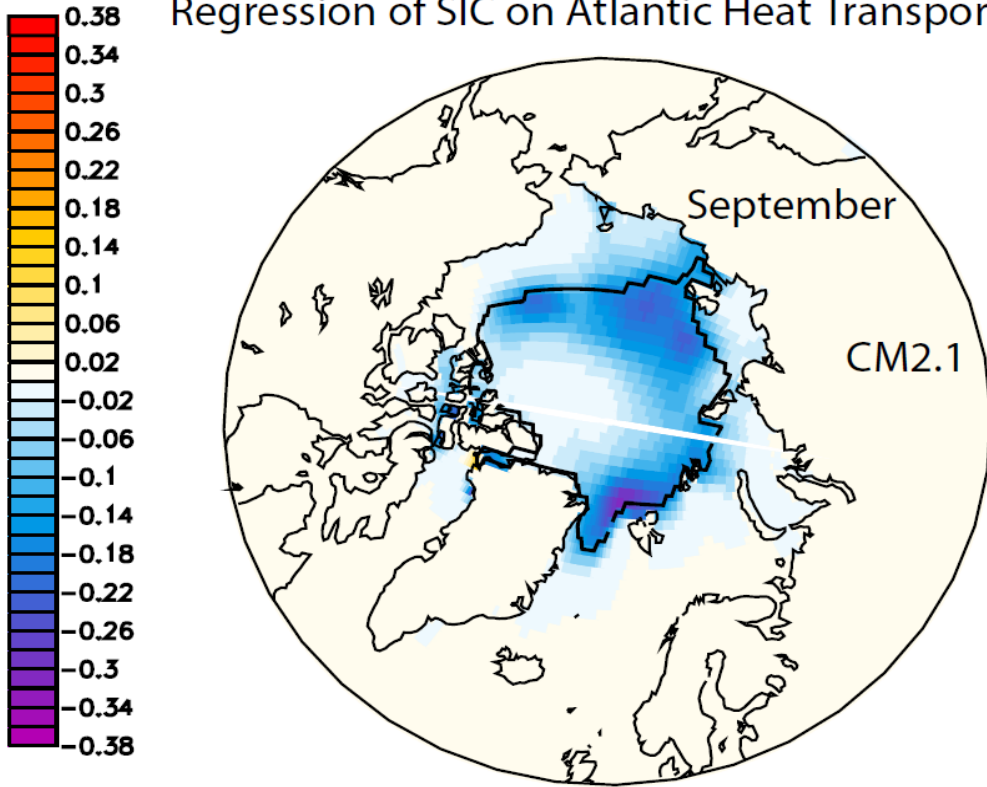
$HT_{ATL}$ : Northward Atlantic Heat Transport across the Arctic Circle



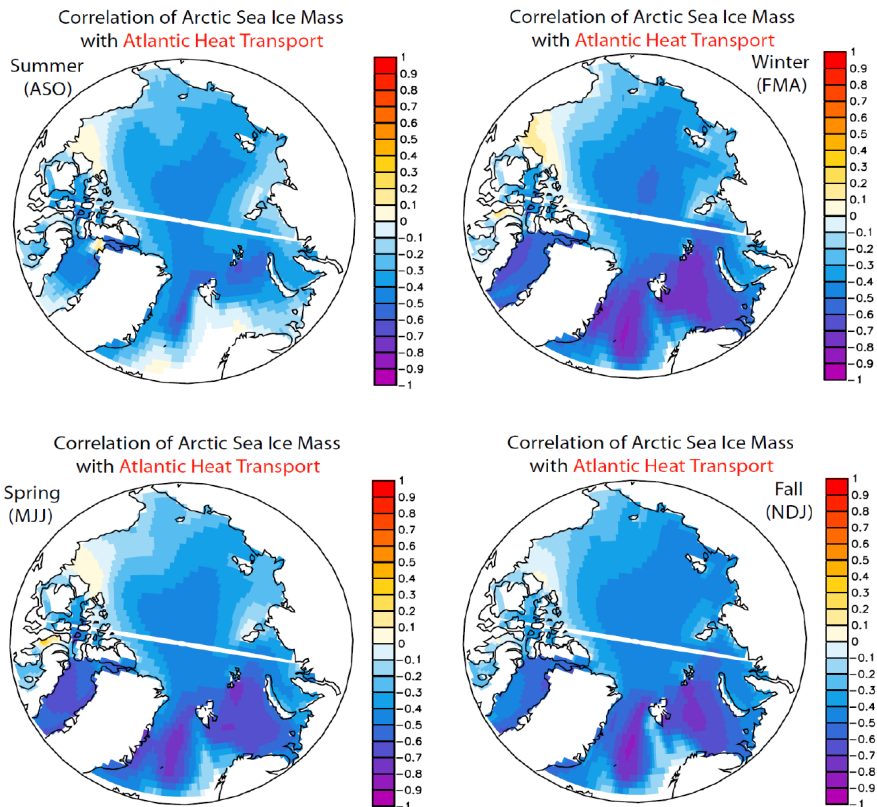
$$r(SIE, HT_{ATL}) = -0.50, \text{ SIE lags by 2 years}$$

GFDL CM2.1 3600-year Control Simulation

# Regression of SIC on Atlantic Heat Transport



## Impact of Atlantic heat transport on Arctic Sea Ice Mass at All Seasons



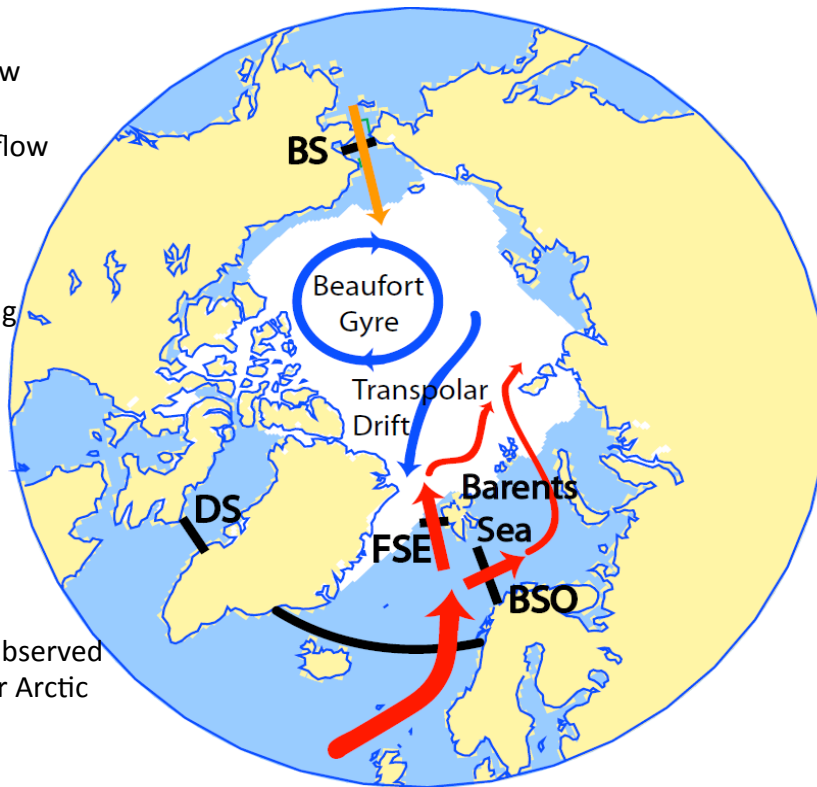
## Schematic of Main Gateways of Atlantic and Pacific Inflow Entering the Arctic and Arctic Ocean Circulation

**Arrow:** Atlantic Inflow

**Orange Arrow:** Pacific Inflow

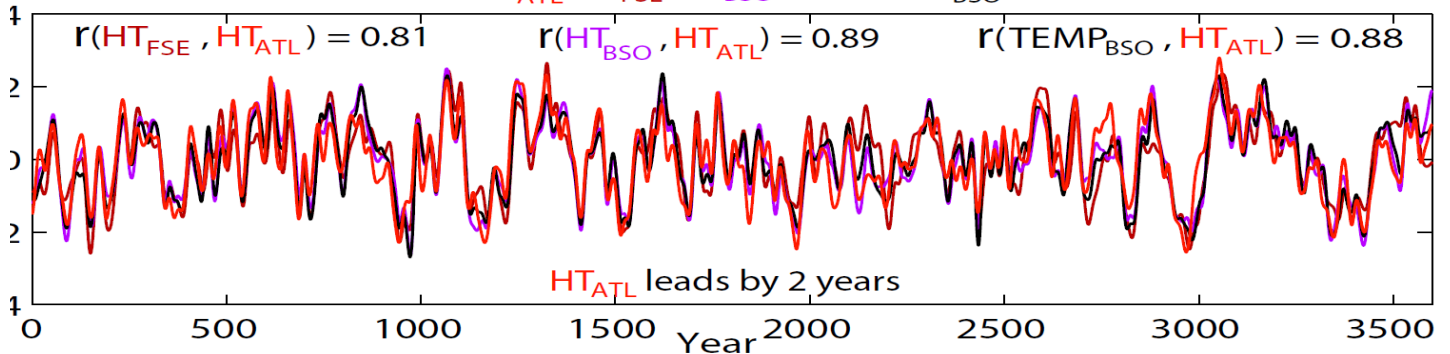
Barents Sea Opening  
East Fram Strait  
Davis Strait  
Greenland Strait

The color represents Observed  
Biological September Arctic  
Ice Extent per 1979-2013

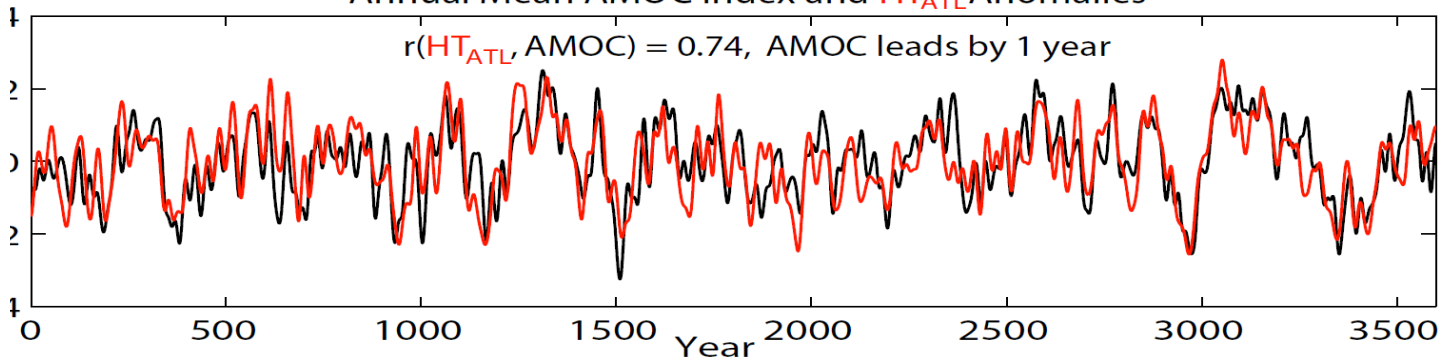


The Atlantic inflow enters the Arctic mainly through the Iceland-Scotland Ridge and further splits into two main branches: one enters the Barents Sea across the Barents Sea Opening (BSO) the other flows northward as the West Spitsbergen Current across the East Fram Strait (FSE); both eventually reach central Arctic.

Annual Mean  $HT_{ATL}$ ,  $HT_{FSE}$ ,  $HT_{BSO}$ , and  $TEMP_{BSO}$  Anomalies



Annual Mean AMOC Index and  $HT_{ATL}$  Anomalies



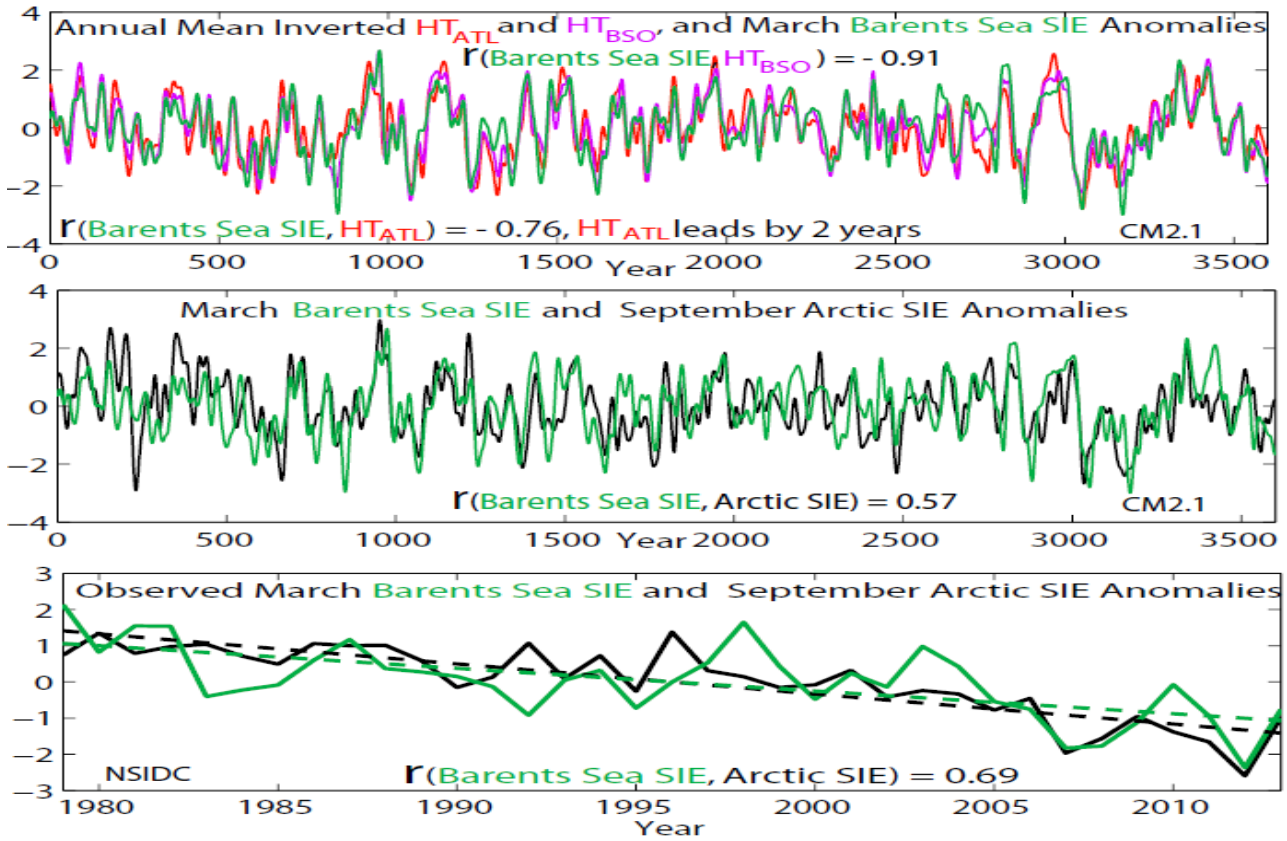
$HT_{BSO}$  : Eastward Heat Transport across BSO;  $HT_{FSE}$  : Northward Heat Transport across FSE

$TEMP_{BSO}$  : averaged Atlantic Water temperature at 200m along BSO

AMOC index : maximum of Atlantic meridional overturning streamfunction at 45°N in density space

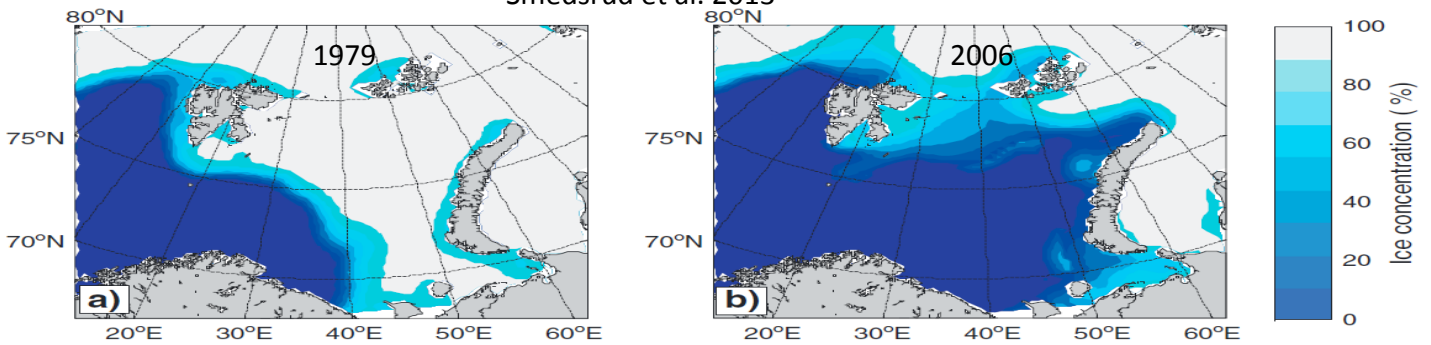


### Linkage with March Barents Sea SIE Variability

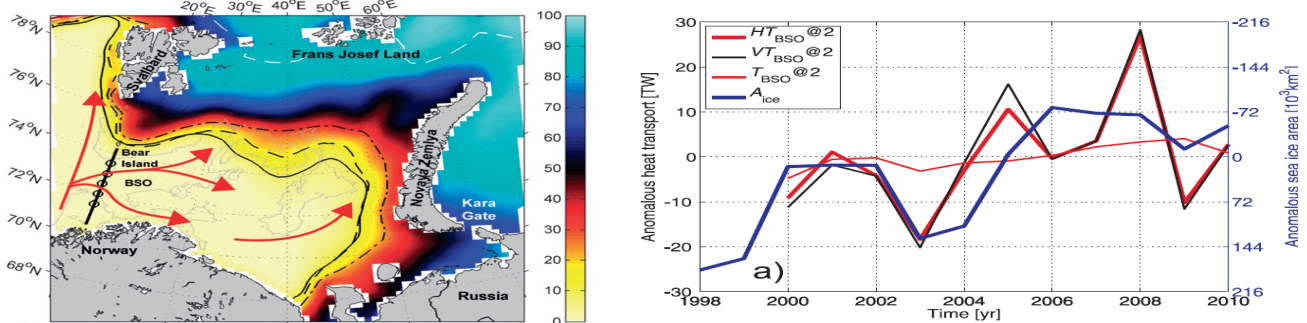


### Observed (NSIDC) Barents Sea Ice Concentration in Late Winter (March-April)

Smedsrud et al. 2013



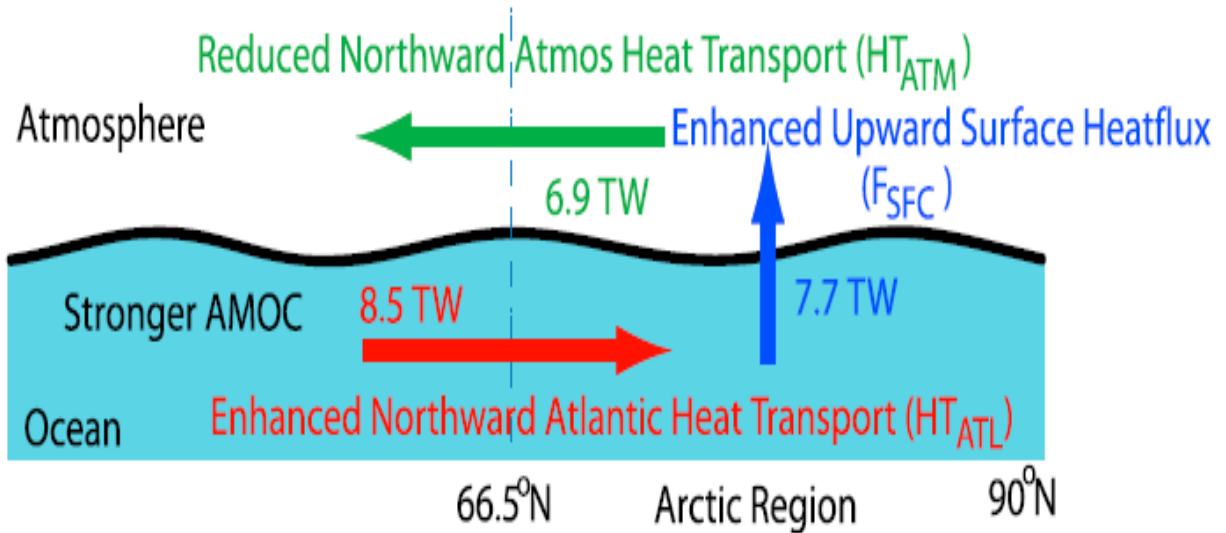
### Observed Ocean Heat Transport across Barents Sea Opening ( $HT_{BSO}$ )



Arthun et al. 2012

The observed increase in  $HT_{BSO}$  is also found as a prime driver for the observed sea ice decline in Barents

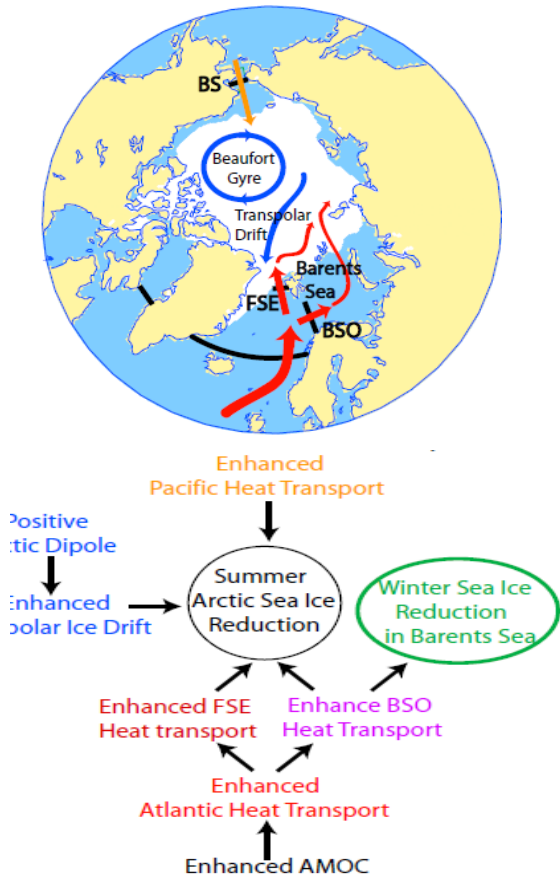
## Response of Atmosphere Heat Transport



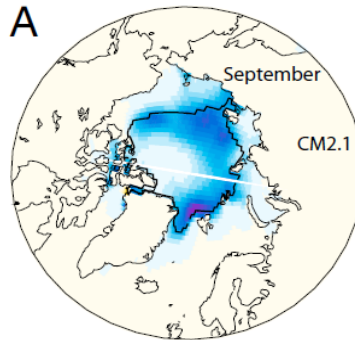
The Bjerknes compensation (Bjerknes, 1964) has been found at decadal time scale (Shaffrey & Sutton, 2010; Jungclaus & Koenig, 2010; Farneti and Vallis, 2013)

At multidecadal/centennial time scale, the coherences among  $HT_{ATL}$ , Arctic SHF, and inverted  $HT_{ATM}$  are higher than those at decadal time scale

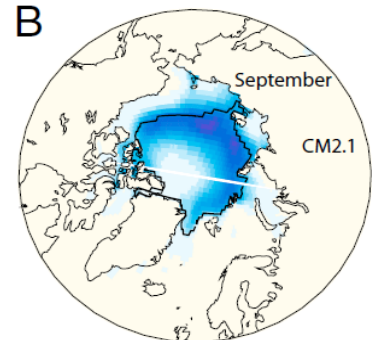
Changes in  $HT_{ATM}$  are forced by anti-correlated changes in  $HT_{ATL}$  thus provide a negative feedback to September Arctic SIE variations



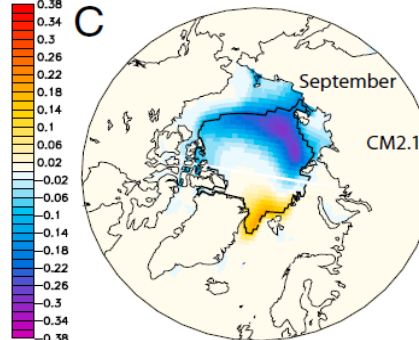
Regression of SIC on Atlantic Heat Transport



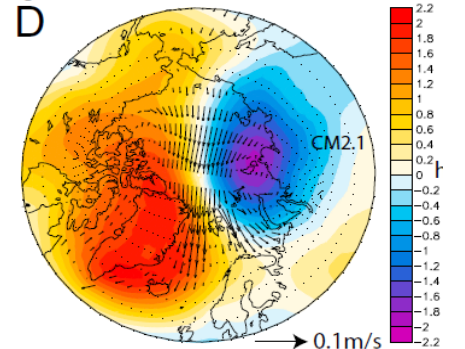
Regression of SIC on Pacific Heat Transport



Regression of SIC on Arctic Dipole



Regression of AMJJ SLP/Ice Motion on Arctic Dipole



## Summary and Discussions

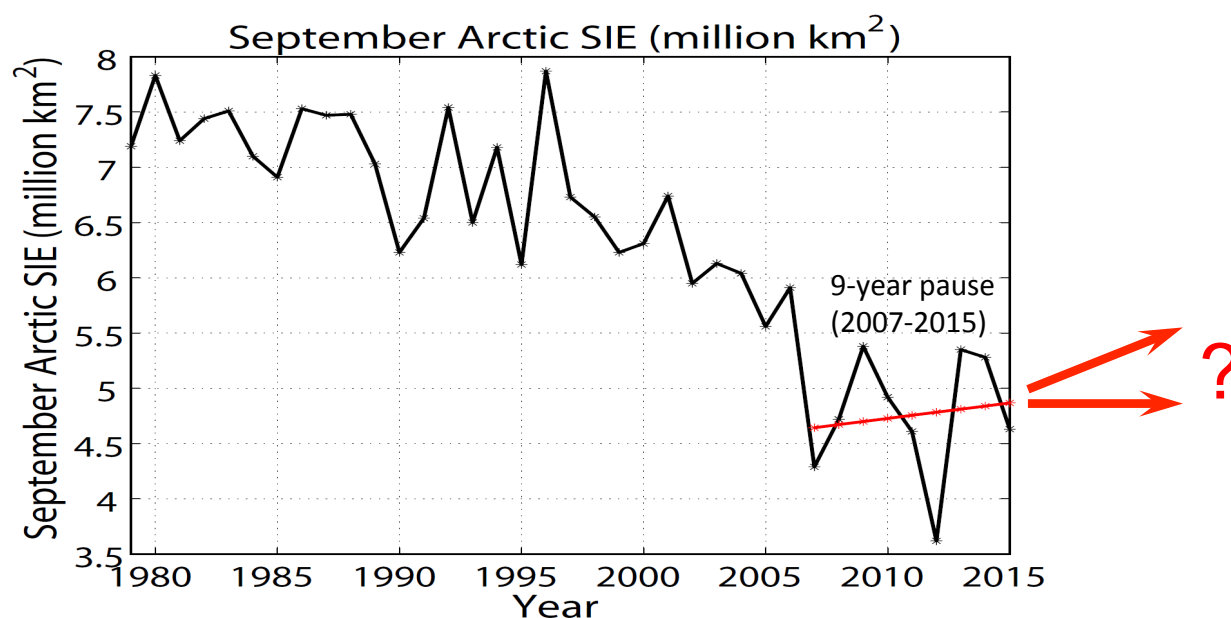
The AMOC variability and the associated Atlantic heat transport into the Arctic have played a significant role in the low frequency variability of summer Arctic SIE

Summer Arctic SIE variations are significantly correlated with winter SIE variations in Barents Sea in both modeled results and observations, indicating the importance of the Atlantic heat transport into the Arctic

AMOC fingerprints indicate a strengthening of AMOC since mid 70's, consistent with the observed decline of Arctic sea ice

At low frequency, changes in atmosphere heat transport into the Arctic are forced by anti-correlated changes in the Atlantic heat transport into the Arctic, they provide a negative feedback to changes in summer Arctic SIE

Enhanced Pacific heat transport into the Arctic and Positive Arctic Dipole also contribute to summer Arctic sea ice decline



Very recent study identified a 7-year pause (2007-2013) in summer Arctic sea ice decline (Swart et al. 2015). If the AMOC continues to weaken in the near future, there might be a longer *hiatus* in the September Arctic SIE decline